HYDROGEOLOGIC INVESTIGATION OF HARBOUR RIDGE LIMTED ST. LUCIE COUNTY, FLORIDA

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HYDROGEOLOGIC INVESTIGATION OF HARBOUR RIDGE LIMITED ST. LUCIE COUNTY, FLORIDA

INTRODUCTION

In 1981, Harbour Ridge Limited, contracted Geraghty & Miller, Inc., to conduct a hydrogeologic investigation of its St. Lucie County property, the location of which is shown in Figure 1.

The first phase of the investigation consisted of the installation of six two-inch-diameter test wells to obtain initial on-site hydrogeologic data. From the data obtained in Phase I, locations were selected for Phase II, a more detailed testing program consisting of one test-production well, three observation wells, and three salt-water monitoring wells. After the installation of all wells in Phase II, a 72-hour constant-rate pumping test was conducted.

This report contains a summary of all data obtained in the testing programs, a description of the hydrogeologic system that occurs at the site, an analyzis of the hydrologic coefficients derived, and a model of the impacts of a proposed well field.

FINDINGS

- l. The water-table zone of the shallow aguifer extends from 20 feet to 60 feet below land surface. The water-table aquifer consists predominantly of fine-grained sand with traces of clay. Below the water-table zone, a layer of sandy clay occurs with a thickness of 6 to 11 feet. The production zone of the shallow aquifer exists below the clay to a depth of 129 to 150 feet below land surface.
- $2.$ The production zone in the shallow aquifer is 80 to 110 feet below land surface.
- $3.$ The ground-water quality in the Harbour Ridge area is generally qood. The water is treatable and a potable product could be delivered.
- 4. In the vicinity of Harbour Ridge, the production zone of the shallow aquifer responds to pumping as a leaky artesian one, with recharge by vertical leakage downward through a confining bed consisting of sandy clay.

 $5.$ The aquifer coefficients are estimated as follows:

Transmissivity = $100,000$ qpd/ft Storage Coefficient = .0002 Leakange = $.002$ gpd/cu.ft.

 $6.$ A diversion of 0.644 mgd can be obtained from the shallow aquifer without causing adverse impacts.

WELL CONSTRUCTION METHODOLOGY

To determine geologic and hydrologic conditions in the project area, a total of 16 wells and 6 test pits were installed during 1981. All wells were installed by Maxson Well Drilling, Inc., of Lake Worth, Florida, under the direction of Geraghty & Miller, Inc., hydrogeologists. Except for the test-production well which was installed by the cable-tool method, all wells were installed by the drive-wash method. The six test pits were installed by Harbour Ridge during their investigation into fill material available beneath the site and were constructed to a depth of two to three feet below the water table. In addition to the wells and test pits, three water-table piezometers were installed adjacent to Wells R1, R2, and TFW1. Locations of wells, piezometers, and test pits are given on Figure 2. All wells were left in place-for-future monitoring use.

Except for the test-production well, all wells were constructed with 2-inch-diameter steel casings and PVC screens. The test-production well was gravel-packed and constructed with a 16-inch-diameter outer casing. The inner casing was an 8-inch-diameter steel pipe attached to a 30-foot-long stainless steel screen. The test-production well complies with the rules of the Florida Department of Environmental Regulation (DER), chapter 17-21, Water Wells in Florida; and chapter 17-22, Water Supplies; and was permitted by DER. Construction details, depths, and screen intervals of all wells are shown in Figure 3.

The salt-water monitor wells were constructed to determine variations of ground-water quality with respect to depth in the vicinity of the North Fork of the St. Lucie River. The wells were installed by the drive-wash method to facilitate water sampling. Two-inch-diameter, open-ended, galvanized casing was driven in 21-foot sections. After each section was driven, the formation samples were collected by washing; and a water sample was collected by air-lift pumping. Upon completion, all wells were surveyed to determine their elevations referenced to NGVD.

HYDROGEOLOGIC CONDITIONS

Shallow Aquifer

The drilling of sixteen wells and the installation of six test pits on the Harbour Ridge property generated a substantial amount of site-specific hydrogeologic data. Geologic logs of wells are given in Appendix A. The data reveal that from present land surface to depths of 20 to 60 feet below land surface, the formation consists predominantly of fine-grained sand with small amounts of clay. Ground water is encountered a few feet below land surface in this zone at the water table. The land saturated section of these suficial sands is the water-table portion of the shallow aquifer. In this interval, the overall water-yielding capacity of the material is fairly low; high-capacity wells cannot be completed in this section. Below the surficial sands, the formation becomes finer and consists of a sandy The clay, which varies in thickness from 6 to 11 feet, acts as a clav. confining layer between the water-table and production zone of the shallow aquifer. The most permeable section of the production zone occurs between 80 and 110 feet below land surface and consists of medium- to coarse-grained, partly-cemented sand, and limestone. Where it occurs, this section can yield large quantities of water to production wells. Below this unit, at a depth from 129 to 150 feet below land surface, the formation is a very fine-grained sand that grades downward to a grayish-olive clay. This clay marks the base of

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the shallow aquifer. The water-table zone, confining bed, and production zone are the major components of the shallow aquifer.

Water-level elevations were measured in all wells and test pits during January 1982. From these data, contour maps of water levels in the water table and production zone were constructed and are shown on Figures 4 and 5. From these two maps (Figures 4 and 5), it can be seen that the natural directions of ground-water flow in both the water table and the production zone are from the southwest to the northeast towards the North Fork of the St. Lucie River. These two contour maps also show that the water levels at the water table are 1 to 2 feet higher than the heads in the production zone.

Floridan Aquifer

The Floridan aquifer consists of a series of soft limestone formations interbedded with dense dolamite and occasional clays. The aquifer is over 1,000 feet thick with the top occurring between 300 feet and 450 feet below land surface. Within the limestone, extensive solution channeling has occurred, creating a high secondary permeability which accounts for the large yields obtainable from the aquifer. Overlying this limestone is a confining layer known as the Hawthorn Formation, consisting primarily of impermeable sand clays.

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The principal use of water from the Floridan aquifer in St. Lucie County is for irrigation. Users obtain their water from individual wells that are allowed to flow freely into ditches for irrigation. The excess water is transported to the St. Lucie River by means of drainage canals.

WATER QUALITY

Shallow Aquifer

To determine if water quality in the shallow aquifer in the area of Harbour Ridge was such that a potable product could be delivered with conventional treatment methods, water samples were collected and analysed from the test-production well and the salt-water monitoring wells.

Test Production Well-The analysis of the water sample collected from the test-production well is presented in Appendix B. The sample was obtained on December 5, 1981, after approximately 72 hours of continuous The water is of good quality. The chloride concentration of pumping. 25 mg/l falls well within the recommended limit of 250 mg/l for public water supply. Total dissolved solids of 558 mg/l are what one might for this area. The iron concentration, although above expect

recommended DER limit for public supplies, is not excessive and can be treated so that a potable product can be delivered to consumers.

Salt-Water Monitoring Wells-To determine the potential for salt-water intrusion into the ground-water system from the North Fork of the St. Lucie River, three salt-water monitoring wells (R1, R2, and R3) were installed on Harbour Ridge property adjacent to the River (see Figure Water samples were collected approximately every 20 feet during the 2 . drilling and analyzed for chloride content by field methods. The results of these analyses are given in Table 1.

As can be seen from the Table 1, chloride concentrations in the shallow aquifer at the locations of Wells R2 and R3 are less than 115 mg/1 and it is likely that no saline water occurs within the water-table zone or production zone in those areas. At the location of Well R1, saline water is present in both the water-table and production zone. The highest chloride concentrations occur in the production zone of the shallow aquifer.

Floridan Aquifer

Detailed water-quality data on the Floridan aquifer are very limited and are reported in the literature only as general trends. Because the Floridan aquifer is principally used for irrigation and because chloride

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TABLE 1

SALT-WATER MONITORING WELLS, RESULTS OF CHLORIDE ANALYSIS HARBOUR RIDGE, ST. LUCIE COUNTY, FLORIDA

I All results are reported in milligrams per liter Notes: as $C1^-$. $2 N/S = No sample obtained.$ 3 The base of shallow aquifer was 129 feet below

land surface at this location.

in the irrigation water, concentrations major are a concern water-quality data most often record chlorides only. The range of chloride concentration in water from this aquifer in St. Lucie County, as reported by the SFWMD in selected wells from 1977 to 1979, range from 280 mg/l (milligrams per liter) to 1250 mg/l (Reece, D. E., et al, Progressively higher chloride concentrations generally occur at 1980). In isolated instances, however, higher chloride qreater depths. at shallower depths than would normally be concentrations occur expected. This phenomenon is thought to be due to localized withdrawals, allowing water of poorer quality from the deeper zones to move upward. The potential for degradation must be considered in the planning of any diversion from the Floridan aquifer.

Ranges in the quality of water in the Floridan aquifer in the vicinity of Harbour Ridge are shown on Table 2. This table was compiled from numerous analyses performed on two wells by the South Florida Water Management District from 1977 to 1979.

To the east of St. Lucie County, a transition zone exists between the relatively fresh water in the Floridan aquifer and salt water. The location of this contact (salt-water interface) is presently unknown, but presumably occurs seaward of the barrier beach where water quality is still similar to that in interior parts of the County.

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TABLE 2

RANGES IN CONCENTRATIONS OF CHEMICAL CONSTITUENTS IN WATER SAMPLES FROM FLORIDAN AQUIFER WELLS IN THE VICINITY OF HARBOUR RIDGE (Reece, D. E., et al, 1980)

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PUMPING TESTS

A step-drawdown pumping test and a constant-rate pumping test were conducted on the test-production well at Harbour Ridge. Both tests were conducted with the same test arrangement.

 \mathbf{A} right-angle drive, gasoline-powered vertical turbine pump was installed in the well, and a six-inch-diameter PVC discharge line was run approximately 300 feet to a nearby shallow pond in a closed depression. The pumping rate was controlled by a gate valve located close to the pump, and a six-inch totalizing flow meter was installed in the discharge pipe so that the pumping rate could be determined. An access for water-level measurements (which were made with an electric probe) was made available between the inner well casing and the pump column.

Step-Drawdown Pumping Test

A step-drawdown pumping test was conducted on the test production well $(TFM1)$ December 1, 1981. The step-drawdown test was conducted for on the following purposes:

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- 1. To establish pumping levels at various production rates to aid in determining the design of a permanent pump,
- 2. to determine the magnitude of water-level response in nearby monitoring wells to pumping stress so that the proper water-level responses in the constant-rate pumping test could be anticipated,
- 3. to set the gate valve and motor speed to a suitable pumping rate for the constant-rate test, and
- 4. to establish a baseline for future performance tests of the well and permanent pump.

Pumping steps were 30 minutes in length and were followed by 30-minute recovery steps. The test data are shown in Table 3. These data indicated that a rate approximately equal to that in the first step would be most useful for the 72-hour constant-rate pumping test.

Interpretation \circ f step-drawdown data by the method of Rorabaugh ("Graphical and Theoretical Analysis of Step-Drawdown Test of Artesian Well," American Society of Civil Engineers, December 1953) shows that the equation: $s = 0.022$ Q + 0.01396 Q^{1.08} can be used to predict drawdown in TPW1 after 30 minutes, and: $s = 0.025 Q + 0.01396 Q^{1.08}$

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TABLE 3

DATA FROM STEP-DRAWDOWN PUMPING TEST OF TEST-PRODUCTION WELL HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

Note: All depths referenced to measuring point 1.35 feet above land surface,

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has been generated using data from the constant-rate pumping test and can be used to predict drawdown at stabilization.

Constant-Rate Pumping Test

A constant rate pumping test was begun on December 2, 1981. Water levels were measured in TFM1 by means of an electric probe, and in monitoring wells by hand-held wet tapes. A layout of the pumping test is shown in Figure 6. Water levels were measured periodically in TFW1 and the monitoring wells in the hour prior to the test. Water levels were stable in all wells during this period.

The test began at $9:00$ a.m. on December 2, 1981. The pumping rate varied between 390 and 400 gpm (gallons per minute) in the first 2 minutes of the pumping before it was stabilized at 400 grm. Pumping continued for 72 hours at this rate $(+/- 0.58)$. Water levels were measured frequently (each minute or more frequently during the first 20 minutes of pumping) in TPW1, M-1, M-2, M-3 and the piezometer adjacent to TPW1. After 72 hours, the pump was shut off and the recovery of water levels were measured for three hours in all wells.

When Test Production Well 1 was turned on, the water levels in all wells monitoring the production zone declined--rapidly at first, and later at

a continually decreasing rate. The largest decline occurred in the pumped well; drawdown was progressively smaller at increasing radial distances from the pumped well. After about one day of pumping, water levels appeared to stabilize in the monitoring wells.

No change in the water level in the piezometer adjacent to TPW1 was observed during the test other than normal daily fluctuations. Due to this and the fact that the shallow pond has bottom peat accummulations from 1 to 3 feet, it does not appear that any substantial amount of water was released into the shallow aquifer during the pumping test.

At the completion of pumping, recovering water levels were measured in several monitoring wells. Their responses mirrored those during the pumping phase; recovery was initially rapid, but slowed at a continually decreasing rate. Measurements of recovering water levels were discontinued before the levels had recovered to the pre-test levels.

When pumped, the aquifer responded as a leaky artesian one. The fine-grained sand, shells, and clay above the production zone appear to act as a leaky confining bed. Below the production zone, the formation contains clay units of very low permeability-so low that this lower confining bed may be assumed to be impermeable.

Drawdown and recovery data from individual monitoring wells and TPW1 were plotted on semi-logarithmic or double-logarithmic graph paper, or

Two methods of analysis are appropriate in determining aquifer both. coefficients from these test data. For data graphed on. double-logarithmic paper, a method described by Walton ("Selected Analytical Methods for Well and Aquifer Evaluation," W. C. Walton, McGraw-Hill Publishing Company, 1962) has been applied. For data graphed on semi-logarithmic paper, the Hantush I method described by Kruseman and DeRidder ("Analysis and Evaluation of Pumping Test Data," G. P. Kruseman and N. S. DeRidder, International Institute for Land Reclamation and Improvement, 1976) has been used. As noted previously, pre-pumping and pre-recovery water-level trends were small and have been disregarded. Normal daily water-level fluctuations were small enough so that the effect on test data is not significant; they also were disregarded.

methods of analysis complement each other although the The two assumptions made about the ground-water system are the same for each method. In the Walton method, the most critical data are collected in the first few minutes of the test, and the data are visually matched to a type curve. Matching is often somewhat arbitrary depending on the skill of the analyst. Conversely, the critical data in the Hantush I method are collected after the first few minutes of testing and until stabilization. Interpretation is based on determining or estimating the value of stabilized water level and on fitting a straight line to the data collected in the middle portion of the test. The data and interpretations from representative wells are shown in Appendix C.

One additional check can be made on the accuracy of the aquifer coefficients interpreted by these two methods. Jacob's modification of Theis equation ("A Generalized Graphic Method for Evaluating the Formation Constants and Summarizing Well-Field History," H. H. Cooper and C. E. Jacob, American Geophysical Union Transactions, vol. 27, 1946) can be applied to data after the first few minutes of pumping or recovery and before stabilization. Although the major assumption upon which this method is based (that the system is a non-leaky artesian aquifer) is not valid at Harbour Ridge, leakage has only a small effect on the shape of the data curve before stabilization. Analysis of transmissivity at the pumped well by Jacob's modified method should produce an approximation of transmissivity obtained by other methods. Storage coefficient and leakance cannot be calculated by this method at any well or for the pumped well because the radius of measurement (effective radius of the pumped well) is unknown. Table 4 summarizes aquifer coefficients determined by the various methods. Values of 100,000 gpd/ft for transmissivity, .002 gpd/cu.ft. for leakance and .0002 for a storage coefficient are considered to be representative of this aguifer.

TABLE 4 AQUIFER COEFFICIENTS DETERMINED FOR INDIVIDUAL WELLS AS A RESULT OF CONSTANT-RATE TEST OF TEST-PRODUCTION WELL HARBOUR RIDGE ST. LUCIS COUNTY, FLORIDA

*Methods of Analysis Used - Hantush I (H), Cooper-Jacob (CJ), or Jacob Modification (JM).

IMPACT OF PROPOSED WITHDRAWAL FROM SHALLOW AQUIFER

The potential amount of water available from the shallow aquifer at the Harbour Ridge site is the amount of fresh water that can be pumped from the aquifer on a sustained basis without causing adverse impacts. At Harbour Ridge, the greatest concern is the potential for salt-water intrusion from the North Fork. The possibility of adverse impacts at Harbour Ridge is real because the property is not large enough for production wells to be located more than a few thousand feet from the North Fork of the St. Lucie River. It is necessary, therefore, to estimate how far in the coastal direction the impacts of the proposed diversions will ultimately extend.

The limit of the catchment area surrounding the well is established by the ground-water divide that develops when drawdowns produced by pumping form a cone of depression around a well. The ground-water divide is where flow in the aquifer changes from towards the river to towards the well, and is marked by a limiting flow line and by zero hydraulic gradient on the aquifer's potentionetric surface.

The catchment area could be approximated from a cone of depression based on drawdowns calculated by appropriate mathematical formulas. These formulas are derived on the basis of simplifying assumptions about the ground-water flow system. For example, an appropriate formula for the flow system at Harbour Ridge would be the leaky artesian aquifer

equation developed by Jacob and Hantush for steady-state flow. The field situation at Harbour Ridge is such that the formula's assumption that the only water available to the well is vertical leakage through the confining bed induced by the cone of depression that is developed is overly conservative. Other factors that can be considered, which are not taken into account in the Jacob and Hantush analytical formulas, are lateral inflow from outside the area of influence of the pumping wells (natural gradient) and natural vertical leakage which occurred even before pumping began. Therefore, the catchment area needed to achieve the desired diversion and consequently, the distance to the ground-water divide will be less in practice than is predicted by the analytical formulas.

Mathematical Simulation

A more accurate mathematical simulation is needed to estimate how far the ground-water divide will ultimately extend towards the areas of brackish ground water occurring along the North Fork of the St. Lucie River. This can be achieved by more realistically simulating field conditions at Harbour Ridge, particularly the areal distribution of inflow and leakage occurring before pumpage starts. This can be done by subdividing the flow system into discrete volume elements and using numerical approximation methods to solve the resulting finite difference equations. Relevant parameters controlling flow in a leaky artesian

aquifer system such as water-level altitudes, soil permeability, and thickness are built into the equations on a volume-element by volume-element basis. Because of the large number of volume elements needed to represent differences in water-level altitude in sufficient detail, and the large number of calculations that must be made upon each element, digital computers are used to solve the finite difference equations. This results in a flow-system model in which differences that occur from one place to another can be accounted for.

The supply source at Harbour Ridge is a leaky artesian aquifer. About 200 volume elements are used to simulate the aquifer being pumped and an equal number volume elements to simulate the leaky confining bed, over a two-square-mile-area (Figure 7). Because no water will be supplied from aquifer storage under steady-state conditions, it is only necessary to use the volume elements to define the geometry of the flow system, the hydraulic coefficients of the aquifer and confining bed, and conditions (heads or flows) at the boundaries of the flow system and the well.

The computer is programmed to achieve an iterative solution of the resulting system of equations by using the method of successive over-relaxation. The correct solution is achieved when successive iterations result in negligible changes in calculated water-level altitudes at all volume elements. As a check on the validity of the iterative solution, the computer is also programmed to calculate a mass balance to be sure that under the resulting head distribution the sum of

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all computed inflows equals the sum of all computed outflows, which must prevail under steady-state flow.

Description of the Model

The geometry of the flow system in the vicinity of the proposed wells, and the node-centered volume elements used to model it are shown in Figure 7. Each element is 754 feet on a side (except where they are bisected by boundaries). Along the lateral boundaries of the model (rows 1 and 15), no flow occurs. Flow of fresh water is simulated as perpendicular to, and toward the river, where the outflow boundary This outflow is approximated to occur along a straight line occurs. lying on the average a short distance off-shore (column o). The production wells that will intercept some of this flow during pumping are located at elements F6 and E13.

The lateral boundaries (rows 1 and 15), are maintained as no-flow boundaries even though some flow across them to the well will occur when pumpage starts. This is done to simplify the model, although it also leads to a conservative estimate, because the effect of assuming no flow across lateral boundaries is to exaggerate the growth of the wells' catchment area toward the area of brackish ground water.

Another boundary on the model shown in Figure 7 is where inflow occurs across the model's southwestern limit (column A) from upland areas. This boundary is simulated by flows calculated from potentionetric head and transmissivity data. However, the potentiometric heads beneath up-gradient areas to the west can only be estimated because no observation wells are located off-site. As topographic maps show that the land surface is about +16 feet NGVD (National Geodetic Vertical Datum) in areas far enough to the west, say 3,000 feet, to not be significantly effected by pumpage, and available data suggest that potentiometric heads are about six feet below land surface, the potentiometric heads in this area are estimated to be at +10 feet NGVD. Therefore, this boundary is simulated by using inflows calculated on the presumption that potentiometric heads will remain constant at about ten what about 600 pampage. feet NGVD in the source area, This inflow is not fixed but will increase as water levels are drawn down by the pumpage along the site's southwestern boundary. It is modeled by using an expression of Darcy's Law for the flow term in the finite difference equation at the volume elements along this boundary.

Before the model can be completed, one must also define conditions along the boundary receiving precipitation recharge, which is the upper limit of the flow system. This flow boundary is not shown in Figure 7 because it affects every volume element. In a leaky artesian aquifer, this can be done by specifying the average water-level distribution through the overlying water table, which is the source of leakage supplying the

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These water levels can be presumed to be constant in the pumped wells. model, because any leakage-induced drawdown of the water table will be offset by a resulting increase in net recharge. This is because the accompanying increase in the depth to the water table, which is close to the surface in many areas, will reduce the natural evaporative loss of ground water. what about wellands?
ground water. This reduction in discharge will be larger relative to the leakage, because evaporative losses from the shallow water-table are high.

In on-shore areas the water table or its outcropping such as lakes is the upper boundary. Along the North Fork of the St. Lucie River where outflow by vertical leakage must occur, the heads along the upper limit of the flow system are those in the St. Lucie River which are taken as 0 feet NGVD. The water-table elevations used to define this boundary on the model are shown in Table 5. They are derived from the measurements made in the water table observation wells and pits during the field investigation (Figure 4).

The other flow limit boundary, the bottom of the shallow aquifer, is implicitly modeled by the inclusion of aquifer thickness in the definition of transmissivity. The lower confining bed is assumed to be impermeable.

Table 6 shows the model input data for each square volume element needed to define the hydraulic characteristics of the aquifer system at the

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TABLE 5

WATER-TABLE ELEVATIONS USED IN FINITE-DIFFERENCE MODEL HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

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		А	в	С	D	E.	F	G	н		J	κ		м	N	\mathbf{o}
		8.3	8.3	8.3	7.7	7	6.8	6.6	6.4	6, 2	4	O	0	0	0	0
	2	8.3	8.3	8.2	7	6.9	6.8	6.6	6.5	6.3	5.3	2	O	0	0	$\mathbf 0$
	з	8,3	8.2	8.1	6.9	6.8	6. .	6.5	6.3	6.1	5.8	3.5	0	0	0	$\overline{0}$
	4	8.2	8.1	8	6.9	6.8	6.7	6.5	6.3	6.1	6	5	1.8	0	0	0
	Б.	8.2	8.1	8	6.9	6.8	6.	6.6	6.5	6.3	6.1	5.1	2	0	0	$\mathbf{0}$
OWS	6	8.2	8.1	8	7.3		6.8	6.8	6.8	6.4	6.1	5.3	2.1	0	0	$\bf{0}$
		8.2	8.1	8	7.7	7.5	7.2		6.8	6.5	6.1	5.5	2, 2	0	0	$\mathbf 0$
ŒΙ	8	8.2	8.1	8	\cdot 8	7.6	7.5	9.25	9.25		6,2	5.3	2.3	0	0	0
	9	8.2	8.1	8	7.8	7.6		.25 19	9.25	7.2	6.3	5.2	2.1	0	O.	$\overline{0}$
	10	8.2	8.1	8	7.8	\cdot 6	5	.25 19.	9.25	7.3	6	5.1	$\mathbf{2}$	0	0	$\mathbf{0}$
	11	8.2	8.1	7.9	7.8	7.6	7.5	9.25	9.25	7.2	5.9	5	2	0	0	$\overline{0}$
	12	8.2	8.1	7.9	7.8	7.6	\cdot 5	9.25	7.5	7	5,6	4	0	0	0	0
	13	8.1	8	7.8	7.8	7.6	7.5	7.3	7.2	6	5.3	2, 5	0	0	0	0
	14	8.1	7.9	7.8	7.8	7.6	7.5	7.5		5.8	5	$\overline{2}$	0	Ω	0	$\mathbf 0$
	15	8	7.9	7.8	7.8	7.6	7.5	7.5	6.2	5.8	4	0	0	0	0	0

COLUMNS

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Note: Locations of columns and rows shown on Figure 7.

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TABLE 6

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AQUIFER SYSTEM COEFFICIENTS IN CONSISTENT UNITS HARBOUR RIDGE . ST. LUCIE COUNTY, FLORIDA

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site, as determined by the field studies, in the consistent units required by the finite difference equations.

The model is not complete until discharge at the offshore outflow boundary is accounted for. However, there are no data defining the discharge-controlling parameters, water levels in the production zone, or confining bed characteristics in this offshore area. The outflow boundary condition is established by extrapolation of present day on-shore water levels and by assuming that they will not be drawn down significantly by pumping because the wells are nearly a mile away.

Calibration

As some of the boundary condition data had to be estimated, it is necessary to check the model's accuracy. Because the model calculates internal heads, a check can be made by determining whether it can duplicate the water levels in the production zone under non-pumping conditions. Table 7 shows the results of this calculation. It agrees closely with the observed water levels (Figure 5) so no adjustment was made in the model input data that had to be estimated. Table 8 gives the flows into and out of each volume element of the production zone under non-pumping conditions.

TABLE 7

WATER LEVELS IN PRODUCTION ZONE UNDER NON-PUMPING CONDITIONS AS CALCULATED BY FINITE-DIFFERENCE MODEL HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

	VULUMNƏ															
		л	в	С	D	Е	F	G	Η			N		м		Ο
		7.971	7.61		7,25[6,88]6,50]		6.11	5,72	5.31	4.891	4.44 3.98			13.5413.131	2.75	2.4
	-2	7.97	7.61		.25 6.88 6.50		6.12	5.72				$5.32 \mid 4.89 \mid 4.45 \mid 3.99 \mid 3.55 \mid 3.13 \mid$				2.4°
	з	7.97	7.61		7.25 6.88 6.50				6.12 [5.73] 5.32] 4.90		4.47	4.01	13.56			2.4
		7.97	7.62		7.26[6.88]6.51[6.13]			5.74	5.34	4.92	4.48	4.03	3.58	3.151		2.4
	Б	7.97	7.62		26 6.89]	6.526.14			5.75 5.35			$4.93 \mid 4.50 \mid 4.05 \mid 3.60 \mid$		3.1		2.4
တ၊	6	7.98	7.63	7.27	16.911				$\left 6.54 \right 6.16 \left 5.77 \right 5.37 \left 4.95 \right 4.52 \left 4.07 \right $				3,61	3.17		2.4
- 31				7.98 7.63 7.28 6.92					6.556.185.795.3914.97			$4.53 \mid 4.08 \mid 3.62 \mid 3.17$				2.4
Ol		7.99 ^T							7.64 7.29 $\left[6.93\right]$ 6.57 $\left[6.20\right]$ 5.82 $\left[5.42\right]$ 4.99			$4.54 \mid 4.09 \mid 3.62$		3.181		2.4
ŒГ	9			7.9917.6517.3016.941				6.58 6.21 5.84 5.43		15.00	4.55 4.09		3.62	3.18		2,4
		8.001							$7.65 7.30 6.95 6.58 6.22 5.84 5.44 5.00 4.55 4.08 3.62 3.17$							2.4
									8.00 7.65 7.30 6.95 6.59 6.22 5.84 5.43 4.99 4.54 4.07				13.61			2.4
	12	8.00		7.65 7.30 6.95 6.58 6.21 5.83 5.41								$\left[4.98\right]4.52\left[4.05\right]3.59$		13.161		2.4
	13	8.00	7.651		7,30 6,95			(6.58)6.2115.81	5.40	4.96	4.50	4.03	13.57	3.151		2,4
		8.00	7.65		7.30 6.94	6.58	6.20	5.81!	5.39	4.94	4.48	4.01i	13.561	3.14 2.76		2,4
	15	8,00		$7.65 7.30 6.94 6.58 6.20 5.80 5.38 4.94 4.48 4.00 3.56 3.14 $												2,4

COLUMNS.

Note: Locations of columns and rows shown on Figure 7.

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TABLE 8

MASS BALANCE OF INFLOWS AND OUTFLOWS PRODUCTION ZONE UNDER NON-PUMPING CONDITIONS AS CALCULATED BY FINITE-DIFFERENCE MODEL HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

 ~ 100 km s $^{-1}$

									<u>culumns</u>							
		А	В		Đ	E	F	G	Ħ					м	N	O
			2366152.3	79.8	62.5138.1					$52.1166.8182.6199.71-33.$			-302 -269	-238	-2091	$-2E3$
						60.91		104.1133.1	180 .	214	129.	-3031	-539	-476	-418	-5E3
	з	4729 89.2				45.2	88.1		148.1	182	203	$-78.$	-541		-4191	$-5E3$
			4716[73.6]	113	2.46 44.1			86.8 116.	147. 180.		1230.	147	-271	-479	-420	$-5E3$
	6.	47081	(72.9)	112.	1.11			42.5 84.9 129. 175.		208.	243	159.	$-243'$	-480	-4201	$-5E3$
œ.			4697 71.9	111				60.0 70.6 97.5 156. 217.		220.	1240.	187.	$-229!$	-482	-421	$-5E3$
		4685	70.8	109	119.	144.	155.	$\overline{1183.1214.}$		232, 238,		216.	-2151	-482	-42	$-5E3$
O!	8	4673	69.8	108.	132.	157.	198.	1521.	582.	306.	1252.	184.	∣−201	-483	-42	$-5E3$
ŒΙ	8	4664	69.01	107	'131	1155.	120.	519.1		$\sqrt{580}$, 335, 266,		169.	-231	-483	-421	$-5E3$
	10		4657(68.4)	106.	$\sqrt{130}$.	154.		$\overline{1195, 1518, 1579.}$		350.	$\overline{1221}$.	155.	-246	-482	-42	$-5E3$
	11		4653 68.1		90.9 130.	1154.	195.	518.	580.	336.	207.	141	-244	-481	-42	$-5E3$
	12		4652 68.0	90.8	130.	154.	196.	1520.	317.	308.	1164.	-7.7	-545	-480	-420	$-5E3$
	13		4646 52.9	75.8	I30.	155.	197.	1226.1	274.158.		122.	-232	-543	-478	-419	$-5E3$
	14	4647	37.9	75.9	130.	T155.		$\overline{197}$, $\overline{258}$, $\overline{245}$, $\overline{130}$,			$\sqrt{78.3}$	-306	-541	-478	-419	$-5E3$
	-51				2320 19.0 38.0 65.0 77.7					$\overline{98.9129.162.3165.5} - 36.$		-304	-270	-239	-209	$-2E3$

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Note: Locations of columns and rows shown on Figure 7,

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It can be seen that there is somewhat more discharge to the St. Lucie River than inflow from upland areas, with the difference supplied by downward leakage through the confining bed. It can also be seen that the production zone is receiving more recharge from leakage in the area of Mile Lake where the water table is highest.

Results of Impact Analysis

After calibration, the ultimate effect of continuously pumping two wells at a combined rate of 86,296 cu.ft./day (0.644 MGD) was determined. Table 9 shows how pumping changes the flow distribution. It can be seen by comparison with Table 8 that most of the water pumped will be supplied by an increase in inflow from upland areas and a decrease in discharge to the River, with some also being supplied by a net increase in downward leakage through the confining bed.

Table 10 and Figure 8 shows the new water levels in the production zone calculated by the model as a result of two wells continuously pumping at a combined rate of 0.644 mgd, and the position of the resulting ground-water divide. It can be seen that none of the water within about 3000 feet of the St. Lucie River should move into the well. A comparison of Figure 5 and Figure 8 shows that water levels in the production zone along the North Fork of the St. Lucie River will be slightly lower but the magnitude of the natural gradient towards the

TABLE 9

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MASS BALANCE OF INFLOWS AND OUTFLOWS PRODUCTION ZONE UNDER STEADY-STATE CONDITIONS WITH 0.644 MGD PUMPAGE AS CALCULATED BY FINITE-DIFFERENCE MODEL HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

	UULUMNS															
		л.	в	С	D	E		G	н					м		Ο.
		3295 130.		166.	156. 137.		151	162.1		$ 170.1176.131.6 -251$			-230	-212		
	2	6608 261.		320.	211	263.1				$309.$ 329. 358. 370.	260.	-198	-461		-196 $-1E3$ -425 -393	$-2E3$
	з	6659 251		312.	207.	264.	312.	330.	338.	345.1	337.	128.6	$-462]$	-425	-393	-253
		6717	1242.	307.	1224.		293.1350.1	357.	354.	353.		370.1256.	-190 [-426	$-393 - 323$	
Ø		6770	1248.	318.	245.	336.1	427.	415.1404.		391.	388.	1271.	$-161!$	-427	-394 $-3E3$	
	6	6783	1249.	322.	319.		$413. -7E4$ 491.		461.	408.1	387.	1299.1		-148 -429 -395 -3E3		
Ωi			6739 245.	314.	362.			436. 495. 467.	441.	$413.$ 381.				$1325. -136 - 430$	$[-395]$ -3E3	
œι	8	$665/1237$.			301.1352.	403. 458. 759. 785.				474.1387.		$\sqrt{289. - 124}$		-432	$[-396; -3E3]$	
	9	6568 229.		7288.		$\begin{bmatrix} 332 & 371 & 340 & 725 & 762 & 489 \end{bmatrix}$					$\overline{393.}$ $\overline{269.}$ $\overline{-157}$ $\overline{-434}$				$1 - 397$	$T = 3E3$
	10.	6497122.		279.	319.	354.				394. 705. 747. 494.	1340.	$ 249. -175 $		-436 -398 $-3E3$		
			6455 219. 261.		316.	351 .				$\overline{1386.1695.1738.1472.321.}$		$\overline{1232}$.		$\overline{ -176 -436 -398 -3E3}$		
	12	6440 218.		262.	322.	365, 391, 694, 470,				439.	274.			79.6 -480 -436 -398		-3E3
	13	6430	203, 249,		1335.	$-2E4$ $ 402.1401.$			425.	287.	228.	-147	'−479∫	$-436 - 398$		$-3E3$
	4	64231	187.1247.		323.	366.	$\overline{1391.1428.}$		394.	257.1	183.	-222		$-478 - 436 - 398 - 323$		
	16.				3206 93.2 122. 159.	1177	[193.							$\sqrt{213. 136. 128. 16.1 -263 -239 -218 -199 -153}$		

COLUMNO

Note: Locations of columns and rows shown on Figure 7.

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TABLE 10

WATER LEVELS IN PRODUCTION ZONE UNDER STEADY-STATE CONDITIONS WITH 0.644 MGD PUMPAGE HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

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Note: Locations of columns and rows shown on Figure 7.

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Geraghty & Miller, Inc.

River will remain the same. From these results, it can be seen that a diversion of 0.644 mgd from the shallow aquifer can be maintained on the Harbour Ridge property without causing adverse impacts.

> Respectfully submitted, GERAGHTY & MILLER, INC.

sG, Whea

James A. Wheatley Staff Scientist

Bermes

Boris J. Bermes Senior Scientist

Geraghty & Miller, Inc.

 $\sim 10^{-11}$

APPENDIX A

 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

GEOLOGIC LOGS

GEOLOGIC LOG OF TEST WELL HR-1 HARBOUR RIDGE
ST. LUCIE COUNTY, FLORIDA

Geologic Log of
Test Well HR-1
Harbour Ridge

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GEOLOGIC LOG OF TEST WELL HR-2 HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

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Geologic Log of
Test Well HR-2
Harbour Ridge

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GEOLOGIC LOG OF TEST WELL HR-3 HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

Geologic Log of
Test Well HR-3
Harbour Ridge

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GEOLOGIC LOG OF TEST WELL HR-4 HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

 $\sim 10^7$

GEOLOGIC LOG OF TEST WELL HR-6 HARBOUR RIDGE ST. LUCIE COUNTY, FLORIDA

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GEOLOGIC LOG OF WELL NUMBER R-1 HARBOUR RIDGE

 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

 $\sim 10^7$

Geologic Log of
Well Number R-1
Harbour Ridge

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 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

Geologic Log of
Well Number R-1 Harbour Ridge

 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

Geologic Log of
Well Number R-1 Harbour Ridge

TOTAL DEPTH: 150

 $\sim 10^{-11}$

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 $\sim 10^{-10}$

GEOLOGIC LOG OF WELL NUMBER $R - 2$ HARBOUR RIDGE

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Geologic Log of
Well Number R-2 Harbour Ridge

Geologic Log of
Well Number R-2 Harbour Ridge

Geologic Log of
Well Number R-2 Harbour Ridge

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GEOLOGIC LOG OF WELL NUMBER R-3 HARBOUR RIDGE

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Geologic Log of
Well Number R-3 Harbour Ridge

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Geologic Log of
Well Number R-3 Harbour Ridge

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GEOLOGIC LOG OF TEST PRODUCTION WELL #1 HARBOUR RIDGE

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Geologic Log of
Test Production Well #1 Harbour Ridge

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Geologic Log Of Test Production Well #1 Harbour Ridge

olive gray; Phosphorite, trace.

Geologic Log Of
Test Production Well #1 Harbour Ridge

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GEOLOGIC LOG OF WELL NUMBER M-1 HARBOUR RIDGE

Geologic Log of Well Number M-1 Harbour Ridge

TOTAL DEPTH: 110

GEOLOGIC LOG OF WELL NUMBER M-2 HARBOUR RIDGE

Geologic Log of
Well Number M-2 Harbour Ridge

Geologic Log of
Well Number M-2 Harbour Ridge

TOTAL DEPTH: 110

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 $\label{eq:2} \frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^{2} \left(\frac{1}{\sqrt{2}}\right)^{2} \left(\frac{$

GEOLOGIC LOG OF WELL NUMBER M-3 HARBOUR RIDGE

 \bullet

Geologic Log of
Well Number M-3 Harbour Ridge

Geologic Log Of
Well Number M-3 Harbour Ridge

Geraghty & Miller, Inc.

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 $\label{eq:2.1} \frac{1}{\sqrt{2}}\int_{\mathbb{R}^3}\frac{1}{\sqrt{2}}\left(\frac{1}{\sqrt{2}}\right)^2\left(\frac{1}{\sqrt{2}}\right)^2\left(\frac{1}{\sqrt{2}}\right)^2\left(\frac{1}{\sqrt{2}}\right)^2\left(\frac{1}{\sqrt{2}}\right)^2.$

 $\sim 10^6$

 $\mathcal{L}(\mathcal{A})$ and $\mathcal{L}(\mathcal{A})$

 $\sim 10^{-10}$

APPENDIX B

 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

WATER QUALITY ANALYSIS OF TEST PRODUCTION WELL

GEOTEC, INC 1602 CLARE AVENUE WEST PALM BEACH, FL 33401

SECONDARY WATER ANALYSIS

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LAB ID 86122

DATE January 11, 1982 BY

LAB ID 86122

GEO FE

PRIMARY WATER ANALYSIS

SAMPLE NO. 112-788

1665 Palm Beach Lakes Blvd. West Palm Beach, FL 33401

PROJECT NO.

DATE COLLECTED BY client TIME

LOCATION JOB # P528DB1 PURPOSE

 \mathbf{r}

DATE $> \alpha$ January 11,1982

 $BY = 8k$

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EVERULADES LABURATURIES, INT. 1602 CLARE AVENUE - WEST PALM BEACH, FL 33401 - 305/833-4200

LAB ID 86109

PRIMARY ORGANICS

METHODS: 1. METHOD FOR ORGANOCHLORINE PESTICIDES IN INDUSTRIAL EFFLUENTS. EPA 1973

- 2. METHOD FOR CHLORINATED PHENOXY ACID HERBICIDES IN INDUSTRIAL EFFLUENTS, EPA 1973
- 3. METHOD 608, ORGANOCHLORINE PESTICIDES AND PCBs FEDERAL REGISTER VO. 44 DEC. 3, 1979
- 4. OTHER:

 $\texttt{DATE}_12 - 26 - 81$

Geraghty & Miller, Inc.

APPENDIX C

 $\sim 10^{-11}$

REPRESENTATIVE DRAWDOWN DATA AND INTERPRETATION

 \mathbb{Z}

Time in minutes

4.

e AS.

Hantush Image Harkour Ridge $\frac{9}{7}$ DRI Method $Q = 2126 m^{3}/day$ $\Delta S_{mp} = 1.77 (m)$ $\Delta S_{m}x = -9.6 \times 10^{-3}$ $2 = -3$ (76) $\Delta S_{MB} = \frac{2.30 \text{ Q}}{2\pi kD}$: $kD = \frac{2.3 Q}{2\pi 4 S_{MB}} = \frac{2.3 (2126)}{2 \pi (1.77)} = \frac{440 m^2 M \omega l}{2}$ (75) $45mx = \frac{QSz}{4\pi (k_0)^2}$: $S = \frac{4\pi (T)^2 \Delta Smx}{Qz} = \frac{4\pi (440)^2 - 96*10^{-3}}{2126 (-3)}$. 37 $T = 440 \text{ m}^2$ /day x 80.52 = [35, 400 gpd/ft] $S = .37$? Recharge Boundary is 300° from production well

 $m-1$ Hantussh Iwage Method Harbour Ridge DRI
 $\frac{155=.012}{5^{11}6}=5\times10^{-6}$ $25m t = \frac{012}{5 \times 10^{-5}} = 2.4 \times 10^{2}$ $r = 100$

 $\frac{5}{1}$

 $\frac{\Delta DD}{\Delta 1/k} = \frac{9\times10^{-3}}{5\times10^{-5}} = 1.8\times10^{-2} = \Delta S_{mt} = 3.8\times10^{-2} (m \text{ day})$

OBS. WELL#M-3 Harbour Ridge DRI

Hautush Image Method

 $r = 633$

 $M^2 \rightarrow M^2 \oplus M^2$ = $\oplus N^2$ = \oplus

PUMPAL \sim cose

Hunkosh Meltud || M-1 | Hirkour Eidge. Rumup T=1%
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$$
S_m = 4.28'
$$
\n
$$
S_p = \frac{1}{2}(s_m) = 2.14' x .3048 = .65 m
$$
\n
$$
C_p = \frac{1}{2}(s_m) = 2.16 = .102 \text{ rad} \times .3048 = .65 m
$$
\n
$$
Q = 2126 m^{3}/day
$$
\n
$$
S_p = 3.18 - 2.16 = .102 \text{ rad} \times .3048 = .31 m
$$
\n
$$
S_p = 3.18 - 2.16 = .102 \text{ rad} \times .3048 = .31 m
$$
\n
$$
S_p = 2.30 (2.097) = 4.82 = e^{-1}/k_{0} (r/L)
$$
\n
$$
2.30 \frac{65}{3} = 2.30 (2.097) = 4.82 = e^{-1}/k_{0} (r/L)
$$
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$$
T = 30.5
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$$
L = \frac{30.5}{.009} = 33.89 \approx 3300 m
$$
\n
$$
T = 30.5
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$$
L = \frac{30.5}{.009} = 33.89 \approx 3300 m
$$
\n
$$
RD = \frac{2.30 Q}{4 \pi k Q} e^{-r/L} = \frac{2.3 (2120)(.941) = 1244 m^{3}/day
$$
\n
$$
1244 m^{2}/day \times 80.52 = \frac{100160 qp/44}{4 \pi (51)} = 1244 m^{3}/day
$$
\n
$$
1244 m^{2}/day \times 80.52 = \frac{100160 qp/44}{4 \pi (51)} = 124 m^{3}/day
$$
\n
$$
S = \frac{4 k D}{r^{2} 2 L} = \frac{4(244) (6.6 \times 10^{-3}) (30.5)}{(30.5)^{2} 2 (3300)} = \frac{16.8 \times 0^{-4}}{16.8 \times 0^{-4}}
$$
\n
$$
T = \frac{16.8 \times 0^{-4
$$

 $\frac{1}{\sqrt{2}}$, $\frac{1}{\sqrt{2}}$

$$
\frac{c \text{backon } \text{Hant } \text{val} \pm M4}{\text{at } t = 0.1}
$$
\n
$$
\frac{c \text{backon } \text{Hant } \text{val}}{4 \text{ (l24)}} = \frac{(30.5)^{2} (1.6 \times 10^{-4})}{4 (124^{2}) \cdot 1} = 2.99 \times 10^{-4}
$$
\n
$$
\omega(u_{1}r/L) = 7.5
$$
\n
$$
\omega(u_{1}r/L) = \frac{2126 (7.5)}{4\pi (124)} = 1.02 \text{ m} \times 328 = 3.34
$$
\n
$$
\omega \text{black} = 7.31
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\omega \text{black} = 7.31
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\omega \text{rank } \text{U} = 7.31
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<u>r=300 Harbour Ridge</u> <u>Ventush</u> I $M_{\rm \star}$ $S_m = 3.15'$ S_{ρ} = 1.575 x.3048 = . 48 m t_{p} = 30 min x 6.944 x 10⁻⁴ = 2.083 x 10⁻² days $\Delta s_p = \Delta s$ par log cycle = 2.15 - 1.05 = 1.1' x.3048 = . \$35n $eq 33: 23 \frac{6p}{\Delta s_p} = e^{7L} K_0(r/L)$ $(a.3)(4\%/3\%) = 3.30 = e^{1/2} K_0(1/2)$ $\pi r e^{r_L} K (r_L) = 33, x = .048 = r_L$ $r = 91.44 \, m \implies L = \frac{1905 \, m}{1005}$ $\frac{6}{4}$ = .048 Δ sp= .335 $e^{-0.048} = .953$ $Q = 2126$
 $Q = 763$ e_{4} 31: $\Delta \phi_{P} = 2.30 Q_{40 K} e^{-1/4}$ $kD = (2.3)(2126)(.953) / 4T(.335) = 1107 m²/day$ $(107)(80.52) = 69,136$ gpd/ft = T $eg. 30: r/21 = \frac{r^2}{4\kappa D t\rho} = \frac{4\kappa D t\rho}{r^2} = \frac{4\kappa D t\rho}{r^2}$ $S = \frac{(4)(107)(2.083 \times 10^{-2})}{(91.44)(3)(905)} = 2.65 \times 10^{-4} = 5$ $c = \frac{12}{kD} = \frac{(905)^2}{1107} = 3278$ days $Leakance = |Y_{C} = 3.05 \times 10^{-4}|$ = 1/_C * 7.48 gpd/f13 = $\sqrt{2.3810^{3}}$ $*$ I did not check this

Hantush I Hartour Riche $M-3$ $r = 632' = 193 m$ $5m = 2.5'$ $Q = 2126$ m³/day $S_p = 1.25'$ x.3048 = .381 m t_{0} = 75 min = 5.208 $\times 10^{-2}$ days $\Delta Sp = 1.0^{\circ} = 0.3048$ m e_4 33: (2.3)(.381)/3048 = 2.875 = $e^{74}K_0(r/L)$ for $c''^{\mu}k_0(r\mu)$ =2.875, $x= .079 = r\mu$ $(r=193 m)$ $L = \frac{193}{1079} = 12443$ m $r/t = 0.079$ = $e^{-r/t} = .924$ C_{φ} 31: KD=(2.30)(2126)(.924) $4\pi(0.3048)^{2/3040}$ $\frac{2.30}{4\pi\Delta^{5}\rho}$ (e^{tr/2}) KD = 1180 $m^2/a_{4} = [94,982,999/4=T]$ $eq30: S = 4kDtp/r2L$ = $(4)(1180)(5.208 \times 10^{-2})$
(193)(2)(2443) $= 2.61 \times 10^{-4} = 5$ $c = L^2/\kappa_D = (2443)^2/(180) = 5058$ days Leakance $\frac{1}{2}$ + $\frac{1}{2}$ + $\frac{1}{2}$ + $\frac{1}{2}$ + $\frac{1}{2}$ = $\frac{1}{c} \times 7.48 \frac{q^{2d}}{43} = 1.5 \times 10^{-3}$ * No checkee

Hintu's L

\n1 M3 Hribov Pids,
$$
-\frac{633'}{193m}
$$

\n5m = 2.5 μ t (.76m)

\n6p = 125 μ t (.38m)

\n4p = 76 min (25p = 18(y)) (.3m)

\n33) 2.30 5μ s, 5μ t. κ (r/L)

\n2.3 $38/3 = e^{-t}k$ s. (r/L)

\n2.4 $38/3 = e^{-t}k$ s. (r/L)

\n2.5 $28/3 = e^{-t}k$ s. (r/L)

\n2.6 $1 = r$ s. 239 m

\n31 Δ s, $-\frac{239}{4\pi k}$ s. $-\frac{1}{2}$ s. $-\$

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LOGARITHMIC 3 x 5 CYCLES
KEUFFEL & ESSER CO. MADE IN U.S.A K∘∑

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Harbour Ridge Well <u>M</u> Hantush Leaky Artesian Aquifer w/ Storage in the Confining Layer $u = 10^{-4}$ pZ84 Feller $Q = 400$ gpm $W(\mu, B) = 10$ 220 Walton $r = 100$ B = 5,0 X 10⁻² $t = 12$ min b' =10 feet $s = 6.9$ feet T = 114.6 Q W(u, B) /s
= 1146 (900) 10 / 6.9 = 166,435 gpd/fl] S' = Tut/2693 r^2 = $66435 (01)(12)/2693 (100)^{2}$
= 3x10-4 $K' = 16\beta^{2} Tb'S/r^{2}S'$ = 16 (.05)² 66,435 (10) (.0003) /100² (.0003) $K = 2.66$ S/g' apd/fiz

46 6210

$$
\Delta 5 = 2.20 \text{ V}_0 = 6000 \text{ T} = 1200
$$
\n
$$
T = \frac{(528)(390)}{2.20} = 93,600 \text{ W/4}
$$
\n
$$
S = \frac{(33,600)(1200)}{(4790)(6900)^2} = 4.93 \times 10^{-4}
$$

the contract of the contract of the

 $\mathcal{A}^{\mathcal{A}}$

$$
25 = 2.47
$$
 $r_0 = 2400$ $t=120$

$$
T=(528)(390) = 83,368974
$$

$$
S = \frac{(83,368)(120)}{(4,790)(2400)^{2}} = 3.63 \times 10^{-4}
$$

$$
\overline{S} = 4.265 \times 10^{-4}
$$

= 89,496

 $\mathcal{L}(\mathcal{$

 $\label{eq:2.1} \mathcal{H}_{\mathcal{A}}(x) = \mathcal{H}_{\mathcal{A}}(x) + \mathcal{H}_{\mathcal{A}}(x) + \mathcal{H}_{\mathcal{A}}(x) + \mathcal{H}_{\mathcal{A}}(x)$

Mater 1005 Harbour Klikits

KOF SEMI-LOGARITHMIC 5 CYCLES X 70 DIVISIONS

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III. feccilitus

 $\mathcal{L}_{\mathcal{A}}$

 $V = 390$ opm = 2126 m/day △っ= . チゥm $r_e = 1720$ m $\Delta 5 = 2.36 / \pi kD \implies KD = 2.38 / \pi \Delta 5$ KD= (2.3)(2126) = 1037.65 m^2 day = transmissivity
= 83,552 gpd/st

 $C = \frac{(r_o / 1.12)^2}{kD}$

 $\mathcal{L}^{\text{max}}_{\text{max}}$ and $\mathcal{L}^{\text{max}}_{\text{max}}$

 $C = (1750/1.12)^{2}$

 $\frac{1038}{1038}$ = 2352 days

and the control of the state of

 $\label{eq:2.1} \frac{1}{\sqrt{2}}\frac{d\mathbf{r}}{d\mathbf{r}} = \frac{1}{\sqrt{2}}\frac{d\mathbf{r}}{d\mathbf{r}}\frac{d\mathbf{r}}{d\mathbf{r}}\,,$

 L eakance = V_c x7.46 = 3.2×10⁻³

 $\mathcal{L}(\mathcal{L}(\mathcal{L}))$ and $\mathcal{L}(\mathcal{L}(\mathcal{L}))$. The contribution of $\mathcal{L}(\mathcal{L})$

 $\mathcal{L}(\mathcal{L}(\mathcal{L}))$ and $\mathcal{L}(\mathcal{L}(\mathcal{L}))$. The contribution of $\mathcal{L}(\mathcal{L})$

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 $\overline{\mathsf{S}}$

 $\ddot{}$

$$
L = 1/2.3 \text{ or } = 1/2.3 \times 10,000 = 4348
$$

\n
$$
C = L^{2}/kD
$$

$$
kD = 2.3 Q/4\pi (2s)
$$
\n
$$
T = 2.3 (2126)/4\pi (32) = 1216 = 97,900
$$

l,

 $\ddot{}$