#### **APT REANALYSIS**

- Palm Springs, Forest Hill Village Wellfield SITE: Section *M.* Township 445 Range 42E
- CH2M-Hill, Hydrogeologic Report, Evaluation of Wellfield Facilities, **REPORT:** Village of Palm Springs, Palm Beach County, Florida, June 1983.

# **GEOLOGIC DATA: pp. 5-10**

Drilling logs for pilot hole to 260' below grade showing: Well  $#5 -$ 



Geophysical, electric and gamma ray logs to 260' below grade

Wells #1-4 - gamma logs

Based on drilling log and gamma logs, aquifer thickness estimated at 234'.

The lithologic and geophysical logs show a good producing zone, probably the Turnpike aquifer, from 120' to 165'BG.

Site elevation is approximately \_\_ NGVD.





Depth to water: 5'

# **INFLUENCING FACTORS:**

 $1)$ Lake Worth Drainage District canals E-3 and E-8, 350' and 600' from the pumping well, respectively. Both canals are about 6' deep.

APT:

Started: 4/13/83 at 0917 Duration: 48 hours Discharge: 1600 GPM to E-3 (308,021 FT3/DAY) Recovery:

Comments:

- Stevens recorders on wells 2, 3, and 4. Chalked steel tape on wells 1, 5, and  $1)$ canal.
- $2)$ Variation in Q at  $t = 8$  hours, "quickly" corrected.
- At 40 hours a 6" irrigation well located 3,000' from well #5 was pumped for 3) three hours at about 300 GPM.

# **CONSULTANT'S ANALYSIS:**

Method: Jacob - distance drawdown, time drawdown

Comments:

 $\Rightarrow$ 

- Pg. 6-7 Consultant corrected drawdown data for declining water level trend  $1)$ based on eight hours of background data. The decline was obvious only for the last two hours, 0600-0800. The decline was attributed to evapotranspiration and barometric pressure change. The arguments are not convincing.
- $(2)$ Poor results for well #4.
- 3) Jacob analysis gave higher T's.
- 4). Analysis of recharge from canals made using Walton's method. This analysis seems inappropriate given that the method assumes a fully penetrating boundary and the canals penetrate only 3% of the aquifer.

# **REANALYSIS:**

Method: Neuman, 1975, Analysis of Pumping Test Data from Anisotropic Unconfined Aquifers Considering Delayed Gravity Response.

# Results:



# Comments:

- $1)$ Drawdown data was used as measured
- $2$ <sup>\*</sup> Wells 2 and 3 had good type curve matches, fairly smooth drawdown curves and similar results. The drawdown curve for well 1 was more erratic, the type curve match was not as good, and the results did not agree well with results from wells 2 and 3. Since well 1 was measured with chalked tape while wells 2 and 3 were measured with Stevens recorders, it is possible that the timedrawdown measurements were not as accurate in well 1. Therefore, the results from well 1 were not used in calculating the average aquifer characteristics.
- $3)$ There was not sufficient late time data to calculate specific yield at these sites.
- Using the assumed thickness of the aquifer, 230', and the average T, 374,000 4) Building the assumed uncluded on the equilibrium,  $\frac{1}{2}$ ,  $\frac{1}{$ completely unrealistic for the sand and unsolutioned rock sections.
- Given the high K computed from this method and the likely presence of the 5) Turnpike aquifer, the Neuman method assuming a homogeneous unconfined system is inappropriate.

Method: Modified Hantush

Results:



Comments:

- There are two possibilities for semi-confined aquifer behavior at the site. First,  $\left( \begin{matrix} 1 \end{matrix} \right)$ if the Turnpike aquifer at the site is sufficiently more permeable than the nonsolutioned zones above and below it, it would act as a semi-confined aquifer. Second, the sandstone/clay layer from 33-46 ft. at the site could act as a semiconfining layer for the aquifer below it. In either case, a semi-confined analysis is appropriate.
- $2)$ as 1) from above.
- $3)$ <sup>\*</sup> as 2) from above.
- The modified Hantush method assumes: 1) there is negligible drawdown in 4) the source bed above the semi-confined production zone and 2) water release from storage in the semi-confining layer is appreciable. Since thre is no data on the upper aquifer zones during the pump test, it is not possible to check these assumptions.
- 5) If the effective producing zone at the site is the Turnpike aguifer with a thickness of 45', the hydraulic conductivity of the zong is 3,840 FT/DAY.

based on the above results

If the effective producing zone is the aquifer below the sandstone/clay layer, with a thickness of 188', the hydraulic conductivity of the zone is 920 FT/DAY.  $6)$ 



Results:



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line.

Comments:<br>1) through 3) as above.

 $4)$ The Hantush-Jacob method assumes:

- $5)$ as above except  $K = 4,470$  FT/DAY.
- $6)$ as above except  $K = 1,070$  FT/DAY.

Method: Numerical Analysis using radial finite element method.

# **RECOMMENDED VALUES:**

Comments:

REFERENCES:

Water Use Permit 50-00036, Staff Report

Palm Springs Pump Test Reanalysis. Constant Flux Nodes - Actual Prod. Well Partially Penetrating Well from  $140 - 170$  Nodes  $5, 6, 7$  $Q = 1600$  GPM = 214 FT<sup>3</sup>/MIN = 308,000 FT<sup>3</sup>/DAY Axisymetric  $\div$  Q by 2T  $Q_A = 49,048$  FT<sup>3</sup>/DAY Nodes 517 get half shares because at top & bottom of well Node  $5 = 7 = 12,262$  FT<sup>3</sup>/DAY Node 6 =  $24,524$  FT3/DAY Hydraulic Conductivity-From PB135, Hantush-Jacob analysis,  $\overline{T}$ = 156,000 FT<sup>2</sup>/DAY<br>Assuming aguiter thickness of 235',  $\overline{K}$ =  $\overline{T}/b$ = 662 FT/DAY  $(s$ ased on litho logs, model setup) Rough cot since used entire thickness I analysis seminantined, using  $k = 600 FT/bAy$   $K_H/K_v = 10$  (assumed, no basis) Drawdowns seem quite small. Double checking model by running fully penetrating well to compare to Theis  $Q = 308,000$  FT<sup>3</sup>/DAY  $\overline{b}$  = 235' Q/5 = 1311 FT<sup>2</sup>/ day<br>Q/62TT = 209 FT<sup>2</sup>/ day Node  $F + c<sub>ov</sub>$ .  $^{\prime}$  Q 2090  $10$  $\mathbf{I}$  $20<sub>o</sub>$  $4180$  $\overline{z}$ 3  $17.5$ 3658  $\overline{\mathcal{L}}$  $15$ 3135  $\mathcal{S}$  $15$ 3135  $15$ 6 3135  $\vec{\tau}$  $12.5$  $2613$  $\overline{8}$ 10  $2090$  $\widetilde{\mathcal{T}}$  $12.5$  $2013$ 10  $5$ 3135  $15$  $\left| \right|$  $3135$  $15$  $12$  $3135$  $13$  $5<sup>5</sup>$  $3135$  $14$  $15$  $3135$  $15$  $12.5$  $2613$  $12.5$ 16  $2613$  $17$ 7.5 1568

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<u>Palm Springs APT Reanalysis w/ MODFLOW</u> A 3-dimensional, model was chosen to reflect the aguiter layering and the partially penetrating pumping well. The model was given 16 layers from 123-133, and 140-170 feet BG, was represented as open to the agoifer from 120 to 170 feet in the model. This makes nodes 9,10,11<br>and 12 pumping nodes. The model represented 1/4 of the aguifer<br>area with the well in I corner and no fluw boundaries on all sides. Dischage: APT Q = 1600 GPM  $D = 4$  for  $4$  area in model =  $400$  GPM = 77000  $FT^3/DAY$ 2) Divide proportionally among pumping nodes Nodes 9 & 10 10' thick Nodes 11612 15 thick  $77000/50 = 1540 FT^3/DRY-FT$  $550'$ Nodes 9 = 10 10' (1540) = 15,400 FT3/DAY<br>Nodes 11 = 15' (1540) = 23,100 FT3/DAY Time: 2 Days There is a continuous, record of water levels Supporting  $Data:$ at a monitoring well in the Palm Springs. wellfield. The data reflect primarily when the pumps are on & off. Data is in strip chart form and data reduction would be extremely tedious, Not worth the effort. Monitor well is completed in the production zone. No shallow monitoring wells, Aguifer Characteristics: Neuman analysis of Palm Springs Obs Wells 1, 2, and 3. showed the following: We11  $\leq$  $\tau$  $K_{\mathbf{p}}$  $(fT^2/PPAY)$ .0038  $\mathbf{L}$  $166,800$  $.013$ 375200  $.022$  $\mathbf{z}$ ,0006 3  $372,764$ ,0086. ,0002

ATORIA (1988) 50 SHEETS 5 SQUARE<br>100 SHEETS 5 SQUARE<br>ATORIAL (1999) 200 SHEETS 5 SQUARE

 $\sqrt{2}$ 

PS APT Rean (cont.)  $\mathsf{Z}% _{\mathsf{M}}^{\mathsf{N}}(\mathcal{M}_{0})\simeq\mathsf{Z}_{\mathsf{M}}^{\mathsf{N}}(\mathcal{M}_{0})$ The lithologic and geophysical logs show a good producing zone Results from Well I were considered quastionable (see APT reanalysis<br>sheet), Average characteristics based on wells 2 and 3 are:  $T = 374,000 F T^2/DAY$  $S = .015$  $K_D = .0004$ Using the assumed thickness of the aguiter,  $230'$ ,  $K = 1626$  FT/DAY,<br>which is clearly ridiculous. Modified Hantush Method  $well$  $\leq$  $\beta$  $\tau$  $204,366$  $\mathcal{L}$ ,000023  $.02$  $245,240$  $,00020$  $\mathbf{z}$  $.02$  $176,570$ 3  $.00012$  $.02$  $A_{\nu g}$  $210,900$  $.0001b$  $02 -$ (estimated Turnpike thickness at site) Assume b of producing zone is 45% Then K of producing zone is  $T/b$  = 4690 FT/DAY. Assume b of producing zone is 188' (thickness of aguifer<br>below sandstone/clay layer). Then K of producing zone is  $T/b = 1120$   $F\dot{T}/DA\dot{Y}$ .

 $\sum_{\substack{n=-\infty\\n\neq n\text{ odd}}}$  42.381 000 SHEETS 5 SQUARE



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Table 2-1<br>SUMMARY OF WELL DATA

			Main Well Field									Forest Hill Well Field				
		Well Field No. 1			Well Field No. 2				Well Field No. 3 $Wc11$ No.							
													$\mathbf{2}$	3		$5^{\rm b}$
	Parameters		$3^{\mathbf{a}}$		5	6	7	8	9	10	11					
Construction Date		1957	1964	1967	1969	1969	1971	1974	1977	1977	1977	1959	1959	1959	1959)	1983
Total Depth, ft		150	222	150	222	205	200	200	210	210	210	135	141	135	93	170
Casing:	Diameter, in Depth, ft	8 140	10 182	10 110	12 183	12 165	12 161	12 160	12 104	12 102	12 104	8 100	8 113	8 109	8 90	14 115
Screen:	Material Size, Slot Depth, ft	None $\equiv$	Everdur -- 182	Everdur 40 110	Everdur $\qquad \qquad \cdots$ 182	Stainless -- 165	Stainless 35 161	Stainless 40 160	-- 170	Open Hole Open Hole 170	Stainless 40 170	Unk Unk $100 - 135$	Unk Unk $113 - 141$	Unk Unk 109-135	Unk Unk $\sim$ $-$	Stainless 80
Pump:	Manufacturer Model	Peeriess <b>BMA</b>	Peerless 10L8	Peerless 10MA	Deming 15M8E1	Deming 20MBE1	Johnston GD3620	Courbin 20M8E1	Johnston 10DS	Johnston 10DS	Johnston 10DS	None $\qquad \qquad -$	None $\hspace{0.05cm}$ $\hspace{0.05cm}$	None $\frac{1}{2}$	None	None
Well Yield, gpm		400	400	500	500	500	500	600	700	700	700	500	500	400	--	1,400
Static Water Level/ Date		$8.4/1 - 83$	$10.0/1 - 83$ 7.3/1-83		$14.5/1 - 83$	$5.3/1 - 83$	$7.2/1 - 83$	$6.2/1 - 83$		$8.0/1-83^{\circ}$ 10.3/1-83 8.7/1-83			$5.35/1-83$ 6.92/12-82	$8.65/12 - 82$	--	$5.05/4 - 83$
Pumped Water Level/ Date			$27.2/1 - 83$ 40.0/1-83	$34.6/1 - 83$	$41.3/1 - 83$		$22.7/1 - 83$ $42.1/1 - 83$						28.2/1-83 10.6/1-83 17.9/1-83 18.2/1-83 14.15/1-83 10.67/12-82 11.58/12-82			$-- 16.96/4-83$
Maximum Drawdown, ft		18.8	30.0	27.3	26.8	17.4	34.9	22.1	2.6	7.6	9.5	8.80	3.75	2.93		11.91
Flow Rate, gpm		325	360	480	675	585	75	715	700	700	700	620	500	400	$\sim$ $\sim$	1,600
Specific Capacity, gpm/ft		17	12	18 (20) <sup>d</sup>	$25(16)^d$	$34(42)^d$	$\mathbf{2}$	$32(42)^{d}$	269	92 $(110)^d$	74	70	133	137		134

 $a_{\text{Well No. 2, driiled in 1964, was abandoned in 1974, used as monitoring well with recorder 1980 to 1983.}$ 

 $b$ <sub>Well</sub> No. 5 is recently constructed Test Production Well, TPW-1.

 $\textdegree$ Static water level above the base of below-grade pump pit.

d<sub>Number</sub> in parenthesis is original specific capacity given where data available.

 $e$ Well logged to 93 feet, however, postulated to be deeper (see Figure 5-4). This well will be abandoned.

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The Forest Hill WF, located approximately 2 miles west of WF No. 3, clearly develops water from the Turnpike aquifer. This highly permeable section of the Anastasia Formation probably extends from WF No. 3 westward through the Forest Hill Village site, terminating in the vicinity of the Florida Turnpike.

# Aquifer Performance Test

During the rehabilitation of Forest Hill Village Wells No. 1, 2, and 3, it became clear that the well yields at this site were quite high (70 to 137 gpm/ft). Following the rehabilitation of these existing wells, a new well was constructed to complete the well field facility. This well, identified as Well No. 5, was used as the pumping well during a 48-hour APT conducted April 13 to April 15, 1983.

In order to conduct the test, a 12-inch vertical turbine pump with diesel engine was installed in Well No. 5. A total of 350 feet of 10-inch PVC pipe was laid from the well eastward to LWDD Canal E-3 (see Figure 2-2).

A Stevens Type F continuous water level recorder was installed at Well No. 2 one day prior to the planned start of the test to collect background water level data. However, the pen malfunctioned and no record was produced. On April 12, 1983, Stevens recorders were installed at t Wells No. 2, 3, and 4 and gear ratios were set to run o 4 hours at full time scale using a 1:1 gauge scale ratio.<br> $C_2$  Static water levels in Wells No. 1 through 5 were measured, as was the "static" water level in Canal E-3 adjacent to Well No. 4. The test was officially started at 1446 hours  $ss$ at a withdrawal rate of 1,800 gpm. Approximately 30 minutes  $f_{q1}$ into the test, a "familiar" pattern was observed on the water level recorder charts. This pattern indicated that the pump was cavitating and the resulting water level response was a series of rapid, cyclic fluctuations. **The** test was terminated and rescheduled for the next day after determining a more suitable withdrawal rate (1,600 gpm).

The APT was restarted on April 13, 1983, at 0917 at the rate of 1,600 gpm. Figure 6-1 illustrates the background water level response at Well No. 2 (300 feet north of Well No. 5) just prior to the start of the test.

Data collection during the test was accomplished with Stevens recorders (Wells No. 2, 3, and 4) and chalked, steel tape (Wells No. 1, 5 and the canal). At the start of the test each well, the flow measuring device, and the engine were manned. The start of the test was signaled at the pumped well just after static water levels were measured (all wells and canal). Simultaneously, stop watches were started at all wells. The pumping rate quickly stabilized

at 1,600 gpm. Flow was measured by an 8-inch orifice plate attached to 10-inch pipe and piezometer.

Water levels in Wells No. 1 and 5 were measured using a CHALLED, STEEL tape at regular (logarithmic) intervals.<br>
3  $\frac{1}{x}$  Appendix C lists time-drawdown values obtained during the test. At Wells No. 2, 3, and 4, equipped with Stevens recorders, a different technique was used. Here, at time,  $R \cap$  $t = 0$ , the recorder pen was lifted off the chart. The pen was subsequently dropped and lifted at  $t = 15$ , 30, 45, and 60 seconds. As the test proceeded, the pen was dropped at  $1\frac{1}{2}$ , 2,  $2\frac{1}{2}$ , 3, 4, etc., minutes for every minute until 10 minutes into the test. After marking 12, 15, and 17 minutes, the pen was dropped and raised every 5 minutes until 40 minutes into the test, at which time the pen was dropped and allowed to track a continuous record. At approximately 4 hours into the test, recorder charts were replaced and time scales changed to 24-hour, full scale. Figure 6-2 illustrates the type of data this method produces. During the initial portion of the test, when water levels are dropping (or recovering) rapidly, a single point at a known time is made. The drawdown can be accurately scaled from the 1:1 chart, and the method results in very accurate early time-drawdown/recovery data. Later in the test, as water levels change less rapidly, the pen can be dropped to produce a continuous record requiring only periodic checking rather than continuous staffing. For recovery, the reverse procedure is used, i.e., 24-hour recorder gears are replaced with 4-hour gears, and the pen dropping maneuver is employed at the cessation of pumping.

Throughout the test, recorders were checked at regular intervals and the water levels in the pumped well, Well No. 1, and the canal were measured. Also, flow rate from the pumped well was checked periodically. The withdrawal portion of the APT lasted approximately 48 hours, and recovery was tracked for 4 hours. Appendix C also includes time-recovery data from the pumped well and observation wells.

There are several pertinent observations that can be made regarding the APT based on a review of the continuous water level record at Well No. 3 (see Figure 6-2). These are as follows:

The initial segment of the water level record 1. (just prior to the start of the test) traces a slow, steady decline. This decline represents the aquifer response to evapotranspiration. Although no rainfall occurred during the test, approximately 2 inches fell the previous week. Since the shallow aquifer in eastern Palm Beach County is recharged directly by rainfall, continuous water level records plot the rise of

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the aquifer water levels after rainfall (and the decline when no rain falls).

Extending the pre-test water level decline to the end of the test results in the projection of a 0.5-foot decline in water level over 24 hours, using the same slope as the initial segment. This would be valid except that the water level decline is influenced by other factors including barometric pressure. Changes in barometric pressure cause a water level response in aquifers. Increases in pressure result in water level The amount of water level decline decline. attributed to barometric pressure changes can be estimated by comparing the static water level at the beginning and end of the test. Since the test was started and stopped at approximately the same time, it can be assumed that the daily barometric cycles were approximately equal at the beginning and end of the test. Then, comparing the static water levels at the beginning and end of the test, there is approximately a 0.25-foot difference in water level. Comparing this number to the projected slope of the plot on the initial segment suggests that half of the decline is therefore attributable to daily cyclic barometric pressure change and that the other half is attributable to aquifer response due to lack of rainfall.

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- $2.$ At a point approximately 8 hours into the test, the piezometer attached to the orifice which was used to measure flow slipped down approximately This resulted in the appearance that 2 inches. the flow rate had increased by approximately 100 gpm and therefore the engine was throttled back. This error was quickly discovered and the piezometer and flow rate subsequently adjusted to the proper position. The result can be seen on the continuous water level plot. There is a slight recovery of the water level as the engine (and therefore pumping rate) was throttled back and a return to a steady-state drawdown condition as the situation was corrected.
- $3.$ Approximately midway through the test, water levels began to stabilize and even recover slightly. This is due to the fact that at this point, the cone of depression had stabilized and discharge was balanced by recharge and inflow in the production zone. Therefore, the plot<br>represents a "static" water level response although at a lower elevation (approximately 1 foot lower). During this time segment, the

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water level decline due to evapotranspiration is balanced by the effects of cyclic barometer pressure changes. bs

The water level plot has two "peaks," one at the 4. approximate midpoint of the graph, the other approximately 24 hours later. If a line is drawn connecting the crest of these two "peaks," the scope of that line has the same slope as the initial segment of the plot just prior to starting the test. Again this suggests that pumping has been balanced by recharge and inflow and that the plot represents aquifer "static" response.

At approximately 40 hours into the test, a water  $5.$ level decline was observed at all observation wells. After checking the flow rate, it was  $\frac{1}{2}$ determined that another well must have been turned O on. A thorough search of the area was made and a 6-inch irrigation well used to water grass and<br>() shrubs was located. The well was pumping at approximately 300 gpm and was located more than turned 3,000 feet from Well No. 5. This well caused approximately 0.12 foot of drawdown at Well No. 5.  $\leq$  The irrigation well discharged for approximately M 3 hours, after which water levels at the Forest Hill Village site recovered.

> This drawdown-recovery due to the irrigation well affected all of the observation wells, and therefore all subsequent data plots of time vs. water level will depict this response. This can be seen clearly at the end of the data plots illustrated in this report.

Figure 2-2 illustrates the areal relationships among the pumped well, the four observation wells, and canals at the APT site. Figure 6-3 illustrates the vertical relationship of the wells and canal at the site.

#### DATA ANALYSIS

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Time vs. water level data from the pumped well and observation wells were tabulated (Appendix C) and plotted on 3x5 cycle log/log and 5 cycle semi-log graph paper. Both time/ drawdown and time/recovery data were plotted. In addition, distance/drawdown data at specific times were also plotted on 5 cycle semi-log graph paper.

Figure 6-4 illustrates the plot of drawdown versus distance from the pumped well at times of 100, 200, and 400 minutes after the start of the test.

Only the observation wells were plotted at this scale and aquifer transmissivity was calculated using the nonequilibrium formula developed by Jacob. Transmissivity was determined using Equation No. 1:

$$
T = \frac{528 \text{ Q}}{\Delta s} \tag{1}
$$

where

 $\int_0^0 \frac{s^{\delta}}{t^{\delta}} e^{-\frac{1}{2} \int_0^t e^{t^{\delta}} dt}$ 

- $T =$  Transmissivity (gpd/ft)
- $Q =$  Pumping rate (gpm)
- $\Delta s$  = Slope of the distance-drawdown graph expressed as the change in drawdown per log cycle (ft)

From Figure 6-4, transmissivity was calculated as follows:

$$
T = \frac{528 (1,600 \text{ gpm})}{0.79 \text{ ft}}
$$
  
\n
$$
T = \frac{1,070,000 \text{ gpd/ft or}}{7227,270 \text{ ft}^2/\text{day}} = 143,048 \text{ ft}^2/\text{day}
$$

Therefore, from the slope of the distance-drawdown graph, transmissivity is calculated to be 227,270 ft<sup>2</sup>/day.  $143,040$  ft<sup>2</sup>/day

In reviewing the distance-drawdown plots on Figure 6-4 two major facts should be noted: first, drawdown at Well No. 4 was considerably less than would be expected based on the plotted curves using Wells No. 1, 2, and 3. This could be due to its proximity to the canal  $(v50$  feet) and distance from the pumping well (600 feet). However, it is more likely a function of well construction. Recalling Figure 5-4, Well No. 4 apparently has no screen and what appears to be 1 or 2 feet of open hole. This results in a very poor, inefficient hydraulic connection to the aquifer which in turn results in a poor transmission of aquifer water level change to the well. Due to this fact, distance<br>and time drawdown data obtained from Well No. 4 were not used to formulate conclusions regarding aquifer characteristics.

The second observation made from review of Figure 6-4 is that distance from the canals, a line source of recharge, was found to have no greater effect on those wells in close proximity than on those located farther away. Wells No. 2 and 3 were parallel to Canal E-3 but perpendicular to Canal No. 8 (see Figure 2-2). Well No. 1 was located the furthest from either canal.

Since the hydraulic gradient observed from drawdown measurements made at Well's No. 1, 2, and 3 were approximately equal in all directions, time versus drawdown plots were prepared for the pumped well and observation wells (including Well No. 4). Data were plotted on both log/log and semi-log graph paper and used to calculate aquifer characteristics using two different methods.

Semi-log graphical plots were used to calculate transmissivity and aquifer storage using non-equilibrium equations derived by Jacob. Equation No. 2 was used to calculate transmissivity and Equation No. 3 was used to calculate storage, as follows:

$$
T = \frac{264 \text{ Q}}{\Delta s} \tag{2}
$$

where

$$
T = Transmissivity (qpd/ft)
$$

$$
Q = Pumping rate (gpm)
$$

 $\Delta s$  = Slope of the distance-drawdown graph expressed as the change in drawdown per log cycle (ft)

and

$$
S = \frac{0.3 \text{ T } t_0}{r^2}
$$
 (3)

where

 $S =$  Storage coefficient (dimensionless)

 $T = Transmissingity (qpd/ft)$ 

- $t_{\alpha}$  = Intercept of the straight line at zero drawdown (days)
	- $r =$  Distance from pumped well to the observation well where drawdown measurements were made (ft)

Log/log plots were used to calculate aquifer characteristics using graphical methods described by Hantush, Jacob, and others. Equation 4 was used to calculate transmissivity,

Equation 5 to calculate storage, and Equation 6 to calculate leakance, as follows:

$$
T = \frac{Q}{4 \pi s} L(u, v) \tag{4}
$$

see Lohman 1972 

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 $d$ antush-Jacob

- $T = Transmissingity (ft<sup>2</sup>/day))$
- $Q =$  Pumping rate (gpm)
- $s$  = Drawdown at match point (ft)
- $L(u,v)$  = Leakance function of  $u, v$ ; values obtained from match point on the type curve (dimensionless)

and

 $S = 4T \frac{t/r^2}{1/u}$  $(5)$ 

where

S = Storage coefficient (dimensionless)

 $T = Transmissingity (ft<sup>2</sup>/day)$ 

- $t =$  Time from match point (days)
- $r =$  Distance from pumped well to the observation well where drawdown measurements were made (feet)
- $1/u =$  Values obtained from match point on the type curve (dimensionless)

and

$$
k'/b' = 4T \frac{V^2}{T^2}
$$
 (6)

type

where

$$
k'/b' = \text{Leakance (day}^{-1})
$$
  
T = Transmissivity (ft<sup>2</sup>/day)  
v = Values obtained from curve matrix to

curve (dimensionless)

 $r =$  Distance from pumped well to the observation well where drawdown measurements were made  $(\text{feet})$ 

Figure 6-5 illustrates the time-drawdown plot from data collected at the pumped well (No. 5). Although the pumped well data is not the most appropriate application for the above formulas, transmissivity can be calculated using Equation No. 1. From this plot, there appear to be two distinct trends to the points plotted. A best-fit line has been drawn approximating the trend of both sets of points and the slope per log cycle measured. The general shape of the data points and the best-fit lines seems to indicate that a recharge boundary has been reached by the expanding cone of depression. This data plot, or at least the best-fit lines through the points, resemble a typical recharge boundary condition. That is, during the early time (0 to 12 minutes) water is derived from the production zone only, and the slope of this portion of the curve reflects aquifer hydraulic characteristics accurately. After approximately 12 minutes, the cone of depression created by the pumping well begins to become distorted, expanding at a much slower rate due to recharge. This recharge, either from a line source (canal) or from induced infiltration, results in the drawdown being less than it would otherwise be and thus the later time sections of the curve have a much flatter slope than the earlier segments. Since the later time segment does not accurately reflect aquifer hydraulic conditions alone, only the early time segment can be used to determine aquifer characteristics. Table 6-1 lists aquifer characteristics calculated from the data plot of time/drawdown for Well No. 5 (Figure 6-5).

Figure 6-6 illustrates the time/drawdown plot from data collected at Well No. 2. Again, aquifer characteristics were calculated using Equation 1. The storage coefficient can also be calculated from observation well data, whereas it cannot be calculated from pumped well data. Equation 2 was used to determine storage, and the results are listed in Table 6-1.

Figure 6-7 illustrates the time/recovery plot for Well No. 2. In comparing the two curves, the time/drawdown curve again results in a change in slope at approximately 12 minutes into the test, indicating a recharge boundary The time/recovery curve appears to be a smooth condition. plot, with the best-fit line having the same slope throughout.

Figure 6-8 illustrates the time/drawdown data plotted on a log/log scale water level response collected at Well No. 2. These data, when matched to the type curves developed by

# Table 6-1<br>SUMMARY OF AQUIFER CHARACTERISTICS



<sup>a</sup>Distance from pumped well to the observation well where drawdown measurements were made.

<sup>b</sup>Slope of the time-drawdown graph expressed as change in drawdown per log cycle.

<sup>c</sup>Intercept of the straight line at zero drawdown.

Hantush-Jacob has best assumptions

 $d$ Values obtained from matching the  $log/log$  time/drawdown-recovery data to the type curve.

 $\int$  $\overline{5}$ 

Cooper, were used to calculate aquifer characteristics using Equations  $3, 4,$  and  $5.$ 

Similarly, time/recovery data were also plotted on a log/log scale and matched to the type curve; aquifer characteristics were then calculated (see Figure 6-9). Table 6-1 lists the results of these calculations.

Using this methodology, aquifer characteristics were determined based on data collected from all the observation wells used during the APT. Data plots for Wells No. 1, 3, and 4 are included in Appendix D.

Table 6-1 summarizes the results of these time/drawdown recovery calculations.

Aquifer characteristics determined from time/drawdownrecovery calculations were averaged, resulting in the following approximation:

> $T = 215,000 \text{ ft}^2/\text{day}$  $S = 1.5 \times 10^{-4}$  $k'/b' = 2 \times 10^{-3} day^{-1}$

Values obtained from distance/drawdown plots were:

$$
T = 143,000 \text{ ft}^2/\text{day}
$$
  
 $S = 4 \times 10^{-4}$ 

The results using the average values calculated from time/ drawdown recovery rate do not compare well with distance/ drawdown values.

Comparing the average values for transmissivity for each observation well regardless of the method used results in the following:

> $T_{ave}$  @ Well No. 1 = 218,775 ft<sup>2</sup>/day  $T_{ave}$  @ Well No. 2 = 202,575 ft<sup>2</sup>/day  $T_{avg}$  @ Well No. 3 = 198,325 ft<sup>2</sup>/day

Comparing the average values for transmissivity for each observation well for both log/log and semi-log methods results in the following:

 $T_{ave}$  @ Well No. 1 = 250,500 ft<sup>2</sup>/day (semi-log)  $T_{ave}$  @ Well No. 1 = 245,100 ft<sup>2</sup>/day (log/log)  $T_{ave}$  @ Well No. 2 = 235,400 ft<sup>2</sup>/day (semi-log)  $T_{ave}$  @ Well No. 2 = 169,850 ft<sup>2</sup>/day (log/log)  $T_{ave}$  @ Well No. 3 = 232,950 ft<sup>2</sup>/day (semi-log)  $T_{ave}$  @ Well No. 3 = 163,700 ft<sup>2</sup>/day (log/log)

Some observations can be made from these comparisons. In general, a higher transmissivity is obtained using the semilog data plots and Jacob non-equilibrium equations. Also, there appears to be little difference in transmissivity when calculated from data taken at Wells No. 2 and 3. Distance versus drawdown calculated transmissivity (143,000 ft<sup>2</sup>/day) appears to compare well to the log/log calculated values from data at Wells No. 2 and 3 (average =  $166,800$  ft<sup>2</sup>/day).

Comparing time-drawdown to distance-drawdown observations, it appears that the transmissivity obtained from the average value of log/log data plots is a reasonably good approximation. However, the storage coefficient determination using this method is not accurate because of the effects of recharge. Walton describes a procedure for calculating the effect of a recharge boundary groundwater withdrawal. Walton's method assumes that no drawdown occurs along an effective line of recharge. Under this boundary condition, water levels will drawdown at an initial rate under the influence of the pumping well only. When the recharge boundary begins to affect the production well, the time rate of drawdown will change, continually decreasing until equilibrium is reached. The APT site is bounded by two partially penetrating recharge boundaries, and therefore Walton's method may be somewhat inappropriate for this site. To apply this method to the Forest Hill Village site, these two partially penetrating recharge boundaries (Canal E-3 and 8 are approximately 6 feet deep) are theoretically replaced by one single, fully penetrating boundary which would produce the same effect on the site.

Applying Walton's method, a determination of storage coefficient can be made and the results checked by trial and error against actual data. For this analysis, observed drawdown (stabilized) data for Wells No. 1, 2, and 3 were substituted into Equation No. 7 to calculate a value for the distance (a) from the pumped well to the effective recharge boundary as follows:

$$
s = \frac{528 \text{ Q Log } \sqrt{4a^2 + r^2}}{T}
$$
 (7)

where

 $a = Distance to effective rectangle boundary (ft)$  $r = Distance from observation well to pumped well (ft)$  $Q =$  Pumping rate (gpm)  $T = Transmissing (qpd/ft)$ 

The results of these calculations were as follows:



Once a value is known for the distance to the effective recharge boundary (a), the storage coefficient can be<br>determined by substitution using Equations No. 8, 9, 10, 11, and 12.

$$
s_r = s - s_i = \frac{114.6 \text{ Q}}{T} [W(u) - W(u_i)] \qquad (8)
$$

where

$$
u = \frac{1.87 \text{ r}^2 \text{ S}}{T t} \tag{9}
$$

and

$$
u_{i} = \frac{1.87 \text{ r} i^2 \text{ S}}{T t} \tag{10}
$$

and

$$
W(u) = -0.5772 - Ln u \tag{11}
$$

and

$$
W(u_{1}) = -0.5772-Ln \t u_{1}
$$
 (12)

where

 $S_r$  = Drawdown in observation well (ft)

 $s = Drawdown$  due to pumped well (ft)

 $S_i$  = Build-up to image well (ft)

 $Q =$  Pumping rate (gpm)

 $T = Transmissivity (qpd/ft)$ 

 $S =$  Storage coefficient

- $r =$  Distance from observation well to pumped well (ft)
- $r_i$  = Distance from observation well to image well  $(f<sub>t</sub>)$

Using this method, a storage coefficient of 1 x  $10^{-3}$  was calculated using Wells 1, 2, and 3 (see Table 6-2).

Once aquifer characteristics are known, the percentage of water being diverted from a source of recharge can be calculated using Equation No. 13 together with Figure 6-10 as follows:

$$
F_{f} = \frac{1.87 \text{ a}^{2} \text{ S}}{Tt}
$$
 (13)

where

- $a = Distance to effective rectangle boundary (ft)$
- $S =$  Storage coefficient
- $T = Transmissing (qpd/ft)$

 $t = Time (days)$ 



Table 6-2 SUMMARY OF STORAGE COEFFICIENT DETERMINATIONS

Note:  $r_i = 17,074$  feet

 $S = 1 \times 10^{-3}$ 

GNR61

therefore

$$
F_f = \frac{1.87 (8,537)^2 1 \times 10^{-3}}{1,247,664 (0.278)}
$$
  
= 0.39

From Figure 6-10:

$$
P_r = 40\%
$$

where

$$
P_r
$$
 =  $8$  of water directed from a source of rectangle

The sources of recharge are induced infiltration from the overlying permeable sediments and leakance from the canal. Since the canal is not fully penetrating, it recharges the upper water table, which in turn recharges the production zone.

No attempt was made to rigorously determine the actual amount of recharge contributed by the canal. In other parts of the County where transmissivity of the Anastasia Formation is much lower, a pumping well (or well field) will cause a greater head differential between the canal and the producing zone than was experienced at the Forest Hill Village site. The head differential caused by the pumping well (Well No. 5) during the APT at the closest canal was less than 1 foot (see Fiqure 6-11). The reason for this is that the producing zone at this site has much higher transmissivity than is common for the Anastasia Formation.

A very rough approximation can be made regarding how much water is obtained from canal recharge at the Forest Hill Village. If we assume a very simple model, the site can be replaced by a square having a discharge point at the center to simulate the well field center of pumping. The square is bounded on two adjacent sides by a line source of recharge (LWDD canals 8 and E-3) which are considered fully penetrating for this discussion. Then, if 40 percent of the water produced at the center comes from a recharge source, approximately half would come from the canal. Since the model described above does not exactly fit conditions at the site, a reasonable assumption as to amount of water recharged by both canals is 15 to 25 percent.

Having established a value for aquifer transmissivity, storage, and leakance, a series of theoretical distancedrawdown curves were constructed using steady-state leaky artesian formulas. Figure 6-11 illustrates this series of curves for various pumping rates including the rate used during the APT.

Values used to calculate theoretical distance-drawdown curves were:

> $T = 166,800 \text{ ft}^2/\text{day}$  $S = 1 \times 10^{-3}$

 $k'/b' = 2 \times 10^{-3}day^{-1}$ 

Theoretical curves, if based on appropriate aquifer characteristics, should predict aquifer response to pumping. comparison of theoretical versus actual distance-drawdown relationships was made to determine if the aquifer characteristics arrived at were reasonable. Figure 6-12 illustrates the plot of the actual, stabilized distance-drawdown relationship observed during the APT (the solid line). The theoretical distance-drawdown curve (the dashed line), calculated from the steady-state leaky artesian formula at 1,600 gpm, plots almost directly over the actual curve constructed after 48 hours of pumping.

It appears, therefore, that the aquifer characteristics established for the Forest Hill Village site are reasonable and that theoretical curves can be used to predict aquifer response to pumping. Having developed these curves, it is now possible to design a well field for the site. Had wells not already been constructed on the site, the design would focus on well spacing (location) and withdrawal rates which would efficiently develop groundwater within the site boundaries. However, since production wells have already been located, well field design efforts will be directed toward the establishment of withdrawal rates for the wells.

In selecting the pumping rate for a particular well, several factors must be considered, including:

- $\bullet$ Aquifer characteristics
- $\bullet$ Available drawdown
- Casing size  $\bullet$
- $\mathbf{o}$ Screen conditions
- $\mathbf{o}$ Proximity to recharge boundary
- Well efficiency/specific capacity  $\bullet$
- Need for water  $\bullet$

Aquifer characteristics determine the interference effects among wells in the well field, which in turn affect well Since well spacing has already been established. spacing. interference effects can be mitigated only by adjustments to the individual well pumping rate.

Available *Arawdown* limits the water level to which individual wells can be reduced by pumping. In screened wells, the maximum design pumping level including interference effects is 10 to 15 feet above the top of the screen.

Casing size limits the size pump which can be installed in a particular well, which therefore limits the pumping rate.

Screen condition, which depends primarily upon the age, type, and installation method, may limit the rate of withdrawal. The higher the pumping rate, the higher the likelihood that the well will pump sand if the screen is in poor condition, improperly designed, etc.

Proximity to a recharge boundary might result in a well being rated higher than wells remote from the boundary, because induced recharge from a recharge source would reduce the effects of pumping.

Well efficiency/specific capacity, which is a measure of individual well performance, is a function of construction and development rather than aquifer characteristics. Therefore, a well having a low efficiency and/or specific capacity would be rated lower than perhaps might be possible given the aquifer characteristics.

Finally, after considering all of the above factors, the actual water needed from a particular site must be considered.

As discussed above, aquifer characteristics determined for this site are:



Again referring to the theoretical distance-drawdown curves constructed on the basis of aquifer characteristics (Figure 6-11), interference effects can be determined for various pumping rates. Recalling that well spacing has already been established, the determination of recommended withdrawal rates then becomes an iterative process of assigning pumping rates to each well and evaluating interference effects using the theoretical distance-drawdown curves. As an example, if Well No. 5 is assigned a rate of 1,600 gpm, then theoretically (from Figure 6-11) the drawdown at that well would be approximately  $2-1/2$  feet, assuming that no other wells were in use and that Well No. 5 were 100 percent efficient.

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FC50100.E0



FC50100.E0








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- 1

DRAWDOWN (ft)



 $135$ PBC



135 PBC







PBC

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<u>ទ</u> FC50100





Well No. <u>8</u><br>Date Completed <u>1974</u><br>Casing Diameter <u>12 in.</u><br>Casing Depth <u>160 ft.</u><br>Screen Section <u>160-200 ft.</u>

FC50100.E0

Screen Diameter 1<u>2 in, TS</u><br>Screen Slot Size <u>40</u><br>Gravel Pack Size <u>None</u><br>Pumping Test 1/83<br>Duration 3 hr.

Pumping Rate <u>715 qpm</u><br>Static Water Level 6.2 ft.<br>Maximum Drawdown 22.1 ft.<br>→ Specific Capacity 32 gpm/ft.

PBC

CH<sub>2</sub>M<br>HILL

 $\overline{135}$ 

FIGURE E-8. Well Completion Report-Well No. 8.



Well No. <u>9</u><br>Date Completed <u>2/77</u><br>Casing Diameter <u>12 in.</u><br>Casing Depth <u>104 ft.</u><br>Open Hole Section <u>104-170 ft.</u>

 $\overline{\mathbb{L}}$ 

Note: Geophysical Logs from Well 10.

Screen Diameter N/A<br>Screen Slot Size N/A<br>Gravel Pack Size N/A<br>Pumping Test 3/77<br>Duration 4 hr.

Pumping Rate 1,200 gpm<br>Static Water Level <u>8.1 ft.</u><br>Maximum Drawdown 10.9 ft.<br>→ Specific Capacity 110 gpm/ft.



Well Completion Report-Well No. 9.

 $PBC$  $135$ 



Well No. <u>10</u><br>Date Completed <u>3/77</u><br>Casing Diameter <u>12 in.</u><br>Casing Depth <u>102 ft.</u><br>Open Hole Section <u>102-170 ft.</u>

Screen Diameter N/A<br>Screen Slot Size N/A<br>Gravel Pack Size N/A<br>Pumping Test 12/76<br>Duration 23 hr.

Pumping Rate 1,200 gpm<br>Static Water Level 8.1 ft.<br>Maximum Drawdown 10.9 ft. Specific Capacity 110 gpm/ft.

**FIGURE E-10.** 

Well Completion Report-Well No. 10.

 $135$ PBC

|CH2M<br>|**|||**||HILL



Note: Geophysical Logs from Well 10.

**FIGURE E-11.** CH2M<br>Well Completion Report—Well No. 11.

PBC 135

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FORTRAN Coding Form





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FORM 0128



FORM 0128<br>Rev. 7/84

PS11 (cont.)

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FORTRAN Coding Form







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FORM 0128<br>Rev. 7/84

FORTRAN Coding Form

FORM 0128<br>Rev. 7/84

FORTRAN Coding Form

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FORTRAN Coding Form

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$$
= \frac{4 (213360 \text{FT}^{3} / \text{D A} \times \sqrt{.0007} \text{ d}4)}{11.8 (570 \text{ FT})^{2}} = .00016
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= .000105 \text{ d}4 \text{g}^{-1}
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$$

$$
S = \frac{4Tt}{\frac{1}{u}r^{2}} = \frac{4(176,570 \text{ FT}^{2}/D44)(.0007 \text{ day})}{12.2 (570 \text{ FT})^{2}} = .00012
$$



71\





09:06 TIME PLOT NO 0019<br>DATE 87/07/31 TAPENO 6153<br>USER NO SHINE  $\begin{array}{c} \square \\ \square \end{array}$ 

Palm Springs APT Model



**ECL** cue

Pumping Node 15, 15 Obs Well Nodes  $2 - 7$   $15,10$  $5 - 7$   $24.13$  $4 \rightarrow 22, (3 \big/21, 13)$  $3 - 18,13$  $|\rightarrow 10,13$ 

E-3 Nodes

 $1 - 29$ , 11

 $L - 43$  Nobes

 $12,1-29$ 

**DATE ETAG.**  $\mathbf{B}$ CHKD'



Palm Springs Pump Test Reanalysis SEFTRAN Grid



type pftps.out

## SFTPS -> Node & Element Data

SEFTRAN-PC

Written by:

GeoTrans, Inc. 209 Elden Street, Suite 301 Herndon, VA 22070

SEFTRAN: FINITE ELEMENT FLOW AND TRANSPORT VERSION 1.1 COPYRIGHT 1984 GEOTRANS, INC., HERNDON, VA

NUMBER OF PROBLEMS TO BE SOLVED = 1

**PROBLEM NUMBER 1** 

seftran trial palm springs apt

MATERIAL PROPERTY LIST







 $\sim 10^{-1}$ 

DEFAULT INITIAL VALUES (HEAD OR CONCENTRATION)

 $.0000$ 

NO. OF NODES FOR WHICH DATA SPECIFIED .  $=$  $\bullet$ CODE FOR GENERATION OF MESH DATA. . . . = 1

FROBLEM FORMULATION CODE. . . . . . . . = 1 TIME STEPPING INDEX . . . . . . . . . . . = 1 NO.OF NODES FOR WHICH I.C. IS TO BE READ =  $\circ$ 

PRINT OUT DELETION CODE . . . . . . . . = 0



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## NODE NUMBER AND X AND Y COORDINATES





NUMBER OF DIRICHLET BOUNDARY CONDITIONS

NUMBER OF FLUX BOUNDARY CONDITIONS. . . = 17

FLUX BOUNDARY CONDITION DATA



MAXIMUM FULL BANDWIDTH = 37

ELEMENT NUMBERS AND CENTROIDAL COORDINATES



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TIME STEP NUMBER = 1

PRINT CHECK FOR ELEMENT NO. 1

 $SSA = .3704E+02$  THETA = .1000E+01

ELEMENT MATRIX: AA - CC - AI1 - EI1





 $.0000E+00$ 

 $.0000E+00$ 



TIME VALUE = 1070ELA1

type sftps1.out

S E F T R A N - P C

Written by:

GeoTrans, Inc. 209 Elden Street, Suite 301 Herndon, VA 22070

SEFTRAN: FINITE ELEMENT FLOW AND TRANSPORT VERSION 1.1 COPYRIGHT 1984 GEOTRANS, INC., HERNDON, VA

NUMBER OF PROBLEMS TO BE SOLVED = 1

**PROBLEM NUMBER 1** 

seftran trial palm springs apt

MATERIAL PROPERTY LIST





INITIAL TIME STEP SIZE . . . . . . . . = .1500 INITIAL TIME VALUE . . . . . . . . . . = ,0000 TIME MULTIPLIER . . . . . . . . . . . . = 1,500 MAX. VALUE OF TIME STEP . . . . . . . . = 3,000

DEFAULT INITIAL VALUES (HEAD OR CONCENTRATION)

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NO. OF NODES FOR WHICH DATA SPECIFIED  $P = 0$ CODE FOR GENERATION OF MESH DATA. . . . = 1 The process of the first and the process and an analysis and process and

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SOLVING FOR DEPENDENT VARIABLE 1

FRINT CHECK NDFLX-FVAL-QVAL

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and the property of the control of the cona and and the first fill of the first state. **A CALACTE TRACT** T.ILOVETVJ .. UUUUE+UU TIME VALUE =  $.1500E+00$ TIME STEP NO. 1 COMPLETED TIME STEP NUMBER  $=$  2 SOLVING FOR DEPENDENT VARIABLE 1 FRINT CHECK NDFLX-FVAL-QVAL 5 - 1230E+05 .0000E+00 6 - 2550E+05 .0000E+00 7 - 1230E+05  $.0000E+00$ TIME VALUE = .3750E+00  $\sim 10^{-1}$ TIME STEP NO. 2 COMPLETED TIME STEP NUMBER = 3 SOLVING FOR DEPENDENT VARIABLE 1 PRINT CHECK NDFLX-FVAL-QVAL 5 -1230E+05 0000E+00 6 -2550E+05 0000E+00 7 -1230E+05  $.0000E+00$ TIME VALUE = .7125E+00  $\sim 10$ TIME STEP NO. 3 COMPLETED TIME STEP NUMBER =  $4$ SOLVING FOR DEPENDENT VARIABLE 1 FRINT CHECK NDFLX-FVAL-QVAL 5 - 1230E+05 .0000E+00 6 - 2550E+05 .0000E+00 7 - 1230E+05  $.0000E+00$ TIME VALUE =  $.1219E+01$ 

TIME STEP NO. 4 COMPLETED



TIME STEP NUMBER = 5

SOLVING FOR DEPENDENT VARIABLE 1

FRINT CHECK NDFLX-FVAL-QVAL

5 -1230E+05 .0000E+00 6 -2550E+05 .0000E+00 7 -1230E+05

Paria

 $.0000E+00$ 

TIME VALUE =  $.1978E+01$ 

NODE NUMBERS AND CORRESPONDING HEAD VALUES



REQUIRED INFORMATION AT OBSERVATION NODES 

OBSERVED NODE NUMBER = 162

TIME VERSUS NODAL VALUE OF DEPENDENT VARIABLE (HEAD OR CONCENTRATION)

.1500E+00 8000E-09 .3750E+00 -.3390E-06  $7125E+00$  -.1884E-04  $.1219E+01 - .2551E-03$  $.1978E+01 - .1592E-02$ 

OBSERVED NODE NUMBER = 161

TIME VERSUS NODAL VALUE OF DEPENDENT VARIABLE (HEAD OR CONCENTRATION)

 $.1500E+00 - .9223E-09$  $3750E+00 - 3494E-06$  $7125E+OO$   $-1945E-O4$  $1219E+O1$  -.2627E-03  $1978E+01 - 1632E-02$ 

OBSERVED NODE NUMBER = 179

TIME VERSUS NODAL VALUE OF DEPENDENT VARIABLE (HEAD OR CONCENTRATION)

 $.1500E+00 - .1366E-09$  $.3750E+00$  .  $1418E-07$ .7125E+00 -.2339E-06  $1219E+01$  -. 2081E-04  $.1978E+01 - .2510E-03$ 

OBSERVED NODE NUMBER = 213

TIME VERSUS NODAL VALUE OF DEPENDENT VARIABLE (HEAD OR CONCENTRATION)

 $.1500E+00$   $-.5013E-11$  $.3750E+00$   $.1490E-09$  $.7125E+OO$   $-.1061E-OB$  $1219E+01 - 9448E-09$  $.1978E + 01$ .5859E-06

OBSERVED NODE NUMBER = 214

TIME VERSUS NODAL VALUE OF DEPENDENT VARIABLE (HEAD OR CONCENTRATION)

 $.1500E+00 - .4919E-11$  $3750E+00$  .1425E-09  $.7125E+00$  -  $.1131E-08$  .  $.1219E+01$  .  $.2774E-10$  $.1978E+01$   $.7494E-06$ 

C: \JAMES>

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