TECHNICAL PUBLICATION 83-7

October, 1983

DRE 175

FIELD INVESTIGATION INTO THE FEASIBILITY OF STORING FRESH WATER IN SALINE PORTIONS OF THE FLORIDAN AQUIFER SYSTEM, ST. LUCIE COUNTY, FLORIDA

TECHNICAL PUBLICATION #83-7

DRE 175

FIELD INVESTIGATION INTO THE FEASIBILITY OF STORING FRESH WATER IN SALINE PORTIONS OF THE FLORIDAN AQUIFER SYSTEM, ST. LUCIE COUNTY, FLORIDA

bу

Leslie A. Wedderburn and Michael S. Knapp

This publication was produced at an annual cost of \$525.00 or \$1.04 per copy to inform the public. 500 191 Produced on recycled paper.

October 1983

Groundwater Division
Resource Planning Department
South Florida Water Management District
West Palm Beach, Florida

TABLE OF CONTENTS

<u>Page</u>
INTRODUCTION 1
PREVIOUS INVESTIGATIONS
PURPOSE, SCOPE AND METHODS
RESULTS OF INVESTIGATIONS14
Stratigraphy
Introduction14
Cenozoic Erathem14
Eocene Series
Avon Park Limestone14
Ocala Group16
Miocene Series
Hawthorn Formation17
Miocene/Pliocene Series
Tamiami Formation17
Pleistocene/Holocene Series
Anastasia Formation18
Undifferentiated Terrace Desposits18
Hydrostratigraphy
Introduction18
Surficial Aquifer System19
Hawthorn Confining Beds19
Floridan Aquifer System20
Aquifer Parameters23
Water Quality27
Surface Water Quality27
Surficial Aquifer System Water Quality

Table of Contents (Continued)

		Page
	Floridan Aquifer System Water Quality	32
	Floridan Aquifer System Water Levels	35
	Hydrogeologic Model of the Injection Zone	36
	Injection/Recovery Test	42
	Cost Evaluation	63
CONC	CLUSIONS	68
REFE	RENCES	69
APPE	ENDICES	
	Appendix 1 - Lithologic Log, Well SFL-50	.1-1
	Appendix 2 - Geophysical Logs, Well SLF-50	.2-1
	Appendix 3 - Water Quality	.3-1
	Appendix 4 - Aquifer Test Data and Analyses	.4-1
	Appendix 5 - Summary of Packer Test, SLF-50	.5-1
	Appendix 6 - Data from Injection/Recovery Tests	.6-1

LIST OF FIGURES

Figure	<u>Р</u>	age
1	Site location and Hydrogeologic Data in Vicinity of Site	6
2	Location of Wells Tapping the Floridan Aquifer System in the Vicinity of the Injection Site	8
3	Site Plan	10
4	Construction Details, Exploratory Well (SLF-50) and Monitor Well (SLF-51)	11
5	Stratigraphy, Lithology and Hydrogeology of Exploratory Well SLF-50	15
6	Producing Zones in Well SLF-50 and Percent Contribution of Each Zone to Flow in Borehole	2 2
7	Daily Subsurface and Weekly Chloride Concentrations at S-49	28
8	Sediment Concentration and Particle Size, C-24 at S-49	30
9	Producing Zones in Well SLF-50 and Water Quality Profile	34
10	Hydrograph of Well SLF-50, July 20 to August 20, 1982	37
11	Variations in Injection Rate, Injection Pressure and Specific Conductance during Injection Test	43
12	Variations in Injection Capacity and Potentiometric Heads During Injection Test	44
13	Variations in Chloride (field values) with Volume of Water Recovered	46
14	Variations in Specific Conductance (field values) with Volume of Water Recovered	47
15	Variations in Temperature (field values) with Volume of Water Recovered	4 8
16	Relationship Between Specific Conductance and Chloride Concentrations in Recovered Water	49
17	Variations in Sodium with Volume of Water Recovered	50
18	Variations in Potassium with Volume of Water Recovered	51

List of Figures (Continued)

Figure		<u>Page</u>
19	Variations in Calcium with Volume of Water Recovered	52
20	Variations in Sulfate with Volume of Water Recovered	53
21	Variations in Alkalinity with Volume of Water Recovered	54.
22	Variations in Hardness with Volume of Water Recovered	55
23	Variations in pH with Volume of Water Recovered	56
24	Variations in Total Iron with Volume of Water Recovered	57
25	Variations in Total Dissolved Iron with Volume of Water Recovered	58
26	Variations in Total Dissolved Strontium with Volume of Water Recovered	59
27	Variations in Total Dissolved Solids with Volume of Water Recovered	60
28	Recovery Efficiency Based on Chloride Concentrations In Recovered Water	62
29	Costs Per Thousand Gallons of Usable Water Recovered	66
	LIST OF TABLES	
Table		<u>Page</u>
1 .	Summary of Pumping Test Data and Results	25
2	Approximate Chloride Concentrations in C-24 Near Injection Site, and Canal Levels at S-49 (from Bearden, 1972)	27
3	Potentiometric Heads at Injection Site	36
4	Preliminary Conceptual Model of Injection Horizon	38
5	Costs per Well Based on Hydrogeologic Conditions at Well SLF-50	65

ACKNOWLEDGEMENTS

The authors wish to acknowledge the assistance and guidance provided by Mr. Abe Kreitman, former Director, Groundwater Division, South Florida Water Management District, in the development and execution of this study. Thanks are also due to Mr. Fred Meyer, U. S. Geological Survey, Miami Sub-district Office, who willingly gave advice and assistance during the study. This study is one of the many technical investigations authorized by the South Florida Water Management District in the execution of its mandate to provide for the conservation, development, management and proper utilization of the water resources within its boundaries.

SUMMARY

This preliminary report presents data from the initial phase of field investigations into the feasibility of storing fresh water in the saline portion of the Floridan Aquifer System in St. Lucie County. During this study the lithology, water bearing properties, and water quality of the first potential injection horizon within the Floridan Aquifer System at the selected site were investigated. A brief injection/recovery test was performed to assess the operation of the system.

The study indicated that several zones existed within the upper part of the Floridan Aquifer System which could accept and store injected fresh water. Surface or shallow groundwater of marginally suitable characteristics for injection existed in the area. The major drawbacks with the available water were: a) variably moderate to high chloride concentrations (100 to 200 mg/l), b) variable but often high suspended solids concentrations in the surface water, and c) high iron concentrations in both surface and shallow ground water.

At the site chosen for exploratory work, a suitable injection horizon consisting of three producing zones was identified and tested. The injection horizon had moderate transmissivity (45,000 gpd/ft) and total dissolved solids concentration (approximately 2000 mg/l).

Water from a shallow well drilled on the site was injected at rates from 400 gpm to 200 gpm in the zone between 600 feet and 775 feet. Reduction in injection rate and increases in injection pressure indicated plugging of the well bore during injection. This was confirmed by high suspended solids concentration and total iron in backflush from the injection well.

Due to the initial high chloride content of the injected water, approximately 3 percent of the recovered water had chloride concentration below 250 mg/l. However, it was calculated that for an 80/20 blend of

injected water with native water, the recovery efficiency would be 33 percent for the first cycle of injection. At 100 percent recovery the recovered water was still significantly less mineralized than the native groundwater.

Water quality data indicated a gradual increase in mineralization of the recovered water, suggesting that the volume of water which can be recovered would vary in an approximately linear manner with the maximum allowable concentrations of critical water quality parameters (chloride and total dissolved solids). The quantity of recoverable water therefore depends both on the quality of the injected water and the maximum allowable concentrations in the recovered water.

Costs associated with the injection/recovery method are likely to exceed present costs for agricultural water in the area. The study, however, indicated that substantial cost reductions may be possible if suitable sites were obtained at which high rates of injection could be maintained without pretreatment of the injection water or frequent well rehabilitation, and if the quality of the recovered water was not a critical factor.

FIELD INVESTIGATION INTO THE FEASIBILITY OF STORING FRESH WATER IN SALINE PORTIONS OF THE FLORIDAN AQUIFER SYSTEM. ST. LUCIE COUNTY. FLORIDA

INTRODUCTION

In August 1980, the South Florida Water Management District released results of a preliminary investigation into the feasibility of cyclic storage of fresh water in the saline Floridan Aquifer System, Upper East Coast Planning Area (Khanal, 1980). This study indicated that the technique was feasible and recommended a program of field investigations to confirm the efficacy of the method. Field investigations were initiated in 1981 with the following objectives:

- To identify a suitable location for construction of facilities to perform injection/recovery tests.
- 2) To design and construct an exploratory well to determine aquifer parameters, resident water quality, and optimum well design.
- 3) To design and construct a test-injection well and associated monitoring wells for long-term injection/recovery tests.
- 4) To design and implement long-term injection/recovery tests in the test injection well.

This interim report summarizes the results of the first two phases of the field investigations. The report is intended to form the basis for decisions as to the desirability of completing the final phases of the project which would include multiple-cycle long-term injection/recovery tests. It can also serve as supporting documentation for application for a Class V Underground Injection Control Permit as required by Chapter 17-28, Florida Administrative Code, should this be required for construction of facilities for the final phases of the project.

PREVIOUS INVESTIGATIONS

The techniques for injection of fresh water into saline aguifers are similar in many respects to those employed in artificial recharge to fresh water aguifers. Artificial recharge has been practiced successfully in many areas and much experience has been gained in solving problems associated with this technique. Some of the earliest work reported in the United States was done by the U. S. Geological Survey in the Grand Prairie Region, Arkansas (Sniegocki, 1959, 1963a, 1963b; Sniegocki, et al., 1965). These studies identified entrained air, turbidity and micro-organisms as major causes of plugging during injection and suggested methods to overcome these problems. In addition to these problems other technical considerations unique to injection in saline aquifers relate to the difference in water quality between the injected water and the resident aguifer water. Kimbler, et al. (1975). studied the miscible displacement and molecular diffusion processes which take place when a fluid is injected into another fluid of different composition, and defined the primary parameters that affect recovery efficiency of the injected and stored water.

Injection/recovery tests involving fresh water emplacement in saline aquifers have been carried out at a number of locations in Florida by the U. S. Geological Survey with reported recovery efficiencies varying from 0 to 47 percent. A summary of these tests is given by Merritt, Meyer, and Sonntag (in press). Preliminary field experience indicates possible problems where aquifer transmissivity is very high or very low, or when the injected water has high turbidity. Some problems in plugging of the injection well by inorganic precipitates or bacterial activity have also been indicated.

Comprehensive conceptual modeling of recovery efficiency as a function of a variety of hydraulic and water quality parameters has been completed by Merritt (in press). These studies indicated that aquifer permeability,

anisotropy, hydrodynamic dispersion, resident fluid salinity, aquifer storativity, background hydraulic gradients, length of storage period, injection and recovery schedule, wellbore and aquifer plugging and location and operation of wells in multiple-well configurations all have some effect of recovery efficiency. Effects of partial penetration of the aquifer were shown to be negligible, but aquifer stratification (vertical variations in permeability) could have significant effect on recovery of fresh water.

Preliminary hydrogeologic data used in site selection were obtained from Reece, et al. (1980); Brown and Reece (1979); and Brown (1980). Preliminary water quality information on the canals in the area was obtained from Bearden (1972), Pitt (1972), and Federico (1983).

PURPOSE. SCOPE AND METHODS

Cyclic storage of fresh water is a water resources management alternative for efficient utilization of the resource when it is quantitatively adequate but unevenly distributed timewise. Conceptually, the process consists of injection of excess surface or groundwater during periods of availability, storage of this water in the saline aquifer until it is needed, and subsequent withdrawal of the fresh water until the concentrations of critical constituents of the recovered water reach maximum permissible limits.

The purpose of this study was to provide field verification of the feasibility of using this technique in the Upper East Coast Planning Area. The Floridan Aquifer System in this area is an important source of water supply, mainly for agricultural uses. The aquifer system consists of a relatively thick sequence of limestones and dolomitic limestones, confined above by low permeability and clastic, predominantly carbonate sediments and below by dense dolomitic limestones and evaporites. Several discrete producing zones of relatively high permeability occur within this sequence. These zones are separated by less permeable dense, sandy, or chalky limestones.

Groundwater in the system is under artesian pressure. Throughout the area potentiometric heads in the Floridan Aquifer System are above land surface, and wells which tap the system are free-flowing.

Groundwater from the Floridan Aquifer System in the vicinity of the study area is of marginal to poor quality for prevailing agricultural uses (mainly citrus irrigation). Surface water supplies in the area are generally of good quality but inadequate during dry periods. Surface water impoundments as a means of regulating supply are not favored due to cost, safety, environmental and other considerations. Cyclic storage of fresh water in the Floridan Aquifer System, if successfully implemented, could serve as a mechanism for regulating the availability of water, maintaining flowing artesian heads, and improving water quality locally.

The scope of the study included preliminary site selection, drilling and testing of an exploratory well and monitor well, and a short-term injection/recovery test. Based on this study the desirability of constructing a test/injection well for long-term multiple-cycle injection/recovery tests would be evaluated.

Preliminary site selection was based on a review of available pertinent hydrogeologic, geologic, and water quality data from the Upper East Coast Planning Area. The principal criteria used for site selection were:

- 1) An adequate supply of water suitable for injection should be available conveniently and economically. Preferably the chloride content of this water should not exceed 250 mg/l during periods of injection. Suspended solids concentrations during these periods should also be sufficiently low so as not to present major problems with plugging of the rock interstices in the aquifer.
- 2) A suitable injection zone should exist in the upper part of the Floridan Aquifer System. Criteria for suitability of the injection zone include:

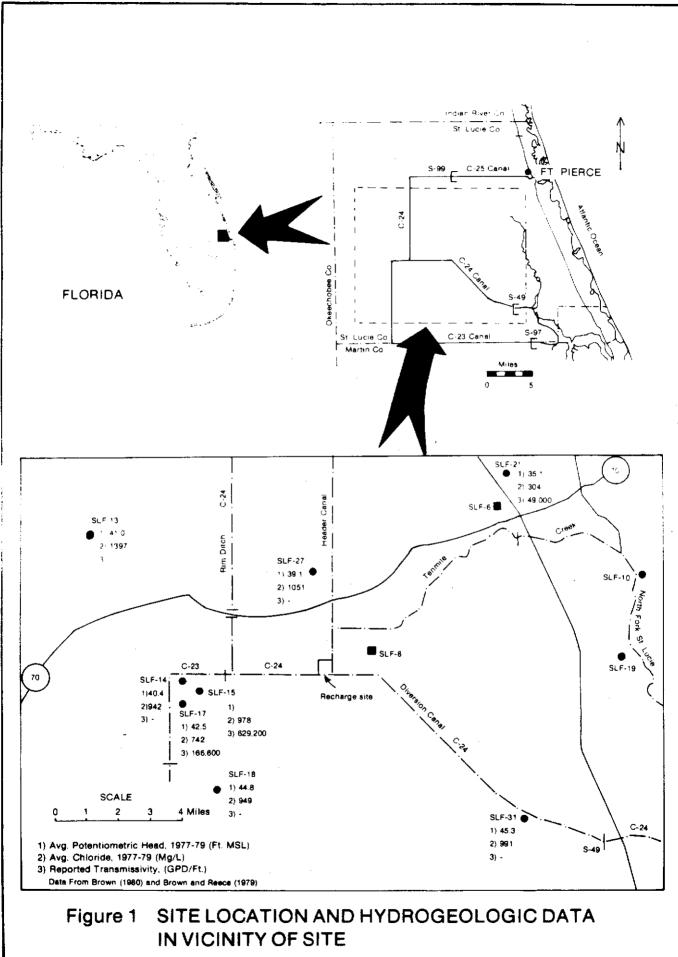
- a) Transmissivity in the range of 50,000 to 150,000 gpd/ft.
- b) Chloride content significantly higher than that of the injected water but not excessive. A range between 1000 mg/l and 5000 mg/l would be considered suitable.
- c) Consideration should be given to selecting an area with low groundwater gradients and low artesian head above land surface.

 Additionally the site should be located at least 1/4 mile from other active wells which might affect the results from tests. Other site considerations would include availability of land, convenience for access, and relevance of the site to present or potential areas of water demand.

The selected site, shown on Figure 1, is located at the northwestern corner of the intersection of Canal 24 and Header Canal, approximately 10 miles west of Fort Pierce, St. Lucie County (T 36S, R 39E, S 14DD; Latitude 27°20'17"N, Longitude 80°29'53"W). Review of available data indicated that the top of the Floridan Aquifer System occurred at approximately 500-550 feet below NGVD (National Geodetic Vertical Datum of 1929, approximately equivalent to mean sea level). The upper producing zone in this system was estimated to extend from the top of the aquifer to approximately 675 feet NGVD (Brown and Reece, 1979).

Total dissolved solids concentrations in the groundwater from this zone were reported as ranging from 1500 to 2000 mg/l. Chloride concentrations ranged from 1000 to 1200 mg/l. Aquifer transmissivities were within the range 100,000 to 500,000 gpd/ft (Brown, 1980) although these values represent contributions from more than one producing zone. Potentiometric heads were between 38 feet and 40 feet NGVD during 1977 (Brown and Reece, 1980).

Chloride concentrations in Canal 24, in the vicinity of the site, were shown by Bearden (1972) to vary between approximately 100 mg/l during high



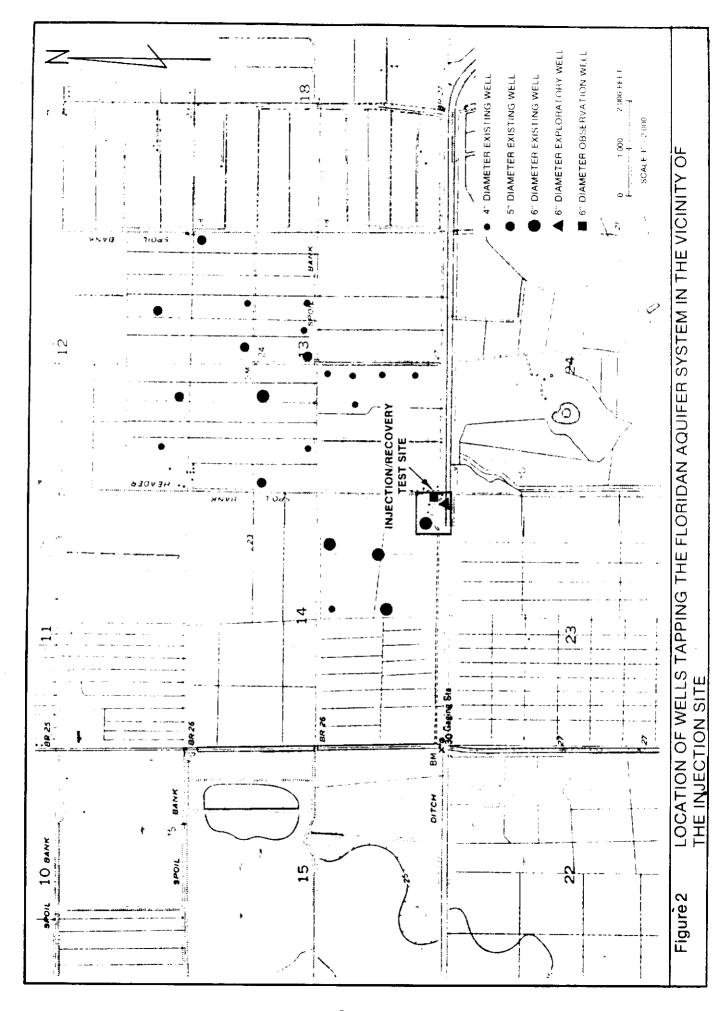
canal stage, and 350 mg/l during low canal stage. Federico (1983) reported minimum and maximum chloride concentrations in C-24 at S-49 of 152 and 534 mg/l respectively, during the period November 1, 1976 to October 31, 1977. Suspended sediment concentrations (Pitt, 1972) were between 2 mg/l during low stage and 360 mg/l during high stage. Ninety-five (95) percent of the sediment ranged in size from 0.004 millimeters (mm) to 0.062 mm.

Figure 2 shows locations of other Floridan wells in the vicinity of the site. Only one well was located within 1/4 mile of the site. This well was located in an abandoned grove, was infrequently used, and valved.

Arrangements were made to utilize this well as a monitor well during aquifer recovery tests. No wells tapping the Surficial Aquifer System were found in the immediate area.

Field work was designed to obtain data on lithology, stratigraphy, hydrostratigraphy, aquifer parameters, and water quality at the selected site. Initially, two deep wells were drilled. The first was an exploratory well (SLF-50) which was drilled to 1000 feet depth below ground level. The lower portion of this hole was later cemented back to 775 feet. The second (SLF-51) was a monitor well drilled to a depth of 775 feet, based on identification of an injection horizon above this depth from data obtained from the exploratory well.

Construction of both wells was done using similar drilling and completion methods. The Surficial Aquifer System was drilled with a 17 inch diameter bit to its base at approximately 130 feet depth, using direct mud circulation. Twelve (12) inch schedule 40 polyvinyl chloride (PVC) casing was set and cemented to this depth. Neat cement was emplaced through an 1 1/4 inch tremie pipe in the annulus between the casing and the open hole. Drilling was then continued using an 11½ inch bit, to the top of the first persistent carbonate sequence (top of the Floridan Aquifer System) at approximately 600 feet depth.

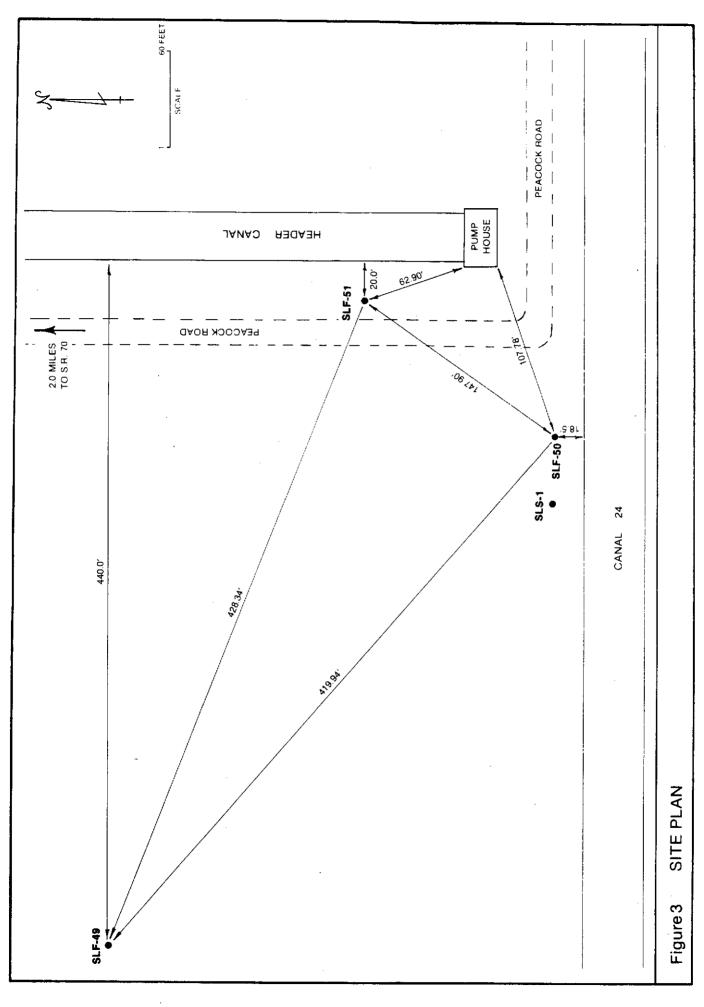


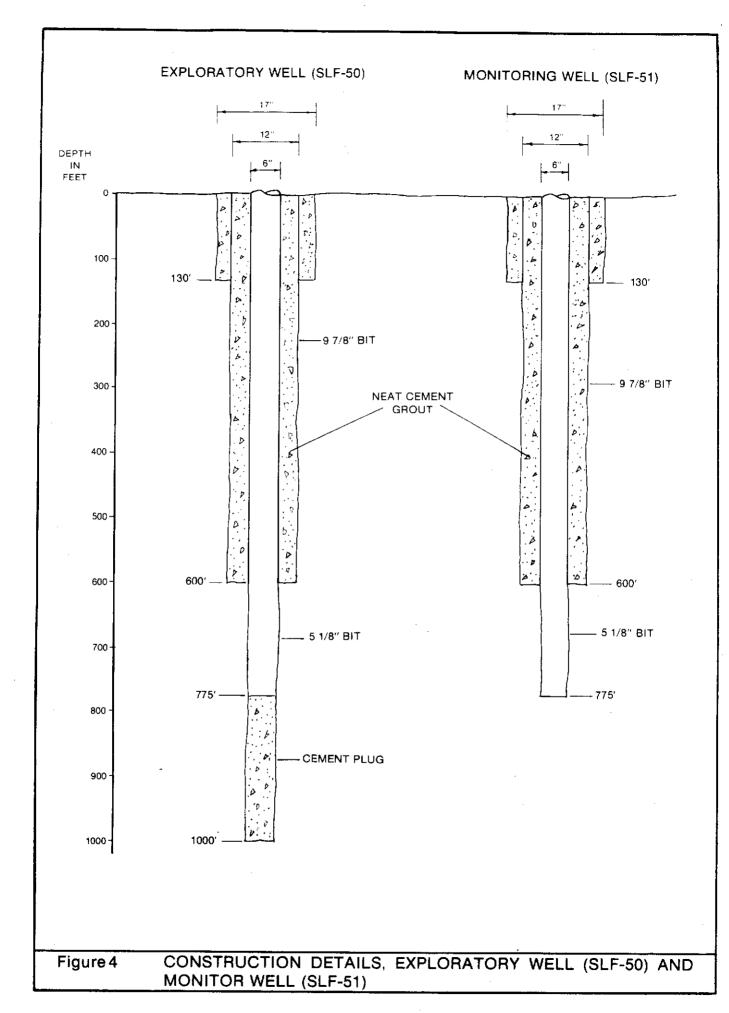
Six (6) inch Schedule 40 PVC casing was set, centralized, and pressure cement grouted. Grouting of the annular space around the 6 inch casing was carried out from inside the casing. The casing was filled with drilling fluid, and a tremie pipe was lowered to the bottom of the hole. Neat cement was pumped in with sufficient pressure to be pushed into the annular space between the nominal $11\frac{1}{2}$ inch drilled hole and the 6 inch PVC casing to displace the drilling fluid. Turbulent flow was maintained during the cementing process.

The final section of the hole was drilled with a 5 1/8 inch diameter bit, using reverse air circulation. In the exploratory well, this was done in stages with breaks at 627 feet, 747 feet and 870 feet depths below ground level to allow for testing of the well at these depths. Figures 3 and 4 show details of the site and finished dimensions of these wells.

During drilling operations cuttings were collected at 10 foot intervals. To ensure that representative samples were collected, the hole was cleared of cuttings after each 10 foot penetration by continuing circulation with the bit stationary until no further cuttings were being discharged. Geologic descriptions of these cuttings are given in Appendix 1.

During drilling of the wells geophysical surveys were run at various stages. In the exploratory well, surveys were run when the well was at 600 feet, 627 feet, 747 feet, and total depth (1000 feet). These surveys provided information which was used to determine casing settings and the termination depth for the well. In the monitor well, surveys were run at 600 feet to confirm the setting depth of the 6 inch casing and at completion of the well at 775 feet. The geophysical surveys included Spontaneous Potential, 16 inch Normal Resistivity, 64 inch Normal Resistivity, Flowmeter, Caliper, Natural Gamma, Neutron Porosity, Temperature Gradient, Differential Temperature, and Fluid Resistivity. The logs obtained from the exploratory well are shown in Appendix 2.





Water samples were collected at 20 foot intervals during reverse air drilling. Groundwater samples were also collected periodically from the overflow from the wells. After completion of drilling, point samples were collected using a geophysical logger and a sampling tube. Samples were also collected during pumping, injection/recovery, and packer tests. The samples were chilled and transported to the South Florida Water Management District Laboratory for analyses. Results of the analyses are given in Appendix 3.

Four aquifer tests were run during and after drilling of the exploratory well and before completion of the monitor well. A 5 inch centripetal pump with a 4 inch drop pipe was used to withdraw water during the tests. Water level changes with time (drawdown and recovery) were measured with an electric tape inserted between the drop pipe and the casing when water levels were below ground surface in the pumping well. Measurements were made at the observation wells using a steel tape secured to a transparent quarter-inch diameter manometer tube attached to a tap on the sealed wellhead and extended approximately 16 feet above ground surface. A manometer tube was also used to measure recovery in the pumped well. Scaffolding was erected to allow for direct reading of water levels in the manometer tube. Data and analyses from these tests are presented in Appendix 4.

Packer tests were run in the completed exploratory well to determine the degree of isolation between the upper and lower producing intervals. The tests consisted of lowering a Tamset 5 7/8 inch diameter retrievable packer through the 6 inch casing to a depth of 776 feet within the open hole and inflating the packer hydraulically to form a seal within the borehole. Water quality and water level information were collected from above and below the packer. Details of the packer test are given in Appendix 5.

After completion of the exploratory well and the monitor well, a 72-hour aquifer test was run to provide more definitive data on the aquifer parameters

of the selected injection zone. Details of this test are given in Appendix 4. During this test, water was withdrawn at a constant rate from Well SLF-51. Water level changes with time were recorded at the pumped well and at two observation wells (SLF-50 and SLF-49). Both wells SLF-50 and SLF-51 were 775 feet deep with casing to 600 feet. Well SLF-49 was 893 feet deep with casing to 560 feet.

Final testing to complete preliminary assessment of the site involved a short-term injection test. To provide water for injection a shallow 8 inch diameter well (SLS-1) was drilled close to the canal (C-24) and the exploratory well (SLF-50). This well was 55 feet deep and was cased to 35 feet. The well was designed to obtain sediment-free water from the canal through induced infiltration. This well produced approximately 800 gallons of water per minute with chloride concentration approximately 200 mg/l. Data and analyses from the injection/recovery test are given in Appendix 6.

Injection was carried out for 75 hours at rates varying from 250 gpm to 500 gpm and pressures ranging from 25 psi to 41 psi. A total of 1.48 million gallons of water was injected. The well was then shut in for 30 days, after which time, recovery was initiated. Recovery was effected by natural backflow over a period of four weeks, until withdrawn water was of approximately the same quality as resident aquifer water. Water quality data were taken at 2 hour intervals during the first part of the recovery cycle. This included onsite chloride, conductivity and temperature readings, and collection of samples for laboratory analysis. Results of all these tests are discussed in detail in subsequent sections of this report.

RESULTS OF INVESTIGATIONS

STRATIGRAPHY

INTRODUCTION

Figure 5 shows the major stratigraphic units encountered in the exploratory well (SLF-50) between land surface (+31.75 feet NGVD) and a depth of 1000 feet. The stratigraphic sequence penetrated ranged in age from Eocene to Recent. Eocene rock units include the Ocala Group and the Avon Park Limestone. The Miocene age Hawthorn Formation overlies the Ocala Group throughout the injection site area. The lowermost beds of the Hawthorn Formation are somewhat similar to lithologies that define the Tampa Formation and Suwannee Limestone (Oligocene) in other parts of Florida. However, due to the phosphatic nature of these beds they are here included in the Hawthorn Formation. Younger formations range in age from late Miocene/Pliocene to Recent and are represented by the Tamiami Formation, Anastasia Formation, and Undifferentiated terrace deposits.

CENOZOIC ERATHEM

Eocene Series

Avon Park Limestone

Applin and Applin (1944) proposed the name Avon Park Limestone to describe rocks of late Middle Eocene age in northern and peninsular Florida. They described the Avon Park Limestone as mainly a cream-colored, highly microfossiliferous chalky limestone, with gypsum and chert. The type locality is in a well in the Avon Park Bombing Range in Polk County, Florida. Where this formation crops out in Levy County, Florida it normally occurs as a tan to brown, very dense, poorly fossiliferous and massive dolomite (Knapp, 1978).

In the exploratory well the Avon Park Limestone was identified as a highly recrystallized, fossiliferous and dolomitic limestone. At total

		LITHOLOGY	DEPTH -FT 	SLF-50 COLUMN	NATURAL Gamma	HYDRO- GEOLOGIC UNIT
•⊣	IFFER-	Fine to medium grained quartz sand with varying percentages of shell. Occasionally interbedded with sandy limestone and/or shell beds.		A7666 12-2	}	SURFICIAL
ANA:		Sandy coquina of mollusk shells with a calcareous cement. Frequently interbedded with sandy limestone, calcareous sandstone and sand	100.0	\$\$\$\$\$\$		AQUIFER SYSTEM
		Sandy and biogenic limestones with minor percen- tages of sparry calcite and dolomite. Phosphate present near base			and the second s	
			200.0			HAWTHOR
			300.0	P P P P P		CONFINING
		Predominantly poorly indurated clayey, silty phosphatic sands. Sandy and phosphatic dolomites and limestones occur near the base. Phosphatic clayey dolo-silts occur frequently throughout the unit.	400.0			BEDS
			600.0			
OH FOR	IVER	Moderately indurated, biogenic, very micritic, coquinoid limestone with a fauna that includes many larger foraminitera.	700.0			FLORIDAN
1		Blogenic and medium grained calcarenitic limestone.	800.0			AQUIFER
		Highly recrystallized biogenic limestone and dolomite.	900.0		and some of the source of the	SYSTEM
	ANAS FORM TAM FORM CR R FORM AVON	ANASTASIA FORMATION TAMIAMI FORMATION CRYSTAL RIVER FORMATION WILLISTON FORMATION	ANASTASIA FORMATION HAWTHORN FORMATION Predominantly poorly indurated clayey, silty phosphatic sands. Sandy and phosphatic dolomites and limestones occur near the base Phosphatic clayey dolo-silts occur frequently throughout the unit. CRYSTAL RIVER FORMATION CRYSTAL RIVER FORMATION ANOMPARK Highly recrystalized biogenic limestone and dolomite. AVON PARK Highly recrystalized biogenic limestone and dolomite.	ANASTASIA FORMATION TAMIAMI Sandy and biogenic limestones with minor percenlages of sparry calcite and dolomite Phosphate FORMATION Predominantly poorly indurated clayey, silty phosphatic sands. Sandy and phosphatic dolomites and limestones occur near the base. Phosphatic clayey dolo-silts occur frequently libroughout the unit. TAMIAMI Sandy and biogenic limestones with minor percenlages of sparry calcite and dolomite Phosphate present near base FORMATION Predominantly poorly indurated clayey, silty phosphatic sands. Sandy and phosphatic dolomites and limestones occur near the base. Phosphatic clayey dolo-silts occur frequently libroughout the unit. 500.0 CRYSTAL RIVER FORMATION WILLISTON WILLISTON Blogenic with a fauna that includes many larger foraminitera. Blogenic and medium grained calcarenitic limestone. 800.0 AVON PARK Highly recrystalized biogenic limestone and dolomite.	ANASTASIA PORMATION TAMIAMI Sandy and biogenic limestones with minor percenlages of sparry calcite and dolomite Phosphale present near base FORMATION Predominantly poorly indurated clayey, silly phosphalic sands. Sandy and phosphalic dolomites and limestones occur near the base. Phosphalic clayey dolo-silts occur frequently throughout the unit. FORMATION CRYSTAL RIVER IN Moderately indurated. biogenic, very micritic, coquincid limestone with a fauna that includes many larger for aminitera. WILLISTON Biogenic and medium grained calcarenitic limestone. Biogenic and medium grained calcarenitic limestone and dolomites.	ANASTASIA ANASTASIA FORMATION Sandy coquency indurated with sandy lime stone: calcareous sandstone and sand limestones with minor percenlages of sparry calcite and dolomite Phosphate present near Dave FORMATION Predominantly poorly indurated clayey, silty phosphatic clayer dolo-silts occur require in base. Phosphatic clayer dolo-silts occur requently throughout the unit. ANON PARK Highly recrystalitzed biogenic limestones and dolomite. Biogenic and medium grained calcarenitic limestone. Biogenic and medium grained calcarenitic limestone and dolomite.

depth, the well had penetrated 240 feet of these carbonates. A distinct lithologic change marks the boundary between the fossiliferous and dolomitic limestones of the Avon Park Limestone and the overlying calcarenitic limestones in the Williston Formation of the Ocala Group. Fossils present in the Avon Park Limestone include echinoids, foraminifera (especially <u>Dictyoconus sp.</u>), mollusks, bryozoans, and other fossil fragments. The Natural Gamma Ray log (Figure 5) shows an increase in radioactivity in this unit as compared to the overlying Ocala Group. Ocala Group

The term "Ocala Limestone" was first used by Dall and Harris (1892) in a discussion of limestones being quarried near the town of Ocala in Marion County, Florida. Applin and Applin (1944) identified an upper and lower member within the "Ocala Limestone." This usage is still followed by the U. S. Geological Survey and many others. Puri (1953) followed Vernon (1951) in recognizing three distinct units that he believed were present within the strata of the "Ocala Limestone." He proposed for these units the names Crystal River Formation, Williston Formation, and Inglis Formation (in descending order of depth) and suggested that the new formations should be included in the Ocala Group. This usage is followed in this report.

The Crystal River and Williston Formations were both penetrated in the test wells. The Inglis Formation was not recognized in the area. The Crystal River Formation occurred as a coquina of larger foraminifera (Lepidocyclina sp.) in a carbonate mud matrix. The Williston Formation is a medium grained calcarenitic limetone. The unit is fossiliferous, but not coquinoid. The Ocala Group is 100 feet thick in well SLF-50. Only 20 feet of the Williston Formation was encountered in this well.

Miocene Series

Hawthorn Formation

The Hawthorn Formation overlies the Crystal River Formation at the test site area. The lowermost beds of this unit may be equivalent to the Suwannee Limestone, Tampa Formation, or Unnamed Limestone of Mooney (1980). Dall and Harris (1892) first used the term "Hawthorne beds" for phosphatic sediments being quarried for fertilizer near the town of Hawthorne, Alachua County, Florida. In recent work by Scott and Knapp (1983), continuous cores were described and correlated with the type and co-type sections. They described the Hawthorn Formation as "consisting of varous mixtures of clay, quartz sand, carbonate (dolomite to limestone), and phosphates." They also divided the Formation into two distinct units in southern peninsula Florida, a lower predominantly carbonate unit and an upper predominantly clastic unit.

As recognized in this report, the Hawthorn Formation is a heterogeneous sequence of phosphatic, sandy, clayey, calcareous, and dolomitic sediments of which the uppermost bed is an olive-gray, clayey phosphatic sand. The lowermost beds are very fine grained biogenic and slightly phosphatic limestones, which may in part be of Oligocene age (Armstrong, 1981). The Hawthorn Formation exhibits high gamma activity (Figure 5) due mostly to the presence of phosphatic sand. The total thickness of the Hawthorn Formation at the test site was 530 feet. The basal limestones were 60 feet thick.

Miocene/Pliocene Series

Tamiami Formation

Mansfield (1939) proposed the name "Tamiami Limestone" for a fossiliferous sandy limestone approximately 25 feet thick, which was penetrated in shallow ditches along the Tamiami Trail (U.S. Route 41) in

parts of Collier and Monroe Counties, Florida. At the exploratory site the Tamiami Formation was recognized as a sandy, very fine grained limestone with very little phosphate. It was 30 feet thick in this well and extended from 100 feet to 130 feet below land surface.

Pleistocene/Holocene Series

Anastasia Formation

Sellards (1912) used the name Anastasia Formation for exposures of coquina rock extending southward along the Atlantic coast of Florida from St. Augustine. The rock is composed of sandy limestone, calcareous sandstone, and unconsolidated sand and shell (Puri and Vernon, 1964). In Well SLF-50 the Anastasia Formation occurs as a calcareous sandstone and shell bed approximately 70 feet thick between 30 and 100 foot depths below land surface.

<u>Undifferentiated Terrace Deposits</u>

The surficial deposits at the test site are composed of medium-grained sand, shell and calcareous clayey material about 30 feet thick. Some of the sediment penetrated by the exploratory well was evidently spoil material from the adjacent drainage canal.

HYDROSTRATIGRAPHY

INTRODUCTION

Three major hydrostratigraphic units were identified at the test site, extending from land surface (elevation +31.75 NGVD) to a depth of 1000 feet. These are the Surficial Aquifer System, Hawthorn Confining beds, and the Floridan Aquifer System. The Surficial Aquifer System was not examined in detail due to the scope of this project. The Hawthorn Confining beds have an overall low permeability and effectively separate the Surficial Aquifer System from the much deeper Floridan Aquifer System. The Floridan Aquifer System was

found to contain multiple producing zones that correlated closely with highly porous beds identified from well cuttings and geophysical logs.

SURFICIAL AQUIFER SYSTEM

The Surficial Aquifer System is 130 feet thick at the test site and includes sediments associated with the Tamiami Formation, Anastasia Formation, and Undifferentiated deposits. It is underlain by the less permeable clayey sands and phosphatic dolo-silts of the Hawthorn Formation. Some of the individual limestone and sandstone beds within this system exhibit high permeability. A gray calcareous sandstone was penetrated between 30 and 50 feet which appeared to have high permeability. A shallow well completed subsequently in this zone produced in excess of 800 gallons per minute. The associated porosities are commonly of the intergranular, moldic, and intercrystalline varieties. The presence of shell beds with high intergranular porosities also contribute to the high permeability of this unit.

HAWTHORN CONFINING BEDS

The Hawthorn Confining beds are equated with the upper portion of the Hawthorn Formation and are contained wholly within that formation. The lowermost limestone beds of the Hawthorn Formation which are in hydraulic connection with the underlying Eocene sediments are considered part of the Floridan Aquifer System. The Hawthorn Confining beds are composed principally of phosphatic, clayey, dolomitic quartz sands. The presence of interbedded dolo-silts and the clayey matrix of the quartz sands give this unit an overall low permeability. The lithology of this unit is, however, not uniform and there are several well indurated porous limestone and dolomite beds within it, such as at 380 to 390 feet, 440 to 450 feet, and 540 to 550 feet. These beds may be capable of producing small quantities of water, but due to their relative thinness and apparent low permeability they are not considered a

significant water supply source. These zones, along with the less permeable sandy and silty zones in the Hawthorn Formation, are generally cased off in wells which tap the Floridan Aquifer System. The base of the Hawthorn Confining beds is marked by a green clayey dolo-silt which extends from 580 feet to 600 feet. At the exploratory well site the Hawthorn Confining beds are 470 feet thick, and the top was logged at a depth of 130 feet.

FLORIDAN AQUIFER SYSTEM

The term "Floridan Aquifer" was established by Parker (Parker, et al., 1955) to describe water bearing rocks associated with the Lake City Limestone, Avon Park Limestone, Ocala Limestone, Suwannee Limestone, Tampa Limestone, and the lower permeable parts of the Hawthorn Formation which are in hydrologic contact with the underlying units. In describing this unit in southern Florida, Brown (1980) and Wedderburn, et al. (1982), used the term "Floridan Aquifer System" in recognition of the presence of multiple producing zones and confining zones within the sequence of rocks. This terminology is used in this report.

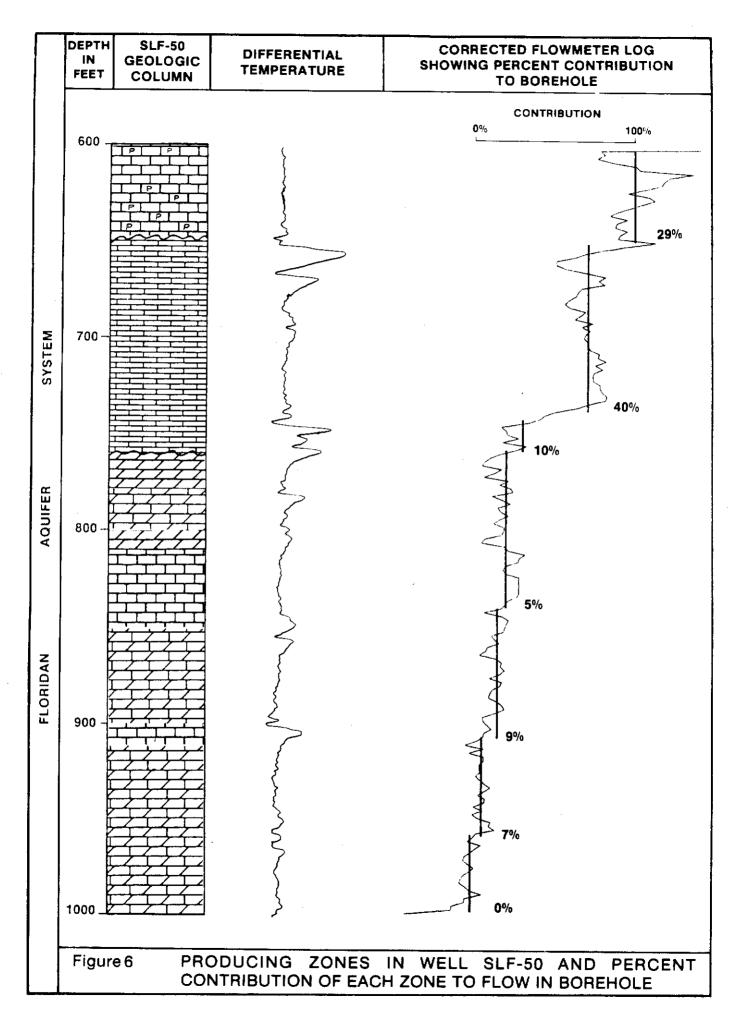
At the exploratory well site, the Floridan Aquifer System consists of a thick sequence of interbedded limestones and dolomites of Eocene to lower Miocene age extending from a depth of 600 feet to the bottom of the well (1000 feet). The well did not, however, penetrate the full thickness of the system. Previous work indicates that this aquifer system is areally extensive throughout south Florida and may be as much as 3000 feet in thickness in this area (Miller, 1982).

The top 60 feet of the system occurs within the Hawthorn Formation and consists of well-indurated fossiliferous limestones exhibiting moderate (approximately 15 percent) moldic and intergranular porosity. The base of this interval is marked by a pronounced gamma ray "kick." This section of the

Hawthorn Formation corresponds to the "Unnamed Limestone" identified by Mooney (1980) as Oligocene in age.

The coquinoid limestone beds of the Ocala Group and the highly recrystallized limestone and dolomite beds within the Avon Park Limestone are also water producing within the Floridan Aquifer System. Geophysical logs indicate that the major water production is confined to relatively narrow intervals which tend to correspond to stratigraphic contacts. Figure 6 shows Differential Temperature Log. Corrected Flowmeter Logs and lithology within the upper part of the Floridan Aguifer System. The Differential Temperature Log shows a number of peaks which indicate significant temperature differences within the system. These peaks generally correspond with zones of flow within the wellbore as detected by the Flowmeter Log. Six zones which contribute significant flow to the wellbore under natural flowing artesian conditions are indicated on the Flowmeter Log which has been corrected to minimize the effects of variations in borehole diameter. The first significant producing zone in the Floridan Aquifer System occurs between approximately 650 to 670 feet, and accounts for about 29 percent of the flow of the well. The most productive zone occurs between of 730 and 750 feet and accounts for about 40 percent of flow. A third zone which accounts for about 10 percent of flow occurs between 760 and 770 feet. These producing zones occur approximately at the contact between the Ocala Group and the Hawthorn Formation and Avon Park Limestone respectively, and may represent reworking or dissolution at the unconformities which mark depositional breaks in the strata.

The lower portion of the borehole below the third producing zone contributes less than 21 percent of the total natural flow from the well. Minor flow contributions are indicated at a depth of 840 to 842 feet (5 percent) and at a depth of 895 to 905 feet (9 percent). Between 955 and 960 feet some 7 percent of the flow from the well is produced.



Based on the above information, the interval above 775 feet which contains the three major producing zones appeared to offer the best promise for injection of fresh water. The first two producing zones are each approximately each 20 feet thick and the third about 10 feet thick. These zones show a mixture of fracture, solution, and moldic porosity, and flow within then could be expected to differ significantly from isotropic, homogeneous porous media flow. Loss of injected water from this interval to the lower strata would be expected to be relatively small due to the much higher horizontal permeability in the injection horizon as compared to the permeability below this horizon.

AQUIFER PARAMETERS

The aquifer parameters of particular importance in assessing the feasibility of injecting fresh water into saline aquifers are permeability, porosity, longitudinal and vertical dispersivity and the coefficient of molecular diffusion (Khanal, 1980). A series of aquifer tests were designed to provide information on the transmissivity and permeability of the producing zones in the upper part of the aquifer system. Porosity was estimated from examination of the rock cuttings. The scope of the study did not allow for any tests to determine the longitudinal or vertical dispersivity or coefficient of molecular diffusion. These parameters were estimated based on experience in similar lithologies from tests conducted by other researchers.

The first set of 4 aquifer tests was performed on the exploratory well, to aid in selecting a suitable injection zone. Drilling of this well was halted at various depths to allow for the tests to be run. The well was developed, allowed to recover fully, then pumped at a constant rate. Data on water level changes with time were collected from the pumped well (SLF-50) and from an existing irrigation well (SLF-49). The Jacob semi-logarithmic method (Jacob, 1952) was used to analyze the recovery data for aquifer transmissivity. This method assumes interalia that the aquifer tested is

isotropic, homogeneous, fully penetrated and fully confined. Since these assumptions were not expected to be fully met, the values derived from this analysis were expected to be approximate, and were utilized only as indicators of the relative transmissivities of different sections of the strata penetrated.

A fifth pumping test was run after completion of both the exploratory well and the monitoring well. This test was designed to provide more exact information on aquifer parameters in the selected injection horizon. The data from this test were analyzed by the Hantush/Jacob type curve method and the Hantush-Inflection point method (Hantush and Jacob, 1955; Hantush, 1964).

Table 1 summarizes the results of these pumping tests. The four initial pumping tests indicated a general trend toward increased transmissivity with depth of penetration of the aquifer. The first zone tested included the interval from the bottom of the casing (600 feet) to a depth of 627 feet. At this site the calculated transmissivity of this interval, which occurs at the very top of the Floridan Aquifer System, was low (9,428 gpd/ft). The second interval tested included all of the aquifer between 600 feet and 747 feet. The test indicated a significant increase in transmissivity (65,340 gpd/ft). At a depth of 870 feet the well was again tested, and the results indicated a further increase in calculated transmissivity (88,000 gpd/ft). The final preliminary test was run with the well at 1000 feet depth. This test yielded a transmissivity of 107,077 gpd/ft.

Based on the results of these tests and additional lithologic and geophysical information previously discussed, the interval from 600 feet to 775 feet was selected as a suitable injection horizon. This interval was isolated from the lower portion of the borehole using a Tamset retrievable/resettable packer set at 776 feet. The packer tests confirmed the presence of a major producing interval above the setting depth, and indicated

TABLE 1. SUMMARY OF PUMPING TEST DATA AND RESULTS

*Preliminary tests to determine relative transmissivity with depth. Values do not reflect true transmissivity since the effects of leakance are ignored. **Values underlined reflect true aquifer characteristics of injection horizon.

relatively good confinement between this zone and the lower strata (see Appendix 5 for details of the Packer test). The exploratory well was then cemented back to 775 feet and a monitoring well constructed to enable monitoring of the zone between 600 feet and 775 feet depth.

The final pumping test was designed to provide definitive information on the aquifer parameters of the proposed injection horizon (600 feet to 775 feet). Well SLF-51 was pumped 72 hours at a constant discharge rate of approximately 388 gpm. Analysis of the data gave a transmissivity of 43,000 to 48,000 gpd/ft; a storage coefficient of 1.6 X 10-4 to 2.6 X 10-4, and leakance coefficient of .043 to .047 day-1. The lower value of transmissivity, compared to that obtained during the initial testing with the well at approximately the same depth (65,000 gpd/ft), is probably due to the effects of leakance which could not be addressed in the initial analysis due to the absence of a suitable observation well.

Porosity values were visually estimated as between 14 to 18 percent. The dominant types of porosity seen in the well cuttings were intergranular and moldic. No tests were run to determine the hydrodynamic dispersion parameters. Hydrodynamic dispersion refers to the spreading and mixing caused in part by molecular diffusion and microscopic variation in velocities within individual pores (Anderson, 1979). For regional applications, values of longitudinal dispersivity coefficient have been found to vary from tens to hundreds of meters. Segol and Pinder (1976) reported values of 6.7 meters and 0.7 meters respectively for longitudinal and transverse dispersivities in limestones in southeast Florida. Values of longitudinal dispersivity ranging from 3.1 meters to 61 meters have been reported for limestone aquifers (see Mercer et al., 1982). However, values ranging from 6.7 to 13.7 meters appear to match lithologic and hydrostratigraphic conditions similar to those in the study area.

WATER QUALITY

The water quality sampling program was designed to provide information on: a) the suitability of the available surface water for injection; b) water quality variations in the aquifer to assist in selection of an injection zone; and c) provide background data for use during injection/recovery tests. Water samples were collected using standard techniques and analyzed by the SFWMD Laboratories. Results of these analyses are tabulated in Appendix 3.

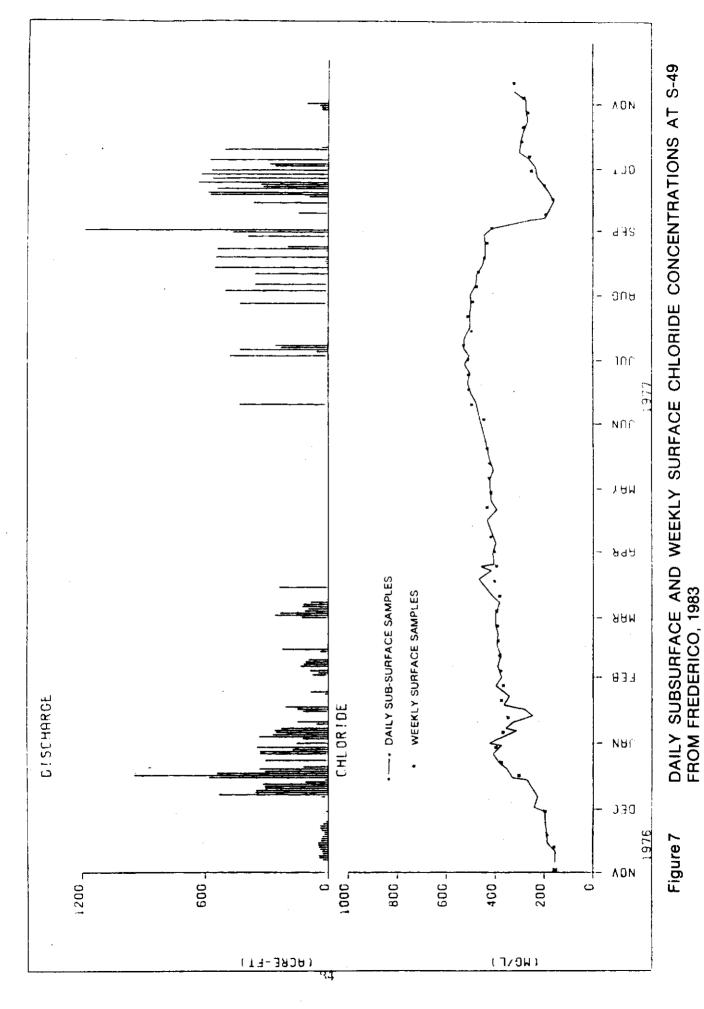
SURFACE WATER QUALITY

Data on chloride concentrations in surface water in the vicinity of the site prior to initiation of the exploratory program have been published by Pitt (1972) and Bearden (1972). Table 2 presents selected data on chloride concentrations in Canal 24, in the vicinity of the site, and associated water levels at Structure 49 (S-49), 11 miles downstream of the investigation site.

TABLE 2. Approximate Chloride Concentrations in C-24 Near Site, and Canal Levels at S-49 (From Bearden, 1972).

DATE	WATER LEVEL AT S-49, UPSTREAM (FT)	CHLORIDE CONCENTRATION IN C-24 NEAR SITE (mg/l)
09/21/67	18.82	150
04/02/68	17.15	360
05/01/68	15.10	250
10/08/68	18.95	150

The data in Table 2 indicate that chloride concentrations vary depending on canal levels; the minimum concentration recorded being 150 mg/l during a period of high water levels and the maximum being 360 mg/l during a period of low water levels. Data from Frederico (1983) (Figure 7) indicate that during 1976/1977 lowest chloride concentrations were found in the canal during the months of November and December 1976 and September, October and November 1977.

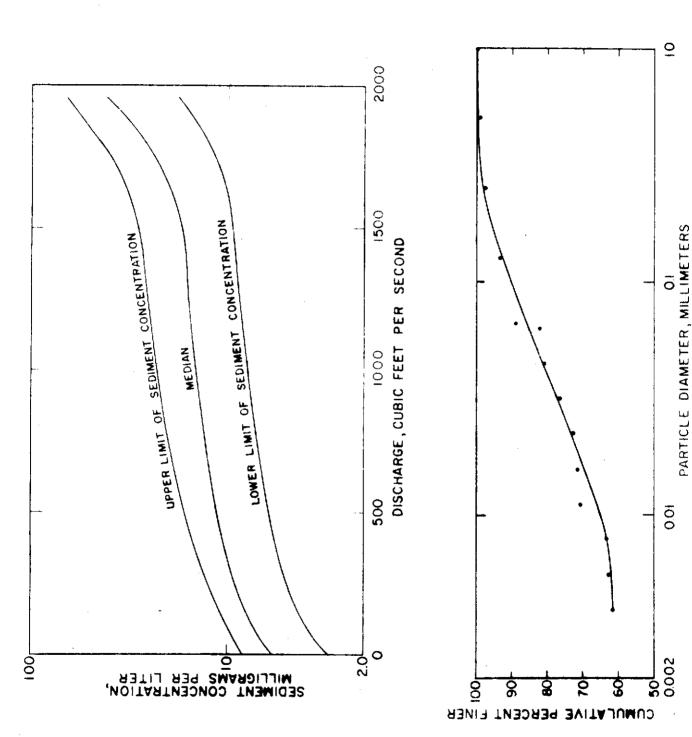


Some previous data on suspended sediment concentrations and particle size distribution in water from C-24 were reported by Pitt (1972). Sediment samples were taken at three different points across the canal 600 feet upstream of S-49 (approximately 11 miles downstream of the investigation site) during the period July to November 1969. Sediment concentrations were found to be the same at all three points. A direct relationship was found between discharge and sediment load. The sediment concentration was found to range from 2 to 360 mg/l during this period. Particle size of the sediment was mostly in the clay size range (.004 to .062 mm) (see Figure 8). According to Pitt, sediment in the canals originates in the cultivated areas from decomposition of vegetal material and soil erosion.

During the present study, water samples were collected from C-24 and Header Canal in the vicinity of the investigation site. All the samples were collected within six feet of the canal bank and less than 1 foot below the water surface. Measured chloride values in C-24 range from 555 mg/l to 107 mg/l. The highest value was recorded during discharge of saline water to the canal. The lowest value was recorded after a period of heavy rainfall (see Appendix 3). On the basis of the historic data and the analyses done during the investigations it was concluded that during the wet season the chloride concentrations in the canal could generally be expected to be below 200 mg/l. Samples taken from Header Canal show higher chloride values than those from C-24 and it was concluded that this canal would be a less suitable source of injection water.

Total Dissolved Solids concentrations in C-24 ranged from 520 mg/l to 1549 mg/l. The lowest value of 520 mg/l is assumed to be representative of uncontaminated surface water. The higher values reflect contamination of the canal by saline groundwater during drilling and aquifer testing. Dissolved iron concentrations in canal waters ranged from 0.41 to 0.44 mg/l. Sulfate

SEDIMENT CONCENTRATION AND PARTICLE SIZE, CANAL 24 AT S-49 (FROM PITT, 1972) Figure 8



concentrations ranged from 42 to 179 mg/l. The higher sulfate values probably reflected contamination of the surface water by water from the Floridan Aguifer System.

In general, available water quality data indicate that water from C-24 was usable for recharge to the Floridan aquifer. However, water with chloride concentrations less than 250 mg/l would probably be available only during the late part of the year (September to December). Treatment of the surface water might prove to be advisable to mitigate problems with clogging. Since the required treatment would represent a significant part of the overall cost of the injection program, an alternate source of water was investigated.

SURFICIAL AQUIFER SYSTEM WATER QUALITY

A shallow well was drilled close to the canal in an attempt to induce infiltration from the canal and reduce suspended solids through the natural filtration by the strata. Examination of the lithologic data from the exploratory well showed that the surficial sediment consisted of a shallow, fine to coarse grained unconsolidated sand from land surface to 10 feet, a thin low permeability micritic limestone bed from 10 to 20 feet, a shell bed with micrite cement from 20 to 30 feet, and a well indurated calcareous sandstone between 30 and 50 feet. The sandstone showed well developed solution features and high moldic porosity, and was therefore assumed to have high permeability. This zone was targeted as a suitable source for withdrawing water for injection. The depth of the canal is about 20 feet (ten feet above the top of the sandstone), and it was therefore expected that some induced infiltration from the canal to the sandstone layer would occur.

The shallow well (SLS-1) was drilled to 55 feet and cased to 34 feet.

The well was developed for two days, until the discharged water was sediment free, and tested at 800 gpm. Average chloride content of water from this well was 200 mg/l; average total dissolved solids content was 1000 mg/l; average

total iron concentration was 5 mg/l; and average sulfate concentration was 200 mg/l.

FLORIDAN AQUIFER SYSTEM WATER QUALITY

As part of the exploratory program water samples were collected during drilling, logging, and pump testing. The sampling program was designed to identify variations in water quality between the producing zones in the aquifer system. However, because of the upward flow of water within the well bore, most of the sampling methods produced a mixture of waters from different zones.

Samples collected during reverse air drilling were expected to be most representative of water quality with depth. In this drilling technique air is injected into the hole during drilling and aquifer water and cuttings are discharged through the drill pipe. Water samples taken from this discharge are generally representative of water in the aquifer at the depth of penetration. Depth samples collected with a point sampler, on the other hand, were mixtures of waters from below the sampler, and except for samples taken near the bottom of the well, would not be representative of water quality in the aquifer at that depth. Samples taken during the aquifer tests consisted of a mixture of waters from all the zones penetrated by the well, with some water which may have leaked into the well from below due to the lowered hydraulic head during pumping.

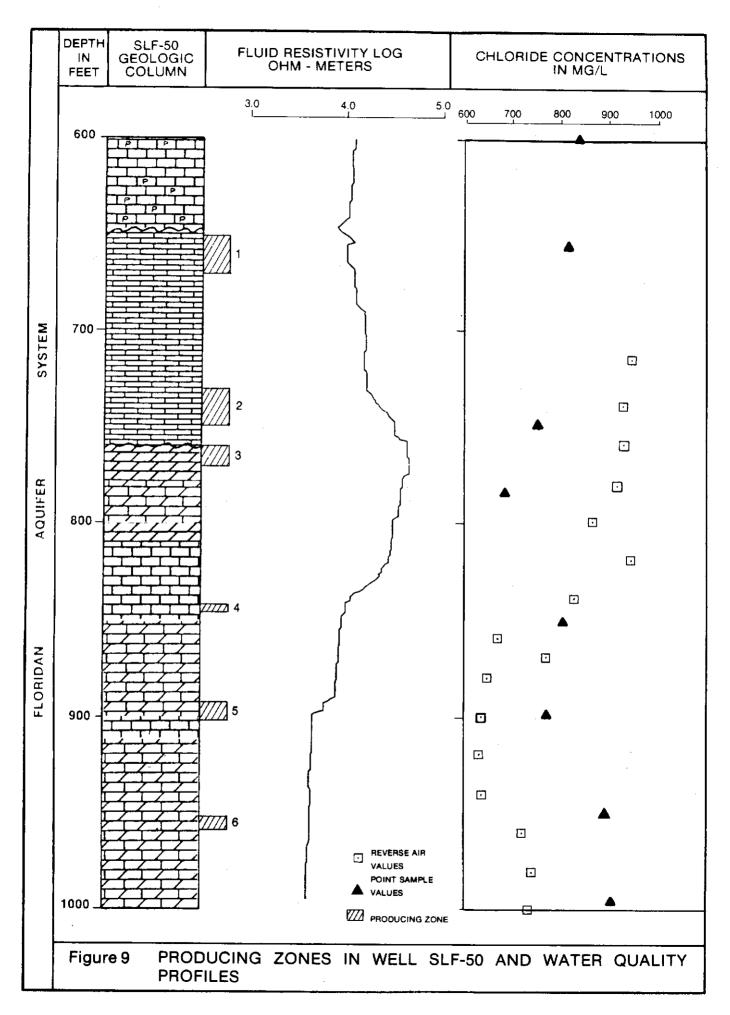
Additional qualitative insight into the variations in water quality with depth was obtained from geophysical logs. Fluid Resistivity Logs run in the flowing well record the resistivity of the mixture of waters below the probe, and would not give an accurate indication of the resistivity of the water in the aquifer at a given depth except near the bottom of the hole. However, variations in the Fluid Resistivity Logs can generally be used to identify the approximate locations of zones of different water quality. This is also true

of the 16 inch and 64 inch Normal Resistivity Logs if variations due to lithologic changes can be screened out or neglected.

All the sampling methods and the geophysical logs indicated non-uniformity in water quality with depth. The general picture shown is one of higher mineralization of the water in the upper part of the uncased borehole, a decrease in mineralization in the mid section and a slight increase in mineralization toward the base of the borehole. The most significant variations are found in temperature and chloride content. The chloride content variations are reflected in variations in conductivity and total dissolved solids. Temperature variations as shown on the Temperature Log (Appendix 2) were generally small (less than 3°F), although variations in water temperature between different producing zones were of sufficient magnitude to be used to identify these zones.

Water quality parameters for each producing zone could not be uniquely determined from any one set of samples due to inevitable mixture of waters during sampling. Based on samples taken during the first two pumping tests with the exploratory well at depths of 627 and 747 feet, it is assumed that the upper two producing zones have similar water quality. Chloride values from these tests varied between 974 and 1048 mg/l and total dissolved solids concentrations varied from 1791 to 2167 mg/l (see Appendix 3). Water quality from producing zone 6 can be determined from point sample values since the upward flow of water would prevent mixture with water above the sampler. Point sample values at depths of 950 and 995 feet give chloride concentrations of 898 and 882 mg/l, respectively, and total dissolved solids concentrations of 2220 and 2200 mg/l, respectively.

Based on the Fluid Resistivity Log (Figure 9), zones 4 and 5 would be expected to have relatively good quality water as shown by pronounced increases in resistivity adjacent to these zones. Similarly, producing zone 3



appears to have relatively good quality water compared to zones 4 and 5 as evidenced by a further increase in fluid resistivity adjacent to this zone. Point sample values at depths between 750 feet and 900 feet, which includes producing zones 3, 4, and 5, indicate that chloride concentrations in these zones is below 800 mg/l.

Compared to water from C-24 and the shallow groundwater, water from the Floridan Aquifer System is high in calcium, sodium, chloride, magnesium, sulfate, and bicarbonate, but low in iron, phosphate, and nitrate. Chloride concentrations in the aquifer are typically more than four times as high as in available recharge water. The level of mineralization of the aquifer water is moderate, which will be a favorable factor in the recovery of injected water. Differences in water quality with depth are relatively small and are not expected to greatly influence recovery efficiency.

FLORIDAN AQUIFER SYSTEM WATER LEVELS

Three factors related to water levels in the aquifer are of importance in cyclic injection and recovery. These are: a) horizontal groundwater gradient, b) vertical groundwater gradient, and c) potentiometric head at the injection well. The horizontal groundwater gradient determines the rate and direction of groundwater flow, assuming aquifer homogeneity and isotropy. During cyclic injection, a steep groundwater gradient can move the slug of injected water down-gradient, thus reducing the amount of water which can be recovered from this well. The vertical groundwater gradient determines the potential direction of leakance between permeable strata. The potentiometric head at the injection well partly determines the wellhead pressure required to inject water into the well. Wellhead pressure is also a function of well construction (head loss due to casing and open hole diameter and roughness), aquifer transmissivity, fluid viscosity, and well losses due to plugging of the aquifer adjacent to the open borehole.

The regional horizontal groundwater gradient in the vicinity of the test site is approximately of 0.5 foot per mile in a northeasterly direction.

Table 3 shows potentiometric heads in wells SLF-50, SLF-51, and SLF-49 for May and September 1982. The measurements indicate that the pressure head above the wellhead at SLF-50 averages about 9.70 feet of water (4.21 psi).

TABLE 3. POTENTIOMETRIC HEADS AT INJECTION SITE

WELL NO.	DEPTH (FT)	WELLHEAD ELEVATION FT NGVD	POTENTIOMETRIC 5/20/82	LEVEL, FT NGVD 9/22/82
SLF-49	893	25.09	41.04	43.76
SLF-50	775	31.75	40.47	42.44
SLF-51	775	26.56	40.18	-

The difference between potentiometric levels in May (dry season) and September (wet season) was 1.97 feet at this well. Data obtained during the packer test indicated that the zone between 775 feet and 1000 feet had a potentiometric head +0.24 feet higher than the zone between 600 feet and 775 feet.

Figure 10 shows potentiometric level variations at well SLF-50 during the period July 20 to August 20. The hydrograph indicates both short-term and longer-term cyclic variations in water levels due to barometric and tidal effects. These variations could introduce small inaccuracies in head measurements during pumping tests, but would have negligible effect on the injection/recovery process.

HYDROGEOLOGIC MODEL OF THE INJECTION ZONE

A preliminary conceptual model of the injection zone was developed to serve as a basis for computer simulation of the injection/recovery process in the aquifer. The model is based on data developed during exploratory work and incorporates the variations in hydrogeologic properties and water quality previously discussed. The model is illustrated in Table 4.

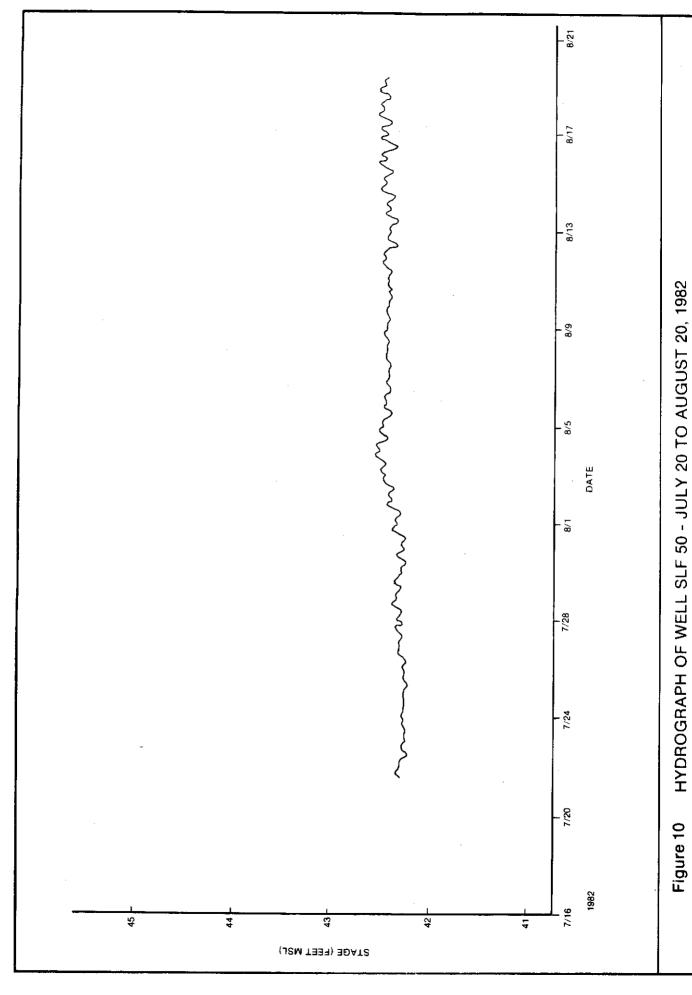


TABLE 4. PRELIMINARY CONCEPTUAL MODEL OF INJECTION HORIZON

TDS L/pm		2050	1900	1900	1900	1900	1500	1600–1900
CJ mg/l		1025	~950	~950	~955	~950	750	800-950 16
π (ft)		50	2	09	20	10	10	20 8
Κ _ν (ft/d)		3.15	103.45	3.15	142.73	3.15	71.36	3.15
K _H (ft/d)		3,15	103.45	3.15	142.73	3,15	71.36	3.15
DESCRIPTION	Non leaky upper confining bed overlying injection horizon	Low permeability zone in the injection horizon	High permeability zone in the injection horizon (Prod. Zone 1)	Low permeability zone in the injection horizon	High permeability zone in the injection horizon (Prod. Zone 2)	Low permeability zone in the injection horizon	High permeability zone in the injection horizon (Prod. Zone 3)	Low permeability semi confining zone beneath the injection horizon
DEPTH (ft)	130-600	600-650	650-670	670-730	730-750	750-760	760-770	770-840
LAYER NO.	ſ	⊷	2	m	4	ĸ	9	1

The injection horizon extends from a depth of 600 feet to a depth of 770 feet. It is bounded above by an essentially non-leaky confining bed, and below by a leaky confining bed with a coefficient of leakance of $.045 \, \mathrm{day}^{-1}$. The basal confining bed is 70 feet thick, thus the vertical permeability of this bed is 3.15 ft/day. This value corresponds closely to average values of permeabilities typical of fine-grained limestones (Bouwer, 1978).

The injection horizon consists of three layers of relatively high permeability (producing zones) and 3 layers of low permeability. The high permeability layers correspond to the producing zones identified on Figure 9. The upper two producing zones are each 20 feet thick (650 to 670 feet and 730 to 750 feet) while the lowermost producing zone is 10 feet thick (750 to 760 feet). Relative permeabilities of the producing zones were calculated based on the following assumptions.

- The intervening low permeability zones have uniform permeability of
 3.15 ft/day.
- 2) The relative permeabilities of the producing zones correspond to their relative flow contributions to the borehole.

The transmissivity of the low permeability beds to the average transmissivity of the injection horizon can be calculated using the equation.

$$T_L = K_L M_L = 3.15 \times 120 = 378 \text{ ft}^2/\text{d}$$
 Where,

 T_L = Average transmissivity of low permeability beds.

 K_L = Permeability of low permeability beds (3.15 ft/d).

 M_L = Total thickness of low permeability beds (120 ft).

Consequently the contribution of the producing zones to the transmissivity of the horizon is:

$$T_P = T - T_L = 6016 - 378 = 5638 \text{ ft}^2/\text{d}.$$

Where,

Tp = Transmissivity of producing zones.

Similarly, the permeabilities of the individual producing zones can be calculated as follows:

$$K_{P1} = \frac{T_H}{M_{P1}} X \frac{Q_{H1}}{Q_T} = \frac{5638}{20} X \frac{29}{79} = 103.45 \text{ ft/d}$$

$$K_{P2} = \frac{T_H}{M_{P2}} \times \frac{Q_{H2}}{Q_T} = \frac{5638}{20} \times \frac{40}{79} = 142.73 \text{ ft/d}$$

$$K_{P3} = \frac{T_H}{M_{P3}} \frac{X}{Q_T} = \frac{5638}{10} \frac{X}{79} = 71.36 \text{ ft/d}$$

Where,

Kp1, Kp2, Kp3 = Permeabilities of producing zones
1, 2, and 3.

Tp = Transmissivity of producing zones $(5638 \text{ ft}^2/\text{d})$.

Mp1, Mp2, Mp3, = Thicknesses of producing zones 1, 2, and 3 (20 ft, 20 ft and 10 ft, respectively).

Qp₁, Qp₂, Qp₃ = Percent contribution to flow of producing zones 1, 2, and 3 (29%, 40%, and 10%, respectively).

QT = Total percent contribution of producing zones (79%).

For the purpose of this model it is assumed that within each zone, vertical and horizontal permeability values are the same. This assumption is based on the observed essentially isotropic distribution of porosity within these zones.

Water quality parameter values were assigned to the various layers as follows. The first low permeability zone within the injection horizon was assumed to have relatively high mineralization as indicated by water samples taken during the first pumping test (well depth 647 feet) (Chlorides approximately 1025 mg/l, TDS approximately 2050 mg/l). The upper two producing zones and the intervening low permeability zones were assumed to have similar water quality and to be somewhat less mineralized. Based on

reverse air samples taken at depths of 715 feet and 740 feet, the chloride concentration in these layers averaged about 900 to 950 mg/l and total dissolved solids averaged 1800 to 1950 mg/l. The lowermost producing zone and the semi-confining bed below the injection horizon was assumed to have an average chloride concentration of 700 to 800 mg/l and a total dissolved solids concentration of 1600 to 1750 mg/l, based on values from point samples.

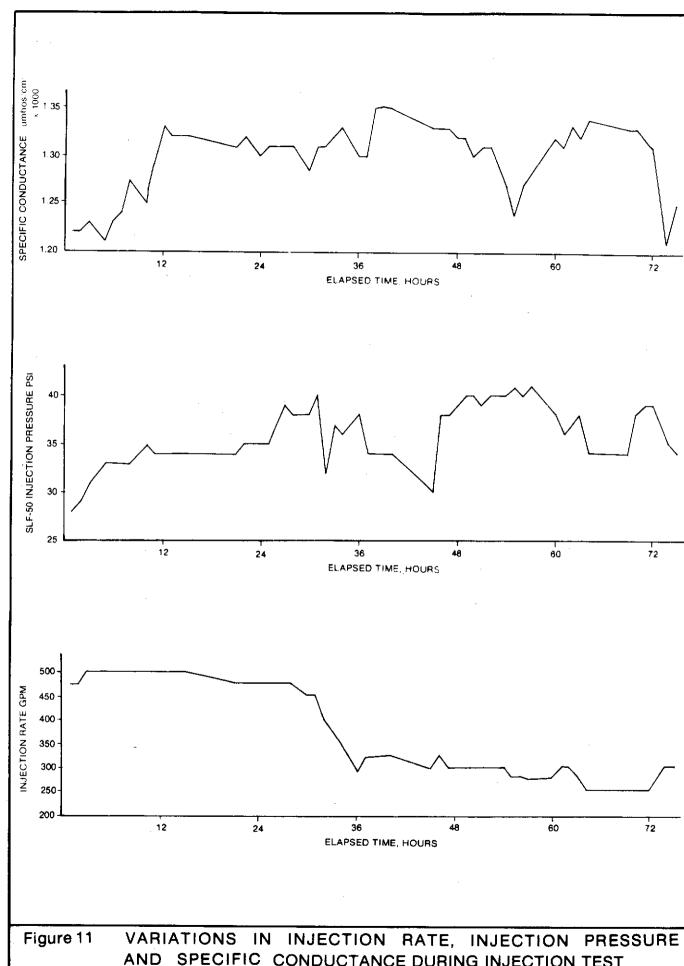
Potentiometric surface elevation at the exploratory well was measured on September 22 as 42.44 feet NGVD. Land elevation was measured as 31.13 feet NGVD, giving a head above land surface of 11.31 feet (4.90 psi wellhead pressure). The regional groundwater gradient estimated from potentiometric maps by Brown and Reece (1979) was 0.5 foot per mile in a northeast direction.

No estimates of parameters governing dispersive mixing were obtainable from the exploratory work. Estimates of longitudinal and transverse dispersivity from tracer tests and model studies for various materials are summarized by Mercer et al. (1982). They showed measured values of longitudinal and dispersivity in carbonate rocks as ranging from 1.0 meters for an intact chalk aguifer to 38.1 meters for fractured dolomite, with an average value for carbonate rocks of 13.7 meters. Segol and Pinder (1976) estimated longitudinal and transverse dispersivities for the limestones in southeast Florida as 6.7 and 0.7 meters, respectively. On this basis a preliminary estimate of longitudinal transmissivity for the injection horizon was assumed to be in the range of 6 to 15 meters. Transverse dispersivity was assumed to be in the order of 1 meter, based on the estimate by Segol and Pinder. Dispersivity, along with flow anisotropy affect the recoverability of injected water. In aguifers with fracture or solution porosity, such as is usually found in the limestone aguifers of the Floridan Aquifer System, mathematical formulations for recovery efficiency may inadequately represent the natural flow system resulting in significant differences between field results and theoretical estimates of recoverability (Merritt, 1982).

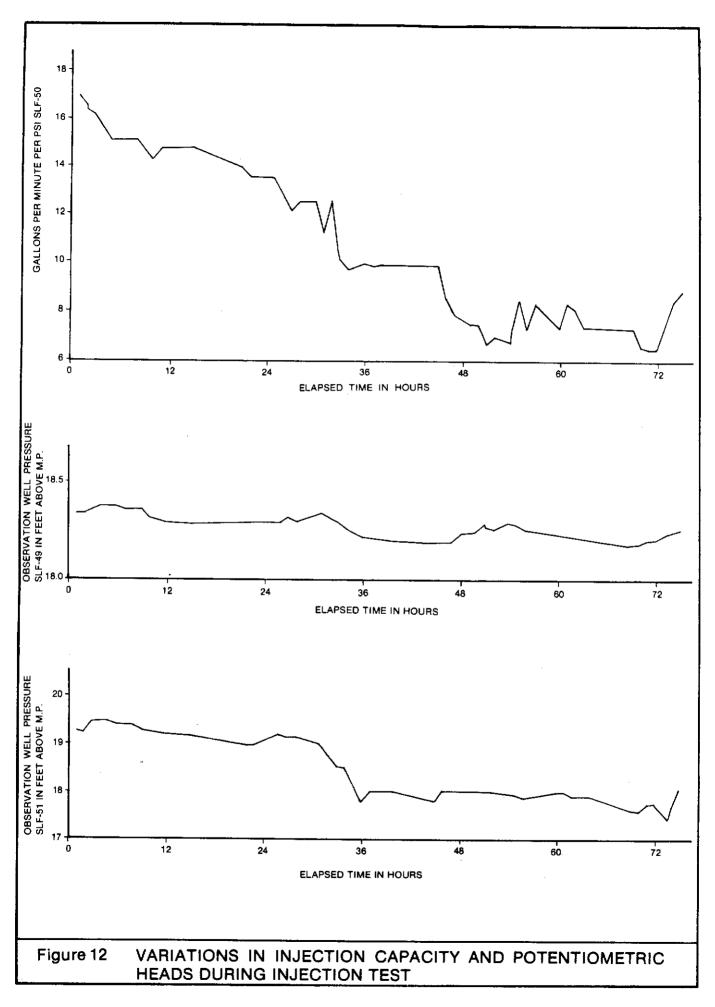
INJECTION/RECOVERY TEST

A short-term injection/recovery test was performed at the exploratory well (SLF-50) to evaluate problems associated with the technique and provide preliminary data for comparison with theoretical results. The test consisted of injecting approximately 1.5 million gallons of water into the exploratory well, and recovering this water by natural backflow after a residence period of 30 days. Data from this test are summarized in Appendix 6. Figures 11 and 12 illustrate data relevant to the injection phase. The data indicate a steady decline in injection capacity with time (Figure 12). It should be noted that the pumping rate was adjusted to maintain injection pressures below approximately 40 psi. The well was backflowed after 3392 minutes for approximately 30 minutes, and backflowed and surged with the pump at 4320 minutes for 1 hour. On both occasions there was some improvement in injection capacity of the well. Water extracted during backflowing and surging was murky with a reddish precipitate. Chemical analyses showed high total iron content in a sample taken from the backflow (39.42 mg/l total iron), indicating that iron precipitation may be a factor in plugging of the well.

Figure 12 also shows variations in potentiometric levels in monitor wells SLF-49 and SLF-51. Very little response was noted in well SLF-49 which is located 420 feet northwest of the injection well. Well SLF-51, which is located 148 feet northeast of the injection well, showed potentiometric head variations consistent with plugging of the injection well. In spite of increased injection pressures needed to force water into the injection horizon, water levels in monitor well SLF-51 showed a consistent decline. This was due to the reduction in the volume of water which entered the aquifer and, consequently, the lower pressures developed in the aquifer away from the injection well.



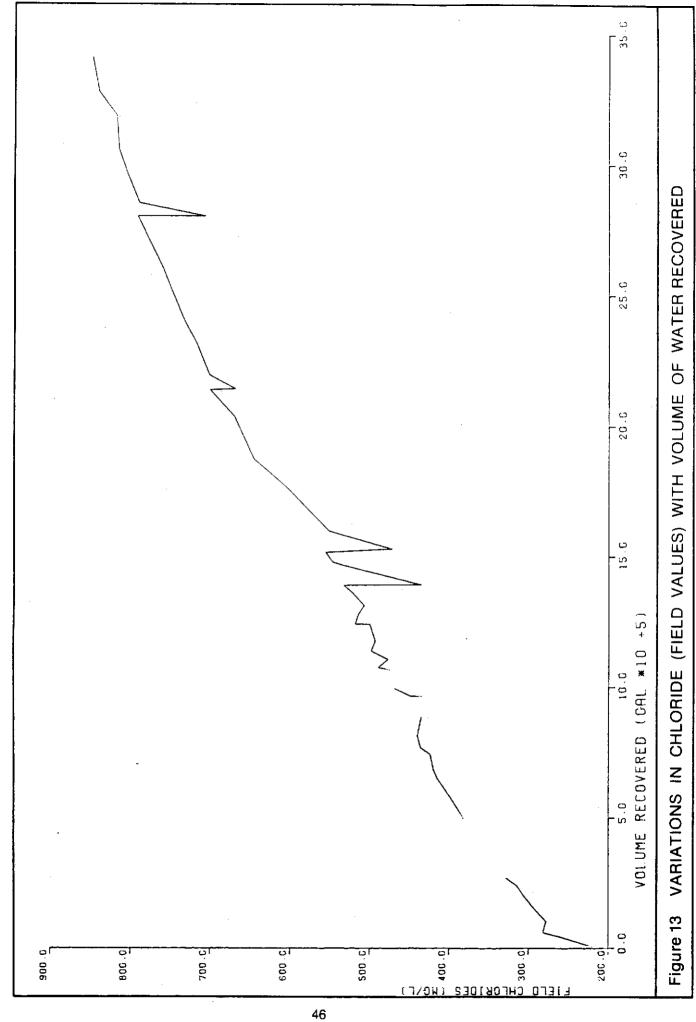
SPECIFIC CONDUCTANCE DURING INJECTION TEST

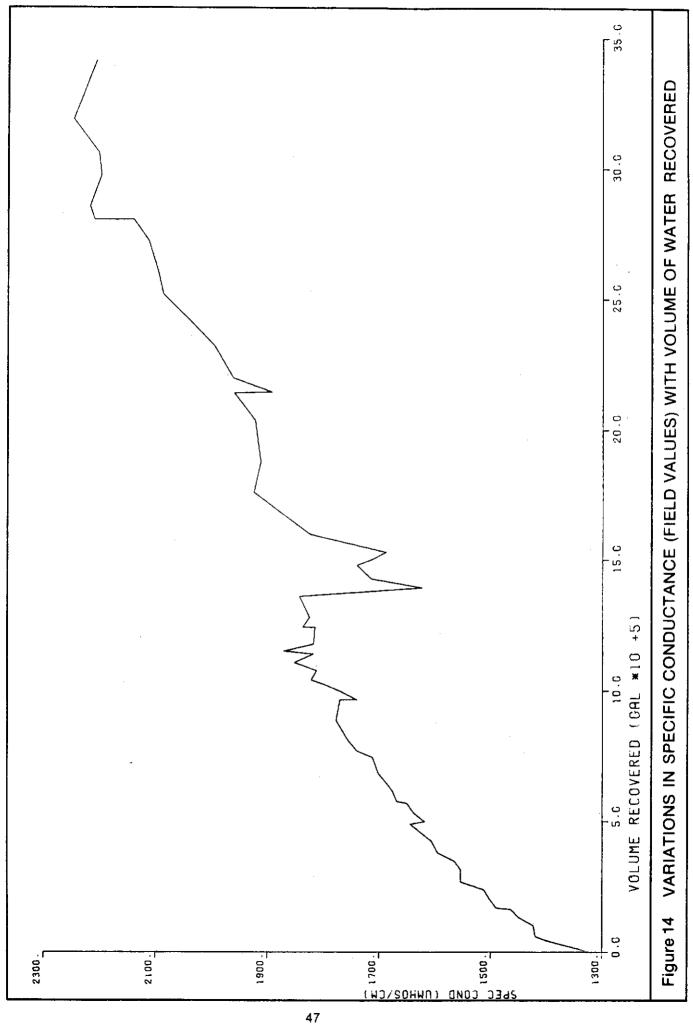


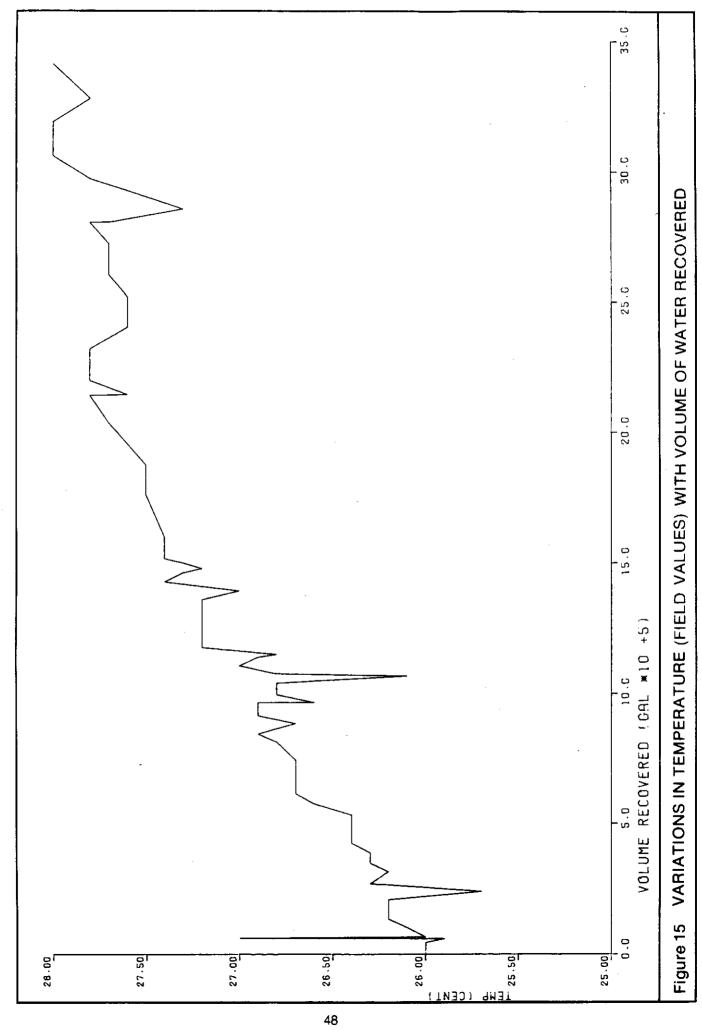
Injection was terminated after 4500 minutes. Recovery began 30 days after the end of injection. Recovery was carried out over a period of six weeks at rates varying from 140 to 250 gpm. Chloride concentrations and specific conductance and temperature were measured at frequent intervals in the field during recovery. Variations in these parameters as a function of volume of water recovered are shown on Figures 13, 14, and 15. Chloride concentrations showed good correlation with specific conductance values as shown on Figure 15. In general the recovery data showed a gradual increase in mineralization in the water as recovery progressed.

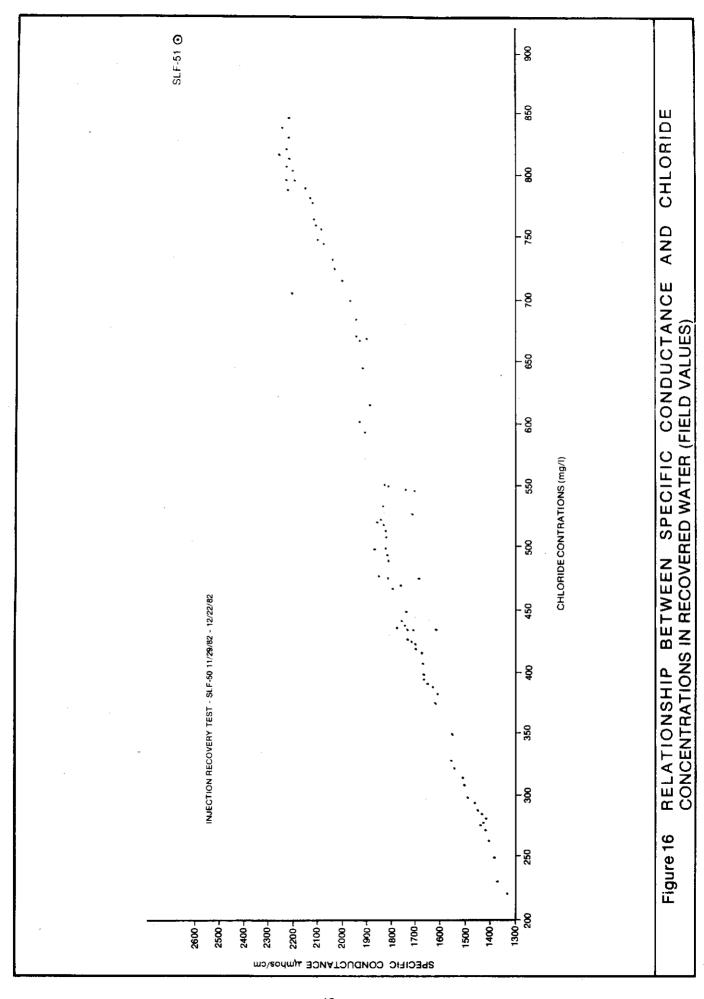
Figures 17 to 27 show plots of the variations in selected water quality parameters measured in the laboratory. The majority of chemical parameters show a gradual increase as recovery progressed, reflecting the higher percentage of more mineralized native water in the recovered water. Calcium and alkalinity concentrations, however, declined as recovery progressed. Sulphate concentrations remained relatively stable during the initial part of the test, but showed a sudden decrease after a break of 1 month in the recovery cycle when sulphate concentrations fell below the levels in both the injected water and native water. Assuming that this apparent reduction was not due to sampling or analytical error, a possible explanation could be conversion of sulphates to sulphides by bacterial activity and the release of hydrogen sulphide from the samples.

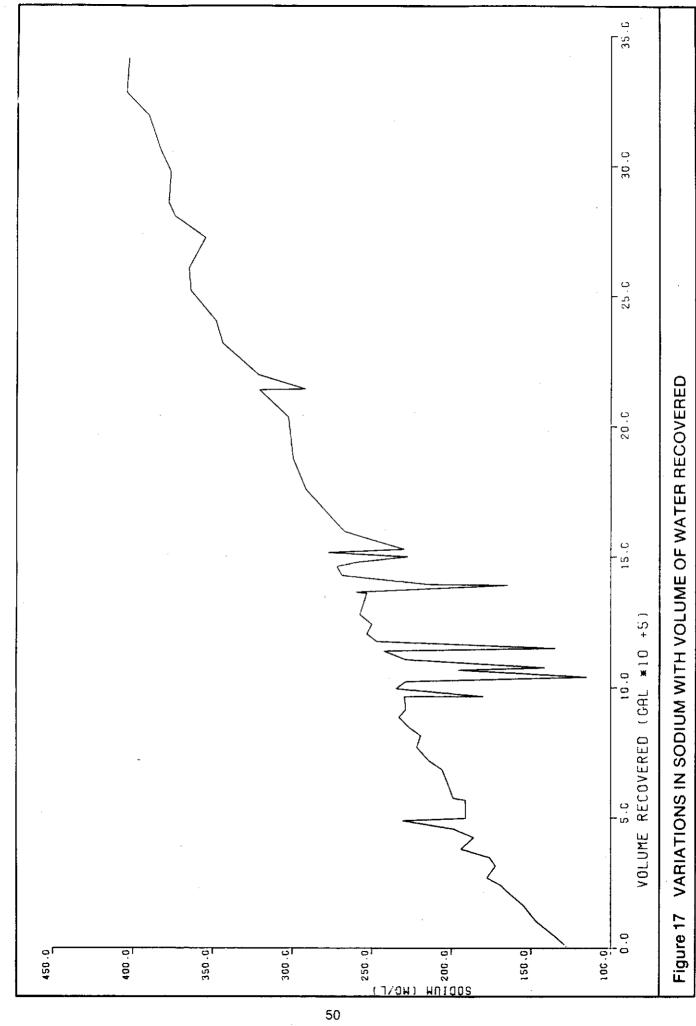
Total iron concentrations varied from 39.42 mg/l to less than 0.03 mg/l during recovery (see Appendix 3). The highest concentration was found during initial backflowing of the injection well at the end of the injection cycle. This sample was reddish in color and contained a reddish gelatinous precipitate. After approximately 14 minutes of backflow the reddish coloration and precipitate were no longer apparent. Other anomalously high values of total iron were found, mainly in samples taken after breaks in the

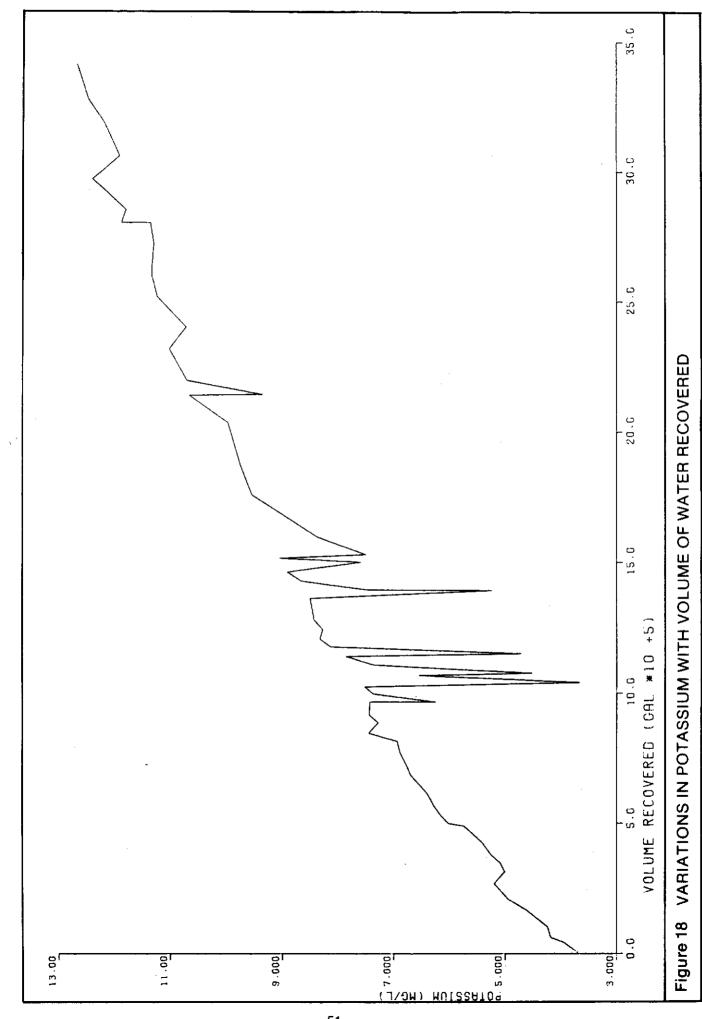


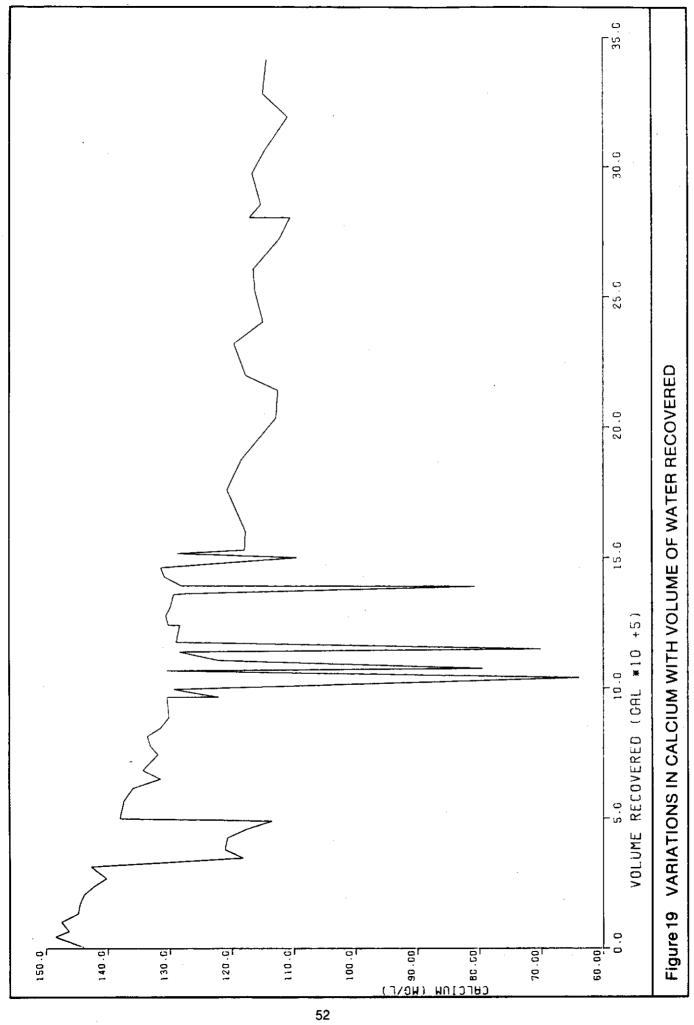


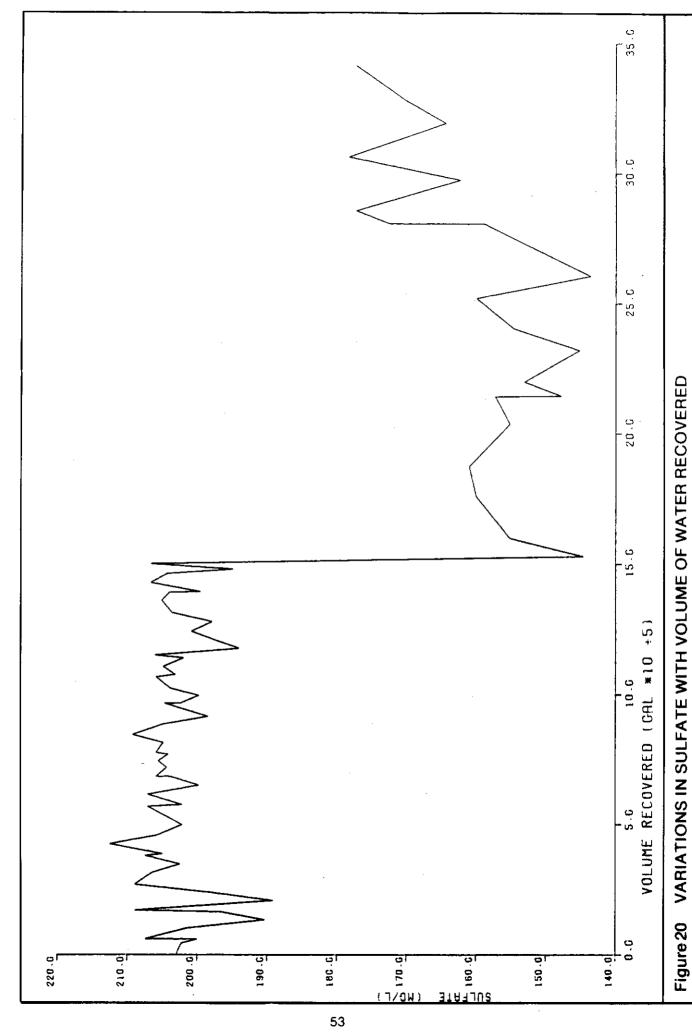


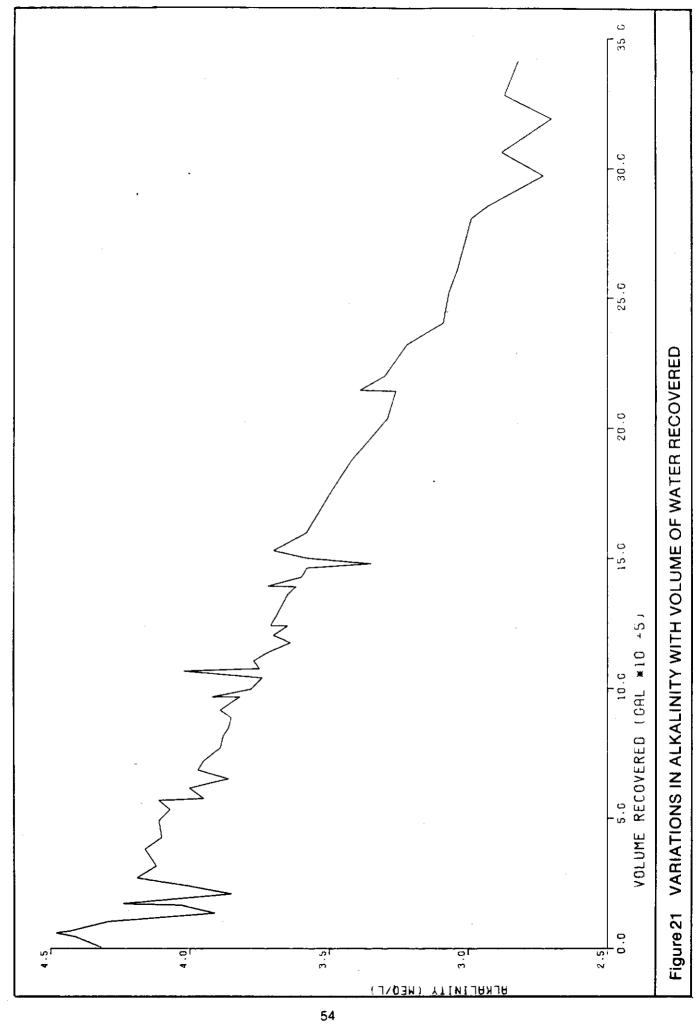


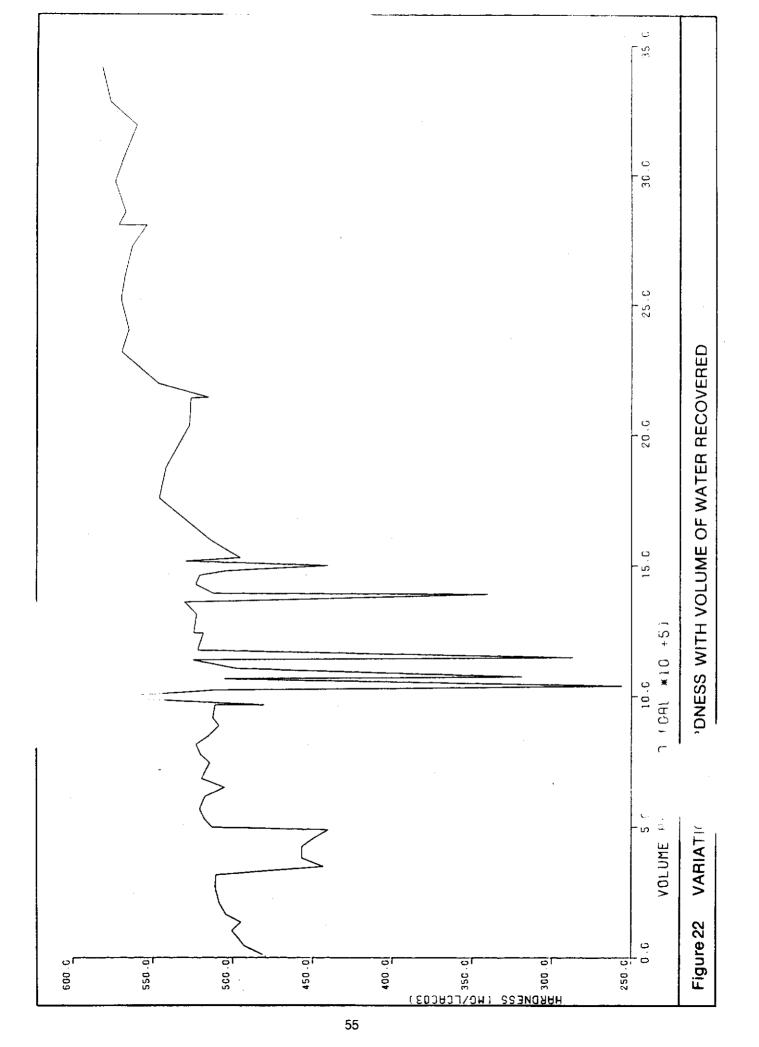


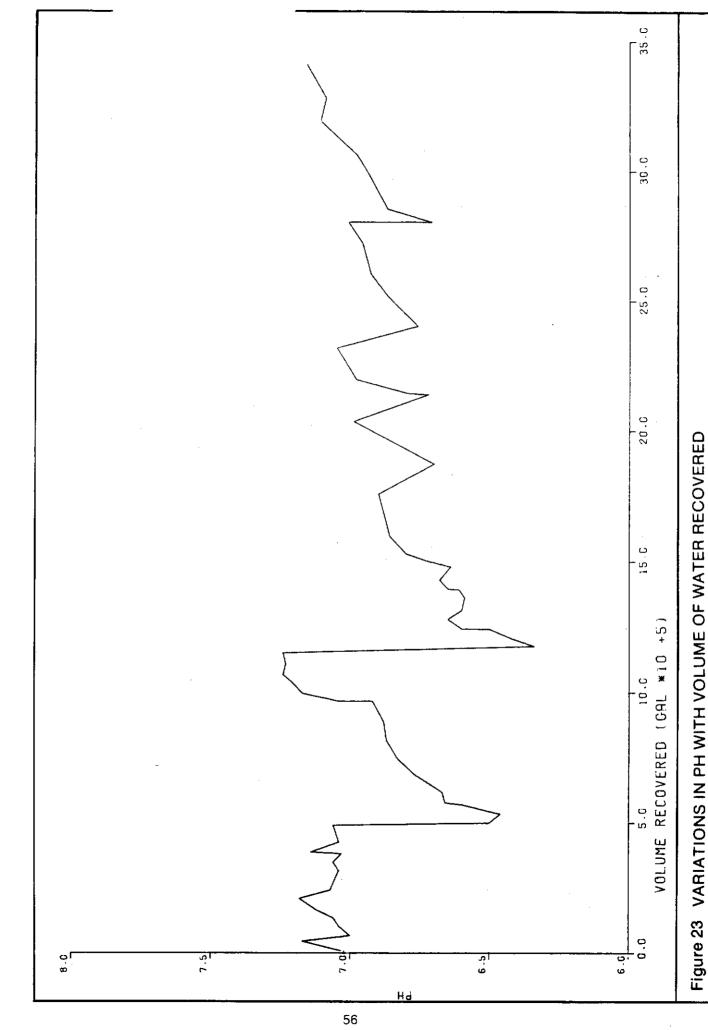


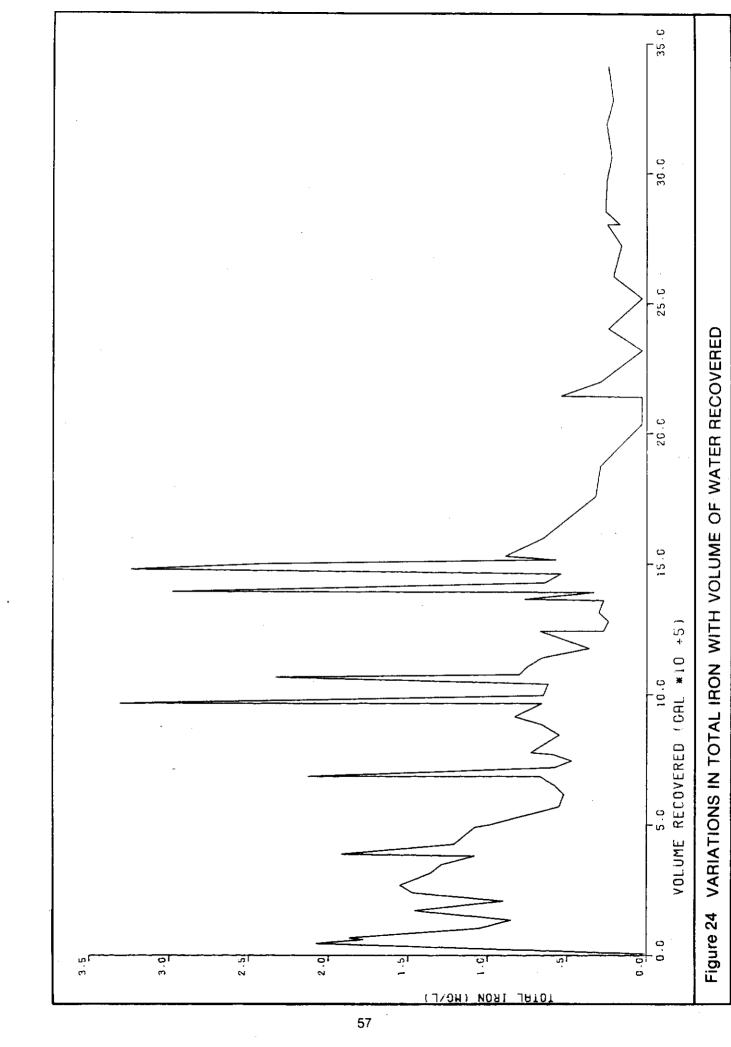


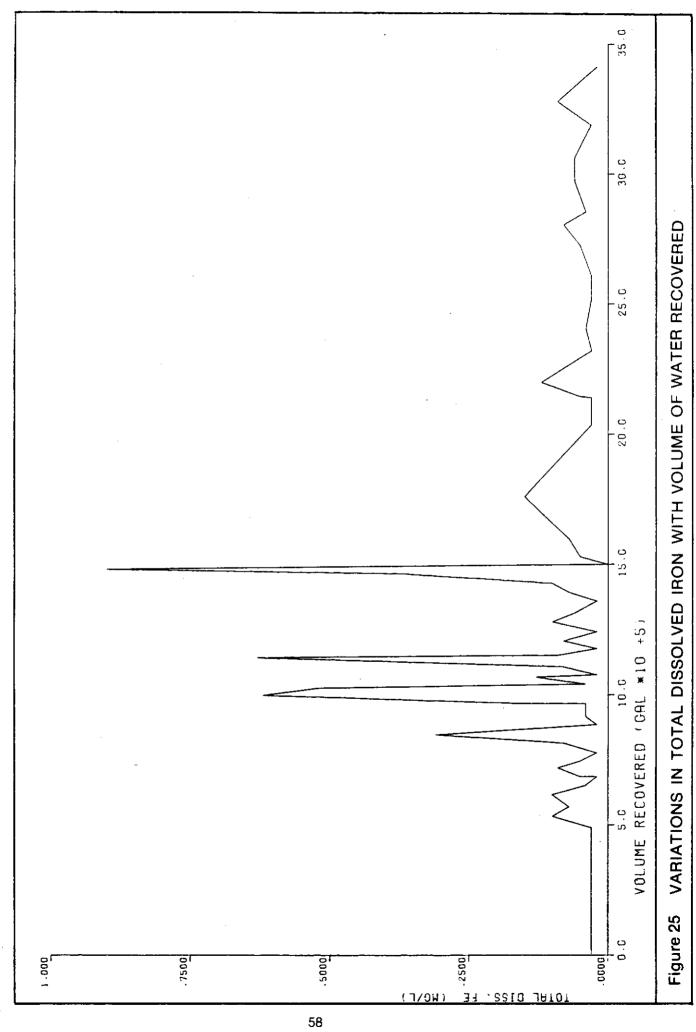


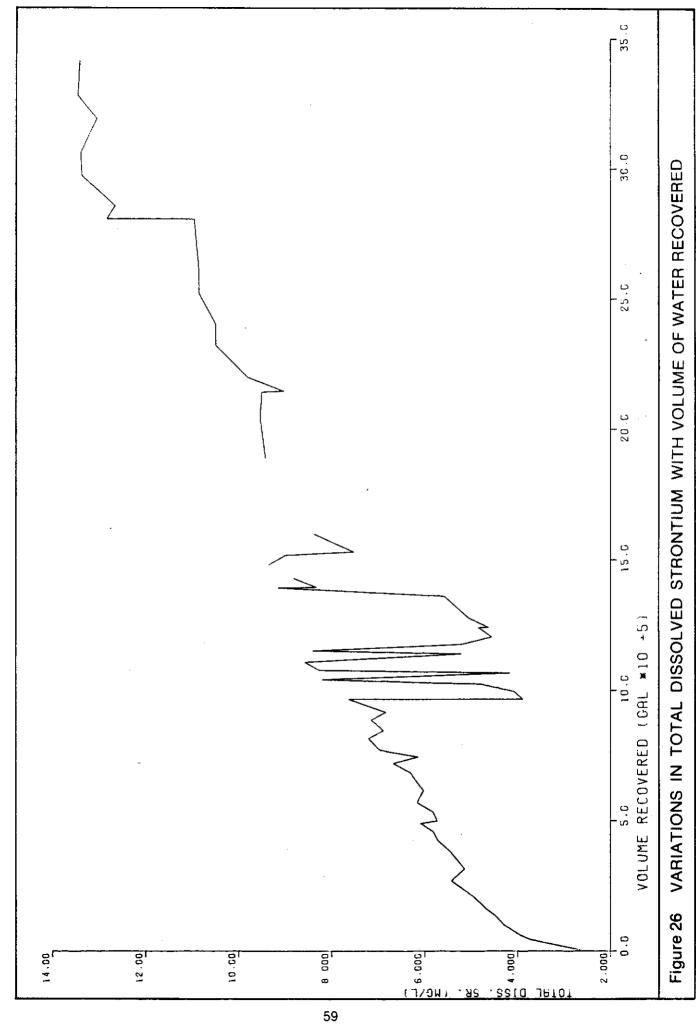


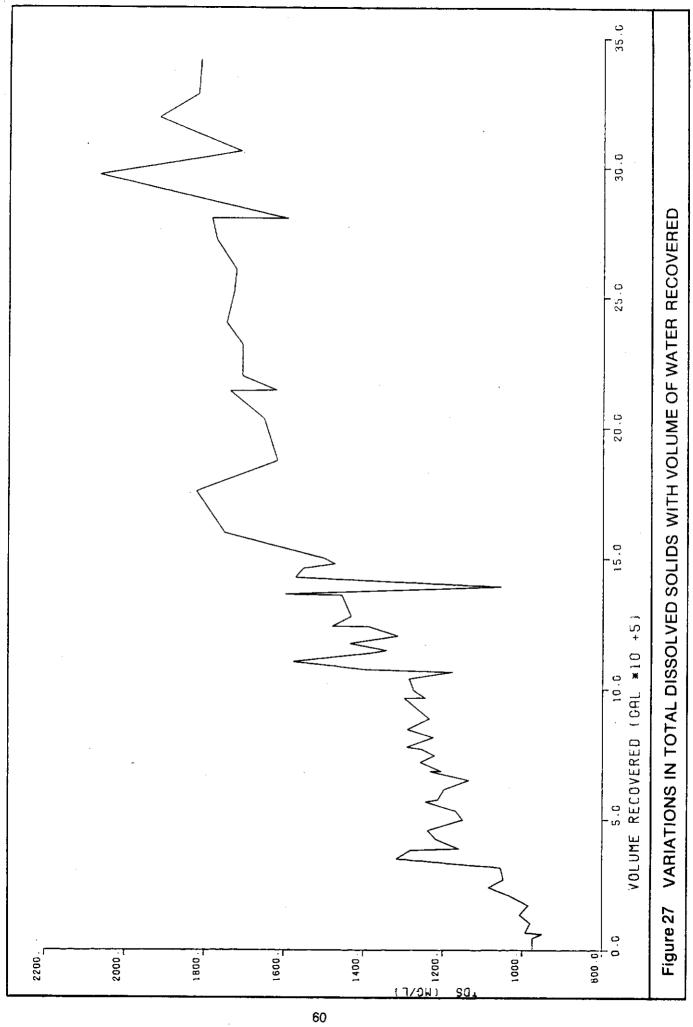










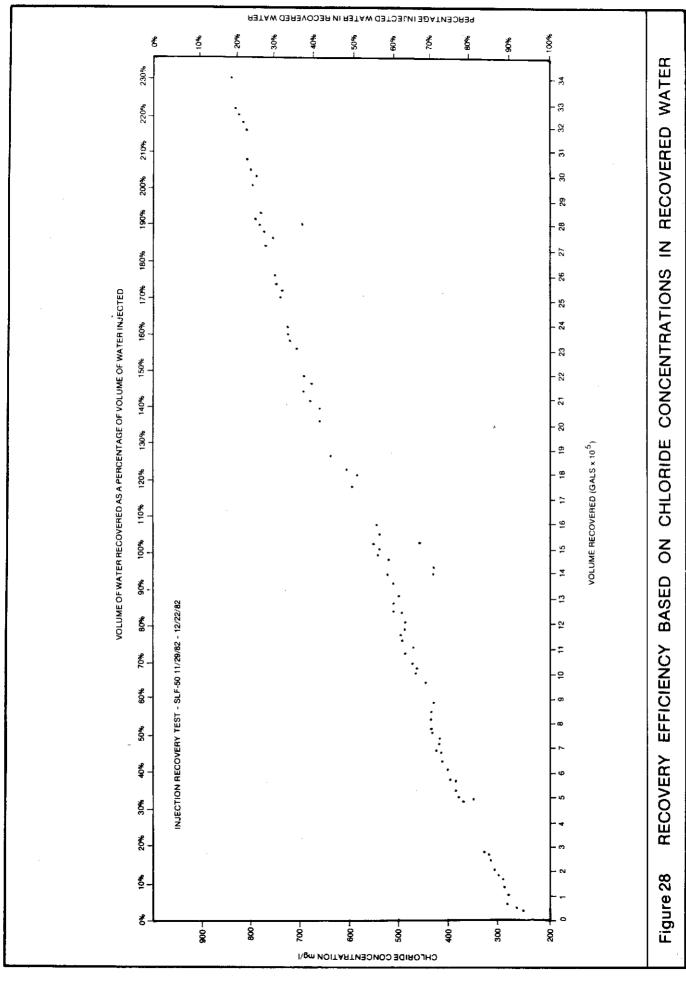


recovery process. In general, however, the data show a pronounced decline in total iron concentrations in the recovered water as recovery progressed.

Total suspended solids concentrations were measured during a part of the recovery test. An initial very high value of 280 mg/l was measured in the sample taken during backflow of the injection well at the end of the injection phase. During the first week of recovery suspended solids concentrations varied between 4 mg/l and 9 mg/l. During later periods of recovery the values varied between 1 mg/l and 2 mg/l.

The water quality data confirm the possibility of plugging of the injection well by suspended material produced as a result of chemical reactions between the injected and native water. Plugging could be minimized by pretreatment of the injected water or by frequent backflushing of the injection well during the injection cycle.

Figure 28 shows recovery efficiency expressed as a percentage of injected water, and the chloride concentration expressed as a concentration ratio of injected water in the native groundwater. The concentration ratios are based on initial concentrations of 200 mg/l and 1000 mg/l for the injected water and native groundwater, respectively. A gradual increase in chloride concentrations is noted, indicating diffusion and mixing of the injected water with the native groundwater. Because of the high chloride content of the injected water only 3 percent of the recovered water had chloride concentrations less than 250 mg/l. Figure 28 may be used to estimate recovery percentages for various values of maximum allowable chloride concentrations in the recovered water. For example, for a maximum chloride concentration of 500 mg/l, the percent recovery would be 45 percent. Figure 28 may also be used to estimate percent recovery assuming that the injected water had a different chloride concentration. For example, if the injected water had a chloride concentration of 50 mg/l, a chloride level of 250 mg/l would represent a 79



percent blend of injected water in native water and the recovery efficiency for a maximum chloride concentration in the recovered water of 250 mg/l would be 33 percent. This latter extrapolation assumes no changes in buoyancy effects or diffusivity as a result of the different ratios of ionic concentrations for the injected and native water.

COST EVALUATION

Costs associated with the injection/recovery method of water storage and retrieval in St. Lucie County were estimated by Khanal (1980) at approximately five cents per 1000 gallons before final treatment and transmission. The present phase of this investigation was not designed to verify these costs since injection and recovery have not been carried out for periods, or at rates equivalent to those used by Khanal in his analysis. However, preliminary indications, based on the experience gained in practical application of the method, suggest that actual raw water costs (before final treatment and distribution) may be much higher than estimated by Khanal. The items which may show significant variation in cost from those estimated by Khanal are:

- The number of wells needed for injection. Nine wells were estimated to inject 1,900 million gallons over a 150 day period and for recovery of 960 million gallons over a 120 day period (one annual cycle of injection/recovery). This assumes a continuous injection rate per well of 1000 gpm. However, due to hydrogeologic conditions, plugging, and the necessity for backflushing, the actual injected volume over a 150 day period could be substantially lower than estimated.
- The analysis by Khanal does not address pretreatment of the injected water which should, at a minimum, involve iron removal and filtration. Pretreatment costs are likely to add significantly to the costs for injection and recovery.

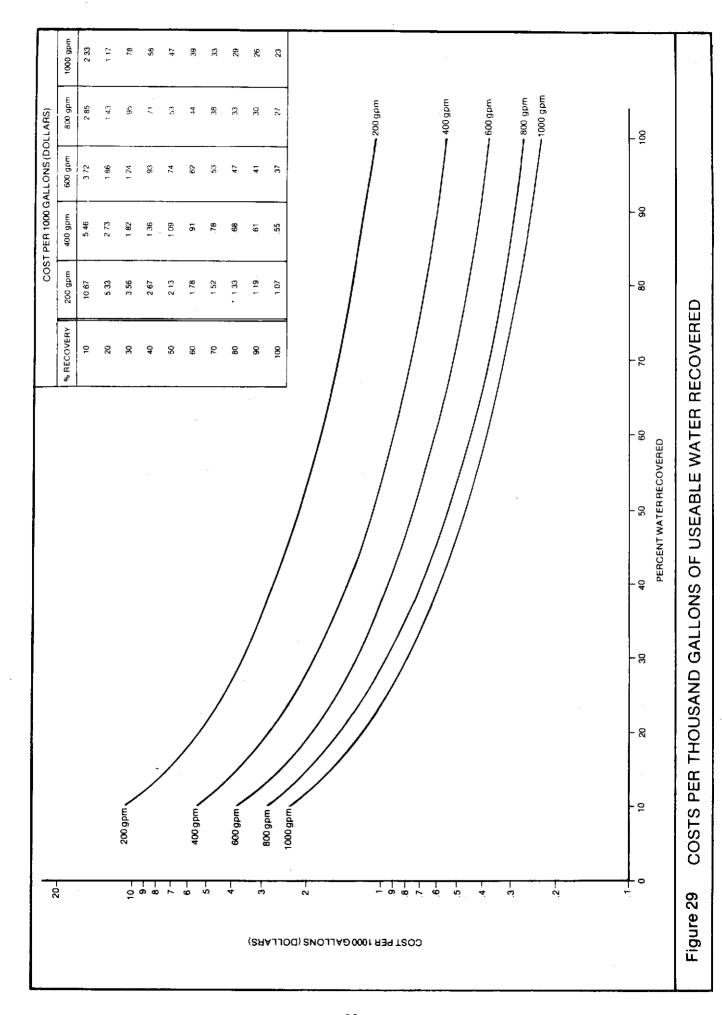
3) Operating costs associated with monitoring injection pressures, temporary halting of injection for backflushing, and resumption of injection are likely to be high. This could add significantly to the overall cost of the injection/recovery system.

Preliminary cost estimates for a single well, based on the aquifer and injection characteristics encountered at the exploratory site, are detailed on Table 5. Based on these estimates a more generalized assessment of costs is developed and displayed on Figure 29. The importance of the injection rate per well and the percentage of usable water which can be recovered in determining costs is well illustrated. In developing this figure only power costs and overheads were adjusted to take into account different injection rates. It was assumed that other costs would remain relatively constant. The period of 150 days of injection used in the calculations does not include time for rehabilitation or backflushing of the well.

For any specific application and hydrogeologic situation, the percent of usable water which can be recovered depends on the water quality which is considered suitable for that use. For example, for domestic use, the chloride limit is generally accepted to be 250 mg/l unless the water is to be blended with water from other sources. At this particular site, due to the relatively high chloride content of the water available for injection, the percent usable water recovered with a chloride concentration at or below 250 mg/l was approximately 3%. Assuming that the initial injection rate of approximately 400 gpm could be maintained, the costs would still be exhorbitant (more than \$10 per 1000 gallons). It should be born in mind, however, that at the site investigated, this injection rate could probably be maintained only if the injected water were pretreated or the injection well were frequently rehabilitated. These operations would add significantly to the costs of the process.

1. Capital Costs

(1 hydrogeologist for 1 month)	150,000 20,000	\$270,000	<pre>pital recovery) interest)\$ 31,714 efficiency 80% per 150 days = 33909 kwh X 4.5¢ per kwh) lush pump and switchgear costs) } 11,026 4,410</pre>	\$ 47,150 1 for 400 gpm injection rate for 150 days, 50% recovery)	000
a. Hydrogeologic Surveys (1 hydrogeologi b. Land Acquisition 1/2 acre at \$5000 pec. C. Well Construction Total Depth 775' Casing - 36" X 50' X 3/8" steel 20" X 130' X 3/8" steel 12" X 600' X 3/8" steel 12" X 600' - 775') Wellhead and Piping 50 X 12" pipe Flow regulator, pressuration	Other Fittings d. Hydrologic and System Testings e. Pump and Switchgear (25 hp 12" X 12" f. Engineering and Legal Fees (25% of a g. Contingency (20% of a, b, c, d, e, f	TOTAL CAPITAL COSTSAnnual Costs	 a. Debt service on \$270,000 + capital recovery Amortized over 20 years, 10% interest b. Operating Costs Power 400 gpm, TDH 100' efficiency 80% Rehabilitation and backflush Parts and Repair (10% of pump and swite Personnel c. Overhead (40% of b)	TOTAL ANNUAL COSTS	tater = 50 $\frac{$47,150}{43,200}$
		2.		ň	



For irrigation use, full recovery (100%) of injected water could be achieved. The recovered water, even in the initial cycle, would still be of better quality than the native groundwater and the quality would be expected to improve with each cycle. Based on the experience at this site and assuming that an injection rate of 400 gpm could be maintained, the cost per 1000 gallons would be approximately 55 cents. However, to maintain this injection rate, pretreatment of the injected water would be necessary. It is estimated that this would add approximately 20 to 25 cents to the cost per 1000 gallons. the total cost of 75 to 80 cents per 1000 gallons would probably not be competitive with other alternatives, such as conserving water by improving irrigation efficiency or utilizing the available resources from the shallow aquifers.

As shown on Figure 29 costs are more favorable where higher injection rates can be maintained; however, the higher transmissivities associated with the more productive wells would generally act to reduce the percentage of water of a given quality which could be recovered. If, as in the case of irrigation use, water quality is not a significant factor compared to storage of additional quantities of water in the aquifer, it could be possible to produce irrigation water at a cost of approximately 23 cents per 1000 gallons. This would, however, depend on locating sites where injection rates of 1000 gpm could be maintained and assumes that, at these sites, no pretreatment of the injected water would be required. Long-term injection/recovery tests under these conditions would be necessary to determine whether the technical problems involved in successfully operating the injection/recovery system (on a scale which would significantly impact the availability of irrigation water) could be solved.

It should be noted that costs for final treatment of recovered water (for domestic use) and distribution are not included in the figures previously

presented. Costs may also vary depending on variations in well design (for example, if deeper or larger diameter wells are required to achieve higher injection rates).

CONCLUSIONS

Based on the experience gained at this site it is concluded that:

- 1) Overall costs for treatment, injection and recovery of groundwater do not appear to be competitive with present costs for domestic or irrigation water.
- 2) Although suitable hydrogeologic conditions exist in the area, a successful injection/recovery system would require the availability of water with low chloride, iron, and total suspended solids concentrations for injection. The most cost-effective method of meeting these constraints would be to develop the program in conjunction with existing or proposed water treatment facilities.
- 3) Suitable hydrogeologic conditions exist in the area which would make large scale injection/recovery systems technically feasible. However, cost and operational complexity make this technique unfavorable at this time as a regional water resource management alternative.
- 4) The technique may have some applicability to specific non-potable uses, due to the high recovery efficiencies which can be achieved if the quality of the recovered water is not a critical factor.

REFERENCES

- Anderson, M. P., 1979. Using models to simulate the movement of contaminants through ground-water flow systems: Critical Reviews In Environmental Control, Vol. 9, p. 97-156.
- Applin, P. L. and E. R. Applin, 1944. Regional subsurface stratigraphy and structure of Florida and southern Georgia: Bull. of Am. Assoc. of Petroleum Geologist, Vol. 28, No. 12, Washington D.C., p 1673-1753.
- Armstrong, J. R., 1981. The Geology of the Floridan Aquifer System in Eastern Martin and St. Lucie Counties, Florida: Unpublished Masters Thesis, Florida State University.
- Bearden, H. W., 1972. Water Availability in Canals and Shallow Sediments in St. Lucie County, Florida: Florida Bureau of Geology Report of Investigations No. 62.
- Bouwer, H., 1978. Groundwater Hydrology: McGraw-Hill Series in Water Resources and Environmental Engineering.
- Brown, M. P., 1980. Aquifer Recovery Test Data and Analysis for the Floridan Aquifer System in the Upper East Coast Planning Area: South Florida Water Management District, Technical Publication #80-1.
- Brown, M. P. and D. E. Reece, 1979. Hydrogeologic reconnaissance of the Floridan Aquifer System, Upper East Coast Planning Area, South Florida Water Management District, Technical Map Series 79-1.
- Dall, W. H. and G. D. Harris, 1892. Correlation papers, Neocene: U. S. Geological Survey, Bulletin 84.
- Federico, A. C., 1983. Upper East Coast Water Quality Studies: South Florida Water Management District, Technical Publication 83-1.
- Hantush, M. S., 1964. Hydraulics of wells, IN: V. T. Chow (editor), Advances in Hydroscience Vol. I: 281-432: Academic Press. New York and London.
- Hantush, M. S. and C. E. Jacob, 1955. Non-steady radial flow in an infinite leaky aquifer: Am. Geophys. Union Trans., Vol. 36: 95-100.
- Jacob, C. E. and S. W. Lohman, 1952. Non-steady flow to a well of constant drawdown in an extensive aquifer: Am. Geophys. Union Trans., Vol. 33: 559-569.
- Kimbler, O. K., 1975. Cyclic Storage of Fresh Water in Saline Aquifers. Louisiana State University, Baton Rouge, LA, Bulletin 10.
- Khanal, N. N., 1980. Advanced Water Supply Alternatives for the Upper East Coast Planning Area: South Florida Water Management District, Technical Publication #80-6.
- Knapp, M. S., 1978. Environmental Geology Series Gainesville Sheet: Florida Bureau of Geology Map Series No. 79.

ŧ

- Mansfield, W. C., 1939. Notes on the upper Tertiary and Pleistocene Mollusks of Peninsular Florida: Florida Geological Survey, Bulletin 18.
- Mercer, J. W., S. D. Thomas and B. Ross, 1982. Parameters and Variables Appearing in Repository Siting Models, NUREG/CR3066: Prepared for the Nuclear Regulatory Commission.
- Merritt, M. L. Cyclic underground storage and recovery of freshwater in south Florida: A digital analysis of recoverability. U. S. Geological Survey Water-Supply Paper, in press.
- Merritt, M. L., F. W. Meyer, and W. H. Sonntag. Subsurface storage of freshwater in south Florida: U. S. Geological Survey Water Supply Paper, in press.
- Miller, J. A., 1982. Thickness of the Tertiary Limestone Aquifer System, southeastern United States: U. S. Geological Survey Open-File Report 81-1124.
- Mooney, R. T., 1980. The Stratigraphy of the Floridan Aquifer System East and Northeast of Lake Okeechobee, Florida: South Florida Water Management District, Technical Publication #80-9.
- Parker, G. G., G. E. Ferguson, S. K. Love, et al., 1955. Water resources of southeastern Florida, with special reference to the geology and groundwater of the Miami area: U. S. Geological Survey Water-Supply Paper 1255.
- Pitt, W. A., 1972. Sediment Loads in C-18, C-23, and C-24, Southeast Florida: U. S. Geological Survey Open-File Report 72013.
- Puri, H. S., 1953. Zonation of the Ocala Group in Peninsular Florida (Abstract): Jour. of Sed. Petrology, Vol. 23.
- Puri, H. S. and R. O. Vernon, 1964. Summary of the Geology of Florida and a Guidebook to the Classic Exposures: Florida Geological Survey, Special Publication #5 (revised).
- Reece, D. E., M. P. Brown, and S. D. Hynes, 1980. Hydrogeologic data collected from the Upper East Coast Planning Area: South Florida Water Management District, Technical Publication #80-5.
- Scott, T. M. and M. S. Knapp, 1983. The Hawthorn Formation in Peninsular Florida: Miami Geological Society Memoir No. 3, in preparation.
- Segol, G. and G. F. Pinder, 1976. Transient simulation of saltwater intrusion in southeastern Florida: Water Resources Research, 12, pp. 65-70.
- Sellards, E. H., 1912. The soils and other surface residual materials of Florida: Florida Geological Survey 4th Annual Report.
- Sniegocki, R. T., 1959. Plugging by air entrainment in artificial-recharge tests: Water Well Journal, V. 13, No. 6, p. 17-18, 43-44.

ì

- Sniegocki, R. T., 1963a. Geochemical aspects of artificial recharge in the Grand Prairie region, Arkansas: U. S. Geological Survey Water-Supply Paper 1615-E, 41 p.
- Sniegocki, R. T., 1963b. Problems in artificial recharge through wells in the Grand Prairie region, Arkansas: U. S. Geological Survey Water-Supply Paper 1615-F, 25 p.
- Sniegocki, R. T., F. H. Bayley III, and K. Engler, 1965. Testing Procedures and Results of Studies of Artificial Recharge in the Grand Prairie Region Arkansas: Geological Survey Water-Supply Paper 1615-G, 56 pp.
- Vernon, R. O., 1951. Geology of Citrus and Levy Counties, Florida: Florida Geological Survey, Bulletin 33.
- Wedderburn, L. A., M. S. Knapp, D. P. Waltz, and W. S. Burns, 1982.

 Hydrogeologic Reconnaissance of Lee County, Florida: South Florida Water

 Management District, Technical Publication #82-1, Parts 1, 2, and 3.

Lithologic Log, Well SLF-50

*-SEGG1

ST. LUCIE CO. TBES RB9E SEC 1400 27 20 17 N 80 29 53 W TUTAL DEPTH- 1000 FT. ELEV.- 25 FT. 100 SAMPLES- 0+ 1000 FT. COMPLETED- 02.02.20 DEPTH WURKED 1000 FT.

ulmer GeormySical LUGS AVAILABLE -

GAMMA NEUTRÜN CALIPHER ELECTRIC TEMPERATURE

HOLL NAMET

RECOVERY TEST WELL (SEF-50), STA.1.0.+11100005G, DRILLER ALVIN WOCSTER. REMARKS-

CUTTINGS DESCRIBED BY MIKE KNAPP (2-22-82), SAMPLE QUALITY (GCUE), REVERSE ALK DRILLING, WATER GUAL, POINT SAMPLE, PUMP TEST, AND OTHER GEOPHYSICAL DATA AVAILABLE.

HYURUGEULUGIC UNITS

U.U 130.0 SURFICIAL AGUIFER SYSTEM 130.0 COV.O HANTHURN CONFINING BEDS 500.0 1000.0 Fluridan Aguifer System

STRATIGRAPHIC FURMATIONS -

U.O- 30.0 UNLIFFERENTIATED SAND AND CLAY

30.0- 100.0 ANASTASIA FÉRMATION

100.0- 130.0 TAMIAMI FERMATION

130.0- 500.0 HAWTHERN FERMATIEN

COU. J- 740.0 CRYSTAL RIVER FURMATION

740.0- 760.0 WILLISTON FORMATION

760.0- 1000.0 AVEN PARK LIMESTENE

LITHCLEGIC LOG

w-secol . ST. EUGLE CO. Tabs, R39E, SEC 1400

- U.O- 10.0 SAND, GRAYISH GRANGE TO GRAYISH BRUAN, 35% PERCSITY, INTERGRANULAR, GRAIN SILE: MEDIUM, RANGE: VERY FINE TO COAKSE, SUB-ANGULAR, ANGULAR, MEDIUM SPHERICITY, UNCONSULIDATED,
- 10.6- 20.0 LIMESTONE, VERY LIGHT CRANGE TO WHITE, 12% PORCSITY,
 INTERGRANLIAR, GRAIN TYPE: CALCILUTITE, GRAIN SIZE: VERY
 FINE, RANGE: MICROCRYSTALLINE TO COARSE, POOR INDURATION,
 CALCILUTITE MATRIX, 20% GUARTZ SAND, MOLLUSKS,

FILL MATERIAL FROM ADJACENT CANAL

20.0- 30.0 SHELL BED, GRAYISH BROWN TO MODERATE GRAY, 25% POROSITY, INTERGRANULAR, POOR INDURATION, IRON CEMENT, CALCILUTITE MATRIX, 26% QUARTZ SAND, 02% CALCILUTITE, MOLLUSKS,

SOME FRAGS ARE WELL INDURATED SANDSTONE

- 30.0- 40.0 SANDŠTŪNĖ, GREENISH GRAY, 12% POROSITY, INTERGRANULAR, INTERCRYSTALLINĖ, PUSSIBLY HIGH PERMEABILITY, GRAIN SIZĖ: MEDIUM, RANGE: VERY FINE IU CŪAKSE, SUB-ANGULAR, RŪUNŪEŪ, MEDIUM SPPERICITY, GCUU INDURATIŪM, CALCILUTITĖ MATRIX, SPARKY CALCITE CEMENT, 10% CALCILUTITĖ, 10% SPAR, MCLLUSKS, FUSSIL MCECS.
- 40.0- 50.0 AS ABUVED
- 50.0- 60.0 INTERMIXED SANDSTONE AND SHELL
- 00.0- 70.0 AS ABOVE.
- 70.0- 80.0 SHELL BED, WHITE TO MODERATE LIGHT GRAY, 25% PORCSITY, INTERGRANULAR, POOR INJURATION, CALCILUTITE MATRIX, 05% CALCILUTITE, 25% QUARTZ SAND, MOLLUSKS, ECHINOID, CORAL,
- 80.0- 100.0 AS ABOVE.
- IGU.U- 110.0 SANDSTONE, WHITE TO LIGHT GREENISH GRAY, 15% PERESITY, INTERGRANULAR, GRAIN SIZE: FINE, RANGE: VERY FINE TO MEDIUM, SUB-ANGULAR, MEDIUM SPHERICITY, POOR INDURATION, CALCILUTITE MATRIX, 15% CALCILUTITE, MOLLUSKS,
- 110.0- 120.0 LIMESTONE, YELLUWISH GRAY, 15% PORUSITY, INTERGRANULAR, GRAIN TYPE: CALCILUTITE, BIOGENIC, SKELETAL, C5% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: MICROCRYSTALLINE, KANGE: MICROCRYSTALLINE TO COARSE, MODERATE INDURATION, CALCILUTITE MATRIX, DOLOMITE COMENT, 10% DOLOMITE, 02% PHOSPHATIC SAND, 25% QUARTZ SAND, MULLUSKS, CORAL,
- 120.0- 130.0 AS ABOVE.
- 130.0- 146.0 SAND, LIGHT OLIVE GRAY TO GRAYISH OLIVE, 12% POROSITY,
 INTERGRANULAR, LOW PERMEABILITY, GRAIN SIZE: FINE, RANGE:
 VERY FINE TO MEDIUM, SUB-ANGULAR, ANGULAR, MEDIUM
 SPHERICITY, POUR INDURATION, DOLOMITE CEMENT, CLAY MATRIX,
 TOX DOLOMITE, 65% CLAY, 07% PHOSPHATIC SAND, CALCAREOUS,
 DOLOMITIC, MOLLUSKS,
- 140.0- 170.0 AS ABOVE,
- 170.0- 180.0 SAND, GRAYISH OLIVE, 12% POROSITY, INTERGRANULAR, LOW PERMEABILITY, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MEDIUM, SUB-ANGULAR, ANGULAR, MEDIUM SPHERICITY, POUR INDURATION, OCLOMITE CEMENT, CLAY MATRIX, 10% DULOMITE, 05% CLAY, 65% PHOSPHATIC SAND, CALCAREDUS, OCLOMITIC, MBLLUSKS,
- 180.0- 190.0 AS ABOVE,
- 190.0- 200.0 DULG-SILT, LIGHT DLIVE GRAY, 10% PORDSITY, INTERGRANULAR, LGM PERMEABILITY, POUR INDURATION, DELEMITE CEMENT, LALCILUTITE MATRIX, CLAY MATRIX, 10% CALCILUTITE, 05% CLAY, 30% QUARTY SAND, 03% PHOSPHATIC SAND, MOLLUSKS, DIATOMS, BENTHONIC FORAMINIFERA,
- 200.0- 240.0 AS ABUVE.

- 240.0- 250.0 SAND, GRAYISH OLIVE, 10% PORUSITY, INTERGRANCER, LOW PERMEABILITY, GRAIN SIZE: VERY HINE, RANGE: VERY FINE TO MEDIUM, SUB-ANGULAR, ANGULAR, MEDIUM SPHERICITY, PEOR INCURATION, CALCILUTITE MATRIX, DOLUMITE CEMENT, CLAY MATRIX, 10% DOLUMITE, CD% CAECILUTITE, CD% CEAY, D2% PHOSPHATIC SAND, DIATOMS, BENTHONIC FORAMINIFERA, PLANKTONIC FORAMINIFERA,
- 250.0- 260.0 AS ABOVE,
- 260.0- 270.0 DOLU-SILF, GRAYISH DLIVE, 10% PÜRÜSITY, INTERGRANDEAR, LÜW PERMEABILITY, POUR INDURATION, DOLUMITE CEMENT, CALCILUTITE MATRIX, CLAY MATRIX, 10% CALCILUTITE, 05% CLAY, 35% SILT, 01% PHOSPHATIC SAND, BENTHUNIC FORAMINIFERA,
- 2/0.0- 290.0 AS ABOVE,
- 290.0- 300.0 ÜÜLC-SILT, GRAYISH ÖLIVE, 10% PÜRUSITY, INTERGRANULAR, LÜN PERMEABILITY, PÜÜR INDURATION, ÜÜLÜMITE CEMENT, CALCILLTITE MARRIX, CLAY MATRIX, 10% CALCILUTITÉ, C5% CLAY, 35% SILT, C1% PHUSPHATIC SAND, DIATCHS, BENTHÜNIC FÜRAMINIFERA, PLANKTÜNIC FUKAMINIFERA,
- 300.0- 310.0 AS ABBVE,
- 310.0- 320.0 SAND, GRAYISH ÜLIVE, 124 PÜRÜSITY, INTERGRANULAR, LUW PERMEADILITY, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TÜ CÜARSE, SUB-ANGULAR, ANGULAR, MEDIUM SPHERICITY, PBOR INDURATIUN, DÜLÜMITE CEMENT, CALCILUTITE MATRIX, CLAY MATRIX, 25% ÜELÜMITE, C5% CALCILUTITE, 02% CLAY, BENTHONIC FÜRAMINIFERA,
- 320.0- 330.0 AS ABOVE.
- 330.0- 340.0 SAND, LIGHT GLIVE, 14% PORGSITY, INTERGRANULAR, LOW PERMEABILITY, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO CHARSE, SUB-ANGULAR, MEDIUM SPHERICITY, POER INDURATION, DELCMITE CEMENT, CALCILUTITE MATRIX, CLAY MATRIX, 25% DOLOMITE, 15% CALCILUTITE, 01% CLAY, 03% PHOSPHATIC SAND, MULLUSKS, BENTHONIC FORAMINIFERA,
- 340.0- 350.0 AS ABOVE,
- 350.0- 360.0 SAND, LIGHT CLIVE, 15% PURCSITY, INTERGRANULAR, LOW PERMEABILITY, GRAIN SIZE: MEDIUM, RANGE: VERY FINE TO CUARSE, SUB-ANGULAR, ROUNDED, MEDIUM SPHERICITY, PCOR INDURATION, DOLOMITE CEMENT, CLAY MATRIX, 15% DOLOMITE, 02% CLAY, 03% PHOSPHATIC SAND, MOLLUSKS, BENTHONIC FORAMINIFERA,
- 360.0- 370.0 DULO-SILT, LIGHT DLIVE, 12% POROSITY, INTERGRANULAR, LCW PERMEABILITY, POOR INDURATION, DOLUMITE CEMENT, CALCILUTITE MATRIX, CLAY MATRIX, 15% CALCILUTITE, C1% CLAY, 02% PHOSPHATIC GRAVEL, 03% PHOSPHATIC SAND, MOLLUSKS, CORAL, BENTHONIC FORAMINIFERA,
- 370.0- 380.0 AS ABDVE

- 360.0- 340.0 BÜLÜMITE, LIGHT GREENISH YELLÜM TU LIGHT BLIVE, 13% PURÜSİTY, İNTERGRANULAR, DC-90% ALTERED, EUHEDRAL, GRAIN SİZE: VERY FINE, RANGE: VERY FINE TÜ MICRECRYSTALLINE, MÜLEKATE INÜLKATIĞN, BULÜMITE CEMENT, CALCILUTITE MATRIX, ZUR CALCILUTITE, 64% PHUSPHATIC SANU, 62% PHESPHATIC GRAVEL, 10% GUARTZ SANU, MÜLLÜSKS,
- 340.0- 400.0 AS ABEVE,
- 400.0- 410.0 BULU-SILT, LIGHT BLIVE, 122 PORBSITY, INTERGRANULAR, LOW PERMEABILITY, POUR INDURATION, DOLOMITE CEMENT, CALCILUTITE MATRIX, CLAY MATRIX, 16% CALCILUTITE, C5% CLAY, 15% QUARTY SAND, D6% PHÚSPHATIC SAND, MULLUSKS,
- 410.0- 440.0 AS ABOVE,
- 440.0- 450.0 Limestune, white, 12% purusity, intergranular, grain type: calcilutite, crystals, 01% alluchemical constituents, grain size: miukucrystalline, range: microcrystalline to very fine, muuekate induratiun, calcilutite matrix, dulumite cement, 15% dulumite, 03% phusphatic sand, 02% guaktz sand,

SAMPLE IS MIXTURE OF LVS AND DOLG-SILT

- 450.J- 460.0 AS ABEVE.
- 460.0- 470.0 DULUMITE, LIGHT ULIVE TO VERY LIGHT GRAY, 12% PORUSITY, INTERGRANULAR, LUW PERMEABILITY, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICRUCRYSTALLINE, MUUERATE INDURATION, DOLOMITE CEMENT, CALCILUTITE MATRIX, 37% CALCILUTITE, OB% PHOSPHATIC SAND, 05% WOARTZ SAND, ECHINOTO, MULLUSKS,
- 470.0- 480.0 AS ABUVE.
- 486.0- 490.0 BÜLÜMITE, VERY LIGHT ÜRANGE, 122 PÜRÜSITY, INTERGRANULAR, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICROCRYSTALLINE, MÜDERATE INDURATION, DOLOMITE CEMENT, CALCILUTITE MATRIX, 35% LALCILUTITE, C3% PHOSPHATIC SAND, MÜLLUSKS,
- 490.6- 510.0 AS ABOVE,
- 510.0- 520.0 SAMPLE IS MIXTURE OF DOLUMITE AND DOLC-SILT
- 520.0- 530.0 AS ABOVE,
- 530.0- 540.0 DOLO-SILT, LIGHT OLIVE, 12% POROSITY, INTERGRANULAR, LOW PERMEABILITY, POOR INDURATION, DOLOMITE CEMENT, CALCILUTITE MATKIX, CLAY MATRIX, 15% CALCILUTITE, 15% CLAY, 06% PHOSPHATIC SAND, 15% QUARTZ SAND, MOLLUSKS,
- 540.0- 550.0 DULUMITE, VERY LIGHT DRANGE, 122 POROSITY, INTERGRANULAR, 50-902 ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICROCRYSTALLINE, MODERATE INDURATION, DOLGMITE CEMENT, CALCILUTITE MATRIX, 25% CALCILUTITE, 04% PHOSPHATIC SAND, 02% QUARTZ SAND, MOLLUSKS,
- 550.0- 570.0 AS ABOVE.

- 570.0- 580.6 BBLUMITE, VERY LIGHT BRANGE, 12% PURUSITY, INTERGRANULAR, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICROCRYSTALLINE, MODERATE INDURATION, DOLOMITE CEMENT, CALCILUTITE MATRIX, 20% CALCILUTITE, 63% PHOSPHATIC SAND, U2% LUARTA SAND, MULLUSKS,
- DBG.G- BYC.O SAMPLE IS MIXT. OF DGCOMITERLYSEAND GREEN CLAY
- 390.0- 600.0 AS ABLVES
- OUC.O- 610.C LIMESTONE, VERY LIGHT CRANGE, 14% PORCSITY, INTERGRANULAR, MULCIC, GRAIN TYPE: CALCILUTITE, SKELETAL, BIOGENIC, 30% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: MICROCRYSTALLINE, KANGE: MICROCRYSTALLINE TO COARSE, GOOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, 15% BOLOMITE, 10% PHOSPHATIC SAND, MULLUSKS, BRYUZGA, BENTHONIC FORAMINIFERA, ECHINGIO,
- 610.0- 620.0 AS ABOVE WITH CHARSE PHUS. (15%)
- 620.0- 635.0 AS ABOVE.
- 635.0- 650.0 V.C. PHOS. AND BIDGENIC L/S IN SAMPLE.
- 650.0- 660.0 LIMESTONE, VERY LIGHT GRANGE, 16% PURESITY, INTERGRANGER, GRAIN TYPE: BIOGENIC, CRYSTALS, CALCILUTITE, 5C% ALLCCHEMICAL CONSTITUENTS, GRAIN SIZE: VERY FINE, RANGE: MICKECKYSTALLINE TO CLARSE, GOOD INDURATION, DOLOMITE COMENT, CALCILUTITE MATRIX, SPARRY CALCITE COMENT, 15% DOLOMITE, BENTHONIC FORAMINTERA, MULLUSKS, ECHINGID, FOSSIL MOLDS,
- 660.0- 670.0 LIMESTONE, VERY LIGHT GRANGE 10 WHITE, 16% PCRCSITY,
 INTERGRANLLAR, GRAIN TYPE: SKELETAL, CALCILUTITE, CRYSTALS,
 65% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: MEDIUM, RANGE:
 MICKECRYSTALLINE TO COARSE, GOOD INDURATION, CALCILUTITE
 MATRIX, BENTHENIE FORAMINIFERA, MELEUSKS, ECHINOID, BRYEZEA,
 - LEPIDUCYCLINA SP., PHOSPHATIC L/S, AND HIGHLY RXTAL DOLO.
- 670.0- 680.0 LIMESTONE, VERY LIGHT GRANGE TO WHITE, 18% PCRCSITY,
 INTERGRANULAK, PGSSIBLY HIGH PERMEABILITY, GRAIN TYPE:
 BICGENIC, CALCILUTITE, SKELETAL, 75% ALLOCHEMICAL
 CONSTITUENTS, GRAIN SIZE: MEDIUM, KANGE: MICROCRYSTALLINE TO
 CUARSE, GCOD INDURATION, CALCILUTITE MATRIX, COQUINA,
 BENTHONIC FORMMINIFERA, MCLLUSKS, ECHINOID, BRYOZOA,
- 680.0- 700.0 AS ABOVE,
- 700.0- 712.0 COGLINA OF LEPS, GOOD PORUSITY.
- 712.0- 720.0 AS ABOVE,
- 720.0- 730.0 LIMESTONE, VERY LIGHT GRANGE TO WHITE, 18% PERCSITY,
 INTERGRANULAR, POSSIBLY HIGH PERMEABILITY, GRAIN TYPE:
 BICGENIC, CALCILUTITE, SKELETAL, 75% ALLOCHEMICAL
 CONSTITUENTS, GRAIN SIZE: MEDIUM, RANGE: MICROCRYSTALLINE TO
 COARSE, GOOD INDURATION, CALCILUTITE MATRIX, COQUINA,
 CHALKY, BENTHONIC FORAMINIFERA, MOLLUSKS, ECHINOID, BRYCZCA,
 CORAL,

740.0- 750.0 LIMESTÜNE, VERY LIGHT GRANGE, 14% PÜRÜSITY, INTERGRANULAR, GRAIN TYPE: BIUGENIC, SKELETAL, URYSTALS, 60% ALLÜCHEMICAL CENSTITUENIS, GRAIN SIZE: FINE, KANGE: MICKECRYSTALLINE TO MEDIUM, GÜĞU INDURATIĞN, CALCILUTITE MATRIX, SPARKY CALCITE CEMENT, BENTHONIC FÜKAMINIFEKA, MÜLLUSKS, ECHINCIG, BKYBZGA, CÜKAL,

MUSTLY SMALLER FURAMS (AMPHISTEGINA SP.()

- 750.0- 760.0 AS ABOVE.
- 760.0- 767.0 BÜLEMITE, MÜCERATE YELLÜWISH BKÜWN, 10% PÜRESITY, INTERCRYSTALLINE, MÜLDIC, INTERGRANDEAR, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, KANGE: VERY FINE TÜ MICRBCRYSTALLINE, GÜÜD INCLRATION, DÜLEMITE CEMENT, CALCILUTITE MATRIX, 10% CALCILUTITE, FLSSIL MÜLDS,
- 767.0- 775.0 AS ABGVE.
- 775.C- 700.C LIMESTONE, WHITE, 12% POROSITY, INTERGRANULAR, GRAIN TYPE: CALCILUTITE, BIOGENIC, CRYSTALS, 30% ALLOCHEMICAL CUNSTITUENTS, GRAIN SIZE: MICKECKYSTALLINE, RANGE: MICROCRYSTALLINE TO MEDIUM, GOOD INDURATION, CALCILUTITE MATRIX, BENTHONIC FORAMINIFERA, FOSSIL MULUS, CONES,

BICTYECHNUS CECKET

- 760.0- 790.0 LIMESTONE, VERY LIGHT DRANGE, 14% POROSITY, INTERGRANULAR, GRAIN TYPE: CALCILUTITE, BIDGENIC, SKELETAL, 55% ALLOCHEMICAL CUNSTITUENIS, GRAIN SIZE: FINE, RANGE: MICROCRYSTALLINE TO MEDIUM, GUOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, 20% DOLOMITE, BENTHONIC FORAMINIFERA, ECHINOID,
- 790.0- 800.0 SOME FRAGS. HAVE GOOD PORESITY
- 830.0- 810.0 DULUMITE, GRAYISH GRANGE TO MODERATE YELLOWISH BROWN, 10% PORUSITY, INTERGRANULAR, INTERCRYSTALLINE, MOLDIC, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICRUCRYSTALLINE, GOOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, 10% CALCILUTITE, FOSSIL MOLDS,
- 810.0- 820.0 CALCARENITE, VERY LIGHT ORANGE, 15% PORGSITY, INTERGRANLLAR, GOOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, 10% DULOMITE, BENTHONIC FORAMINIFERA, MOLLUSKS, ECHINOID,
- 820.0- 830.0 AS ABEVE.
- 830.0- 840.0 LIMESTONE, VERY LIGHT URANGE, 13% POROSITY, INTERGRANULAR, GRAIN TYPE: BIOGENIC, CALCILUTITE, CRYSTALS, 60% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: FINE, RANGE: MICROCRYSTALLINE TO COARSE, GOOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, SPARRY CALCITE CEMENT, 15% DOLOMITE, BENTHONIC FORAMINIFERA, MOLLUSKS, ECHINOID, CONES,

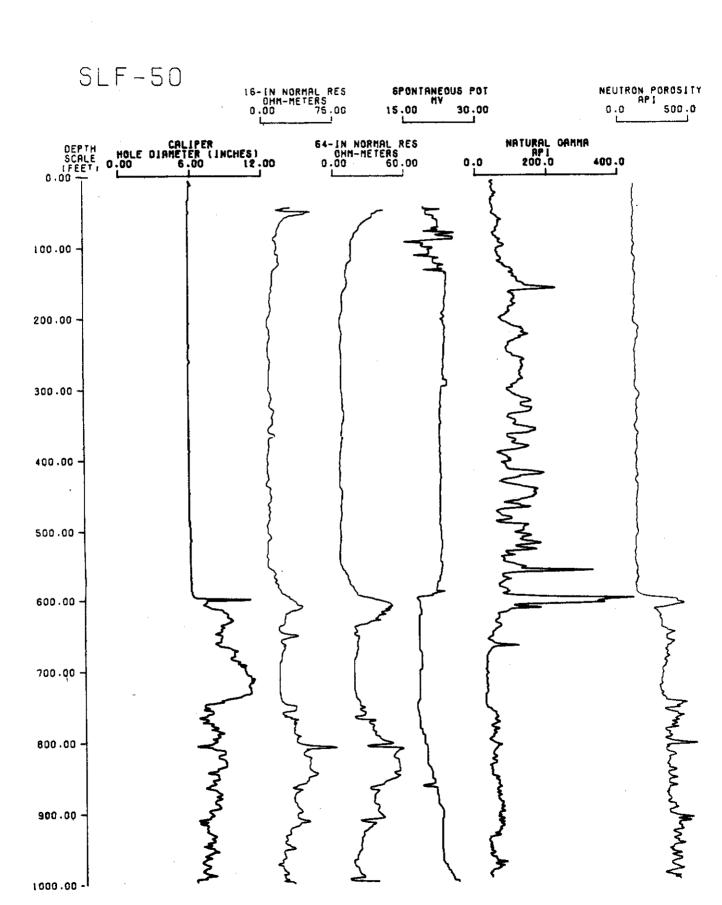
MANY CUNES

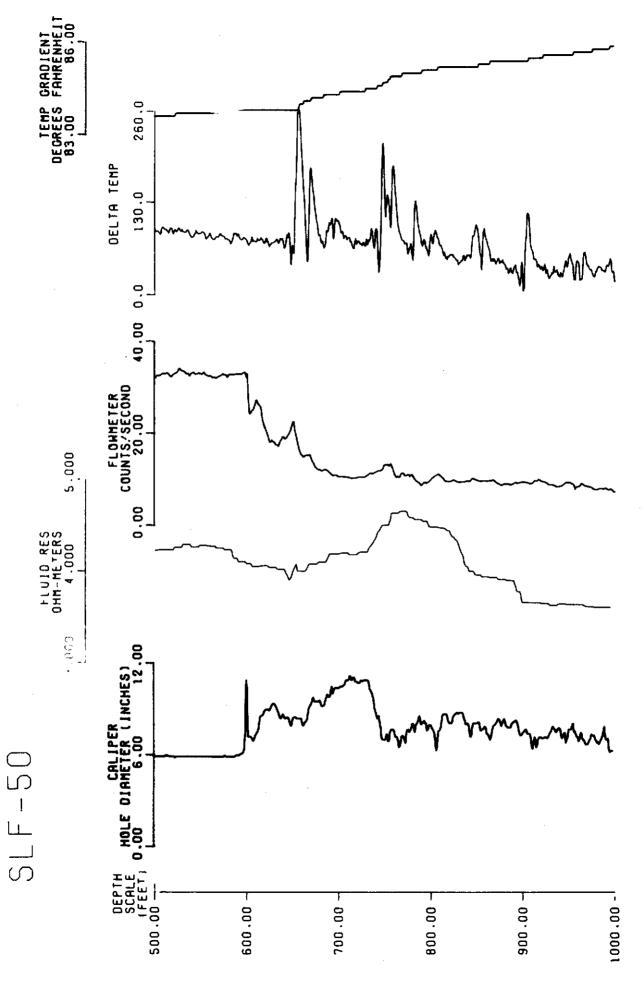
840.0- 650.0 AS ABOVE,

- 050.C- 060.0 LÍMESTŮNÉ, VERY LIGHT ÜRANGE TÜ GRAYISH ÜRANGE, 10%
 PÜRÜSÍTY, INTERGRANULAR, MGLÜLC, INTERCRYSTALLINE, GRAIN
 TYPE: BIÜGENIC, CALCILUTITÉ, CRYSTALS, 40% ALLÜCHEMICAL
 CÜNSTITUENIS, GRAIN SIZE: MICRÜCRYSTALLINE, RANGE:
 MICRURYSTALLINE TO CÜARSE, GLÜÜ INDURATION, CALCILUTITE
 MAIKIX, DÜLUMITE CEMENT, SPARKY CALCITE CEMENT, 30%
 DÜLÜMITE, BENTHÜNIC FÜRAMINIFERA, FUSSIL MELDS,
- BEU.U- B7C.O AS ABEVE,
- 870.0- 660.0 AS ABOVE WITH DICTYOCONUS COLKEL, AND CYPSINA SP.
- SOC.U- SGU.U AS ABOVED
- 890.0- 900.0 LIMESTONE, VERY LIGHT DRANGE TO GRAYISH DRANGE, C8%
 PÜRLSITY, INTERGRANULAR, MOLDIC, GRAIN TYPE: BIDGENIC,
 CALCILUTITE, CRYSTALS, 20% ALLECHEMICAL CONSTITUENTS, GRAIN
 SIZE: MICROCRYSTALLINE, RANGE: MICROCRYSTALLINE TO MEDIUM,
 GUOD INDURATION, CALCILUTITE MATRIX, DELOMITE CEMENT, SPARRY
 CALCITE CEMENT, 30% DOLOMITE, FOSSIL MOLDS, BENTHONIC
 FURAMINIERRA.
- 900.0- 910.0 CALCARENTIE, VERY LIGHT DRANGE, 13% PORUSITY, INTERGRANULAR, GEO INDURATION, CALCILUTITE MATRIX, SPARKY CALCITE CEMENT, BENTHONIC FORAMINIFERA, MOLLUSKS, ECHINCIC, CONES,
- 910.0- 910.0 DÜLÉMITE, GRAYISH BRÜWN, 10% PUROSITY, INTERCRYSTALLINE, INTERGRANULAK, MOLDIC, 50-90% ALTERED, EUHEDRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TÜ MICKUCRYSTALLINE, GUOD INCUKATION, DOLUMITE CEMENT, CALCILUTITE MATRIX, 10% CALCILUTITE, FOSSIL MOLDS,
- 915.0- 920.0 LIMESTONE, VERY LIGHT DRANGE, 12% PORCSITY, INTERGRANULAR, GRAIN TYPE: Blugenic, Calcilutite, Skeletal, 50% Alluchemical Constituents, Grain Size: Fine, Range: Mickecrystalline to Charse, Godd Induration, Calcilutite Matrix, Dolomite Cement, Sparry Calcite Cement, 15% Dolumite, Benthonic Foraminifera, Mollusks, Echinolo, Cones,
- 920.0- 930.0 LIMESTONE, VERY LIGHT DRANGE TO GRAYISH DRANGE, 08% PORDSITY, INTERGRANULAR, LGW PERMEABILITY, GRAIN TYPE: CALCILUTITE, DRYSTALS, 10% ALLOUHEMICAL CONSTITUENTS, GRAIN SIZE: MICROCRYSTALLINE, MANGE: MICROCRYSTALLINE TO FINE, GOOD INDURATION, CALCILUTITE MATRIX, DELOMITE CEMENT, SPARRY CALCITE CEMENT, 25% DOLUMITE, FOSSIL MOLDS,
- 730.0- 940.0 AS ABOVE,
- 940.0- 950.0 DOLOMITE, LIGHT GRAY, C82 PUROSITY, INTERGRANULAR, MOLDIC, LOW PERMEABILITY, 50-902 ALTERED, EUHECRAL, GRAIN SIZE: VERY FINE, RANGE: VERY FINE TO MICROCRYSTALLINE, GCCD INDURATION, DOLOMITE CEMENT, CALCILUTITE MATRIX, 202 CALCILUTITE, BENTHONIC FORAMINIFERA, FOSSIL MOLOS,
- 950.0- 956.0 AS ABOVE,
- 956.0- 96G.O LIMESTONE, VERY LIGHT CRANGE, 12% PORGSITY, INTERGRANULAR, GRAIN TYPE: BIDGENIC, CALCILUTITE, CRYSTALS, 4C% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: MICROCRYSTALLINE, RANGE: MICROCRYSTALLINE TO COARSE, GOOD INDURATION, CALCILUTITE MATRIX, DOLOMITE CEMENT, 15% DOLOMITE, BENTHONIC FORAMINIFERA, MOLLUSKS, ECHINOID,

- 960.0- 970.0 LIMESTONE, VERY LIGHT LRANGE TO LIGHT GRAY, COR POROSITY, INTERGRANULAR, LOW PERMEABILITY, MOLDIC, GRAIN TYPE: DIEGENIC, CALCILUTITE, 10% ALLOCHEMICAL CONSTITUENTS, GRAIN SIZE: MICKUCKYSTALLINE, RANGE: MICKUCRYSTALLINE TO FINE, GOLD INDURATION, CALCILUTITE MATRIX, DOLUMITE CEMENT, 25% DOLUMITE, FUSSIL MOLDS,
- 970.0- 980.0 LIMESTÜNE, VERY LIGHT ÜRANGE, 112 PURÜSITY, INTERGRANULAR, GRAIN TYPE: BIÜGENIC, CALCILUTITE, SKELETAL, 50% ALLECHEMICAL CUNSTITUENTS, GRAIN SIZE: FINE, RANGE: MICROCKYSTALLINE TO COARSE, GOLD INDURATION, CALCILUTITE MATRIX, DOLUMITE CEMENT, 10% DÜLOMITE, BENTHONIC FÜKAMINIFERA, MÜLLUSKS, ECHINUID.
- 980.0- 990.0 AS ABBVE,
- 990.0- 1000.0 LIMESTUNE, VERY LIGHT URANGE, 12% PURDSITY, INTERGRANDLAR, GRAIN TYPE: BIDGENIC, CALCILUTITE, SKELETAL, 60% ALLUCHEMICAL CUNSTITUENTS, GRAIN SIZE: FINE, RANGE: MICKECRYSTALLINE TO COARSE, GOOD INDUKATION, CALCILUTITE MATRIX, DELOMITE CEMENT, 10% DELOMITE, BENTHONIC FURAMINIFERA, MOLLUSKS, ECHINDID,

Geophysical Logs, Well SLF-50





Water Quality

	TUISS FE MG/L	75.	, 4 C	10.	44		. 4.1	5 ^ 5 ·	2.4	. N €	71.	1 - 3	ै •	0 •	44.
Cobe	Tular Fe Muzi	э А					3)**	• (0		77•1					
	FARDNESS MG/LCACC	442.4 442.4 5.62.5	7 4 4 4 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	171.6	141.2	3<0.1	2.145	44.000	1.4.4.0.1.0.1.0.1.0.1.0.1.0.1.0.1.0.1.0.	5.00	7 7 7 4 7	ξ (561.4	477.7	476.7
	ALK REG/L	មាល់ មានក្រុ មានក្រុ	2.72	₩	1.1C	3.24	20.7	2 * * Z 2 * 5 6	\$ C \$	2.66	76.5	2.57	4.73	5 • 5 5	2.5
	SE4 467L	63.7 97.1	175.6	41.6	2.7.5	Z-7C+	20.0	7997	167.6	169.6	172.6	160.4	208.2	161.3	172.1
	7/94 CT	212.6 325.2 426.6	1 1 1 2 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	158.2	112.1	451.4	5.462	7 L L L L L L L L L L L L L L L L L L L	7 C C C C C C C C C C C C C C C C C C C	876.1	1	1142.0	195.8	672.6	9*805
	3 C C C	22.57 31.60	3 3 4 5 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	12.50 13.50	0 • • 0 0 0 0 0	24.50	45.12	7) • 67)) () ()	0 * 0 D	101.80	0.4 0.4 0.4	22.70	63.70	95.50
	CA #61L	44.40 44.60.00	144.10	201.04 04	44.20	96.20	6.5 .	7 0 1 0 4 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	147.00	120.50	00.44 00.44	125.20	103.43	7 0 0 0	0 4 • 4 0
	3 K	~ A 1 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	2000	4 4 0 4 0 7	1.77 1.05	5 Y 5	6.17	7.00	7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7	23.61	15.06	14.24	3 • 4 0	€ € € €	46.44
S	7/9E	120-12	207.23	74.10	34.46	140.40	6.4. SET	130.00	04.063	445.48	39.+30.4 11.00.4	24.204	100.00	323.60	340.00
A. MISCELLANEOUS SAMPLES	TIME HUURAMIN	1350.	1260.	1100.	, 60001	• > 0 C 5	1360.	1910	1200	1345.	- 0.40 - 0.00 - 0.00	1200	430.	• 0011	. 744
A. MISCELLAP	DATE NG/DA/YR	10/27/31	2/16/06	4/27/62	10/14/102	1/19/03	10/23/01	10/////	10/7/71	10/63/01	15/ 5/81	1/ 5/62	70/51/01	20.197102.06	51 1/19/03
	STAFICA	लच	4			₹.		-	4 -4	*	ر بر د خ	4	٦ -		
	4.7	-423	1477	1477	C24-	47-7	1 5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	1	3-1 3-1	SLF-	71.7 11.7	25.1	3 T S-	24.	SLF-

_
ā
ш
⊃
z
=
Ş
NITIO
9
$\overline{}$
ഗ്ര
AMPLES
z
¥
-
S
Š
SUS
Ö
ш
₹
~
_
ш
$\overline{\mathbf{c}}$
Š
₹
-2
ż

IAIIUN Cuük	GATE MO/CA/YR	TIME HUUR, MIN	SP CEND UMHUS/CA	LAB CCND UMHUS/CM	LAB PH	1.565.50 FG/L	NC2 MC N/L		244 FG N/L	0 P	0PC4 MG P/L	NG3 RG N/L	1.015.SF #67L
111111111	10/27/81 12/27/81 12/27/82 13/12/82 14/27/82 14/27/82 10/14/82 10/14/82		1180. 21710. 2170. 510. 1272.	122 122 123 123 123 123 123 123 123 123	200 4 4 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	o o	.151		.16		541 545 541	• 256 • 709	890.0 1589.0 1545.0 1546.0 1580.0 336.0 1212.0
	1 10/27/61 1 10/27/61 1 12/ 2/61 1 12/ 2/61	3300. 1315. 1430. 1500.	1355. 1530. 2246. 2276.	1380. 1480. 2250.	7.59		.211 .219 .004 .004	v v	101H	• • •	650 653 0123 017	457. 457. 9960.	1027.0 979.0 1500.6
1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	10/27/81 12/ 2/81 12/ 2/81 1/ 5/82	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	• • • 	4 4 2 4 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	7.07		* * * * * * * * * * * * * * * * * * *		. 54	v v			2182.0 2801.0 2522.0 2245.0
L - 51	1 10/14/82	930.	1270.	1330.	7.74	12.6		·					932.0 1436.C
15 -47	58/61/1 16	945.	-9597										2104.0

LAIL FEET No.1																				
Lily CEPTH	ARDNE S G/LCAC	13.	92.	382	33.	, ,	18.	5.5	54.											
CATE NA	4 L	7.00		000		200 1	œ	-0	æ											
LAILE DEPTH NA K CA HG/L HG/L HG/L HG/L HG/L HG/L HG/L HG/L	ပ္သပ	00 00 60 00 00 60	90.	50.	70.	9 6	ימי מיני	72.	64.	.DIS.S	005. 696.	109.	000	104. 506.	726.	. 1.59	541.	550	734.	905. 739.
LATE PEPTH NA K CA CALLALIYK FEET NA LA HGAL MGAL MGAL NGAL NGAL NGAL NGAL NGAL NGAL NGAL N	1/9# CI	933	06.	34.	20.	• • • •	35.	20. 35.	33	DISS S FG/L		2.00	2.20		0.70	1.60	0.70	1.90	2.60	3.4°C
LA1E DEPTH NA K CALLLALLY FEET HULL LAB LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAUGH GOLD LAGLE LAGLE LAUGH GOLD LAGLE LAGL	ΣU	3.C	7 2	200	. 6.	, 00 d	9.4	7.5	4.4	D155 F MG/L	0,3	0,0		• •	• •	•	•		•	30
LAIL DEPTH NA K LLIA/14 FEET M6/L M6/L LLIA/14 FEET M6/L M6/L LLIA/14 FEET M6/L L5-60 LLIA/14 FEET M6/L L5-60 LLIA/14 FEET M6/L L5-60 LLIA/14 FEET M6/L L5-20 LLIA/14 FEET M6/L L5-20 LLIA/14 FEET M6/L L5-20 LLIA/14 FEET M6/L LAIL-60 LLIA/14 FEET M6/L LAIL	₹ 5	0.8.0	10.1	98.1	7.2	* O *	0.40	03.3	99.5	A 8	.40	G 0	9	64.	.96	.80	52.	.75	•73	730
LAIE DEPTH NA HULLLATYR FEET MULL HULLLATYR FEET MULL HULLLATYR FEET MULL HULLLATYR FEET MULL HULL HULL HULL HULL HULL HULL HULL	K RG/L	5.25	100	4.00	28.4	1 . d 2 . d 4 . d	7.6	7.0	2.4	CLN JS/C	550	3 C	250	200	950	. 7	72	9	10	C 2
######################################	2. 9	0.4	30.0	17.6	0		9 0	93.2	80.6	P COND MHUS/C					2.6)				4320.
322222222222222222222222222222222222222	DEPTH FEET	715 740	800 800 800	820 840	860 870 880	8 00 C	940	09 66 66	1000	DEPTH FEET	715	760 780	800	840	860 870	88 8	8 8 8 8	940	9 86	100
322222222222222222222222222222222222222	CATE HC/DA/YR	יו פו פי	9 90 90	ap 10 :	စ္ကေတာ့ ႏ	D CD C	တော	90 60	2/ 3/82	DATE MO/LA/YK	0.70	/26/8	120/8	126/8	126/8	3/8	3/8/6/	3/8	9/E/) 3/8 3/8
	SIATION	300	333	200	200	900	, ú	5 0 0 0	2	يد د	22	.	120 4	מי ה	~ ~	Ŋ	v r	S	3	טיט

TATICN CODE	DATE MO/CA/YR	DEPTH FEET	NA MG/L	* #6/L	CA MG/L	7/9W WG/L	1/9# #6/L	SD4 M6/L	ALK MEG/L	HARDNESS MG/LCACD	
LF- 50		382,	478.00	15.60	113.60	83.20	882.7	159.6	2.57	626.1	
LF- 50		.026	491.60	15.60	119.50	81.60	898.4	187.3	2.46	634 • 3	
LF- 50		,006	394.80	13.20	100.40	72.40	751.2	151.1	2 • 65	548.7	
LF- 50		820,	418.40	14.00	104.30	73.60	781.7	161.0	2.64	563.4	
LF- 50		785	365.20	12.40	96.60	65.20	677.6	151.9	2.71	464.6	
LF- 50		750,	385.60	13.20	98.70	70.00	7.047	152.5	2.68	534.6	
LF- 50		.099	438.40	14.80	105.00	74.80	809.1	178.7	2.60	570.1	
LF- 50		T0C	470.00	16.00	108.70	78.00	838.5	182.6	2.57	592.5	
TATION		DEPTH	LAB COND	LAB PH	TDISS FE	TDISS SR					
CGDE		FEET	UNHOS/CM		N6/L	MG/L	1/9H				
1F- 50		995,	3550.	7.53	.05	12.800	2200.0				
LF- 50		950	3550.	7.45	•0•	15.400	2220.0				
LF- 50		,006	2950.	7.45	• 05	12.700	1834.0				
LF- 50		820,	3050.	7.42	.07	13.100	1933.0				
LF- 50		785	2650.	7.47	90•	13.000	1591.0				
LF- 50		,20,	2875.	7.41	• 0•	13.100	1732.0				
LF- 50		.099	3150.	7.39	•0•	13.400	1911.0				
LF- 50		T0C	3275.	7.40	• 03	14.200	2136.0				

HARDNESS MG/LCACE	599.2 590.6 573.0		611.2 571.1	578.6 583.3	594.0 611.8 615.2
ALK REL/L	2.74 2.83 2.76	2.65 2.68 2.44 2.85	2.73	2.68	
. SU4	216.8 216.9 217.8	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4		174.2	
1/9W CL	1046.2 1037.9 1031.5	1006.6 974.9 1019.2 1000.2	540.E	817.1 806.7	
1/9W W6/L	80.00 78.10 74.24		35.00	76.80	00.66 00.66 00.66 00.66
CA MG/L TEST 1	138.10 107.80 107.10	70.20 111.50 113.10 30.80	105-10 105-10 TEST 4	105.83 107.00	111.30 113.60 113.30 113.90
K MG/L PLEPING	15.52	PLMPING 9-10 15-20 15-60 10-30	DOCTO DOCTO DOCTO DOCTO DOCTO PUMPANG	13.60 13.60 PURFING	4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4
NA MG/L	456.00 444.40 444.40	244 464 4664 5664 524 666	423.00 413.00	24 00 04 04	77777777777777777777777777777777777777
TIME HOURS MIN	151c. 160c. 12c.	1136. 1336. 136. 136. 136.		* * 0 9 8 9 9 4	* * * * * * * * * * * * * * * * * * *
DATE MC/DA/YR	1/ 5/02	1/19/82	21 2132	7810T12	0/54/35 0/55/85 8/56/05 0/27/32
STATIEN CUDE	SLF- 50 SLF- 50 SLF- 50	SEF- SEF- SEF- SEF- SEF- SEF- SEF- SEF-	SLF- 90	2111	SLF- 51 SLF- 51 SLF- 51 SLF- 51

STALLUN Cude	Z.	DATE MUZDAZYR	TIME HOURSHIN	CENT	SE CCND LAB UPHOSICE UMED PUMPING TEST 1	LAB COND UMHGS/CM EST 1	LAB PH	1.5US.SC #6/L	TUISS FE MG/L	TD155 SR PG/L	T.015.50 MG/L	ALK MEG/L
SLF- SLF- TT-	7 7 7 7 7 7 7 7	20/4 /1 20/4 /1	1510. 1500. 1710.	7 · 0 ? 7 · 0 ? 7 · 0 ?		4410. 4500.	7.527.427.42		A 20	15.440	1896.0 1931.0 1791.0	2.74
2	2200	7871271 787671 797671 797671 797671		5 H J O O O O O O O O O O O O O O O O O O	PURPING 1 3650. 3610. 3762.	1651 2 3240. 3480. 3480.	7.30	N N N 4	7 V V		2004.0 2085.0 2065.0	20 Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z Z
SLF-	7 O 1 A	79/7/7			- 2456	1EST 3 3500. 3500.	7.48		20° ×	12.500 13.6CC	2115.0 2127.0	2.73
1110	2 2 2 2	20/7/12 2/4/32	• 0 0 8 4 4 • 0 0 8 9 • 0 0 8 9	24.1	PLMPING TEST 4 5290. 3350.	EST 4 3250. 3350.	7.37		< .02 < .62	12.400 12.900	1883.0 1896.0	2.68
1111	4 4 4 4	6/24/62 6/25/62 8/26/62 0/23/62	* * * * * * * * * * * * * * * * * * *		* .02 3	651 5 88 86 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	7.34 7.44 7.45 7.45	1979.0 2286.0 1936.0 1997.0				

E. INJECTION TEST SAMPLES

HARDNESS MG/LCACO	496.2 426.0 475.2 492.8	TDISS SR MG/L	1.760 1.510 1.670 1.710
ALK MEO/L	4.27 3.07 3.89 3.75	TDISS FE MG/L	.20 .15 .25
\$04 MG/L	224.0 186.2 217.2 216.9	TOTAL FE MG/L	5.24 5.53 39.52 422
1/9W	182.7 148.9 197.0 196.0	T.SUS.SD TO MG/L	2.0 3.0 1.0 280.0
NG MG/L	23.50 20.20 22.50 23.40	LAB PH	7.34 7.36 7.61
CA MG/L	160.00 137.30 153.20 158.84	LAB COND UMHOS/CM	1420. 1280. 1425. 1465.
7 % Y	8 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9 9	SP COND UMHDS/CM	1230. 1285. 1240. 1300.
NA MG/L	108.90 103.80 112.70 118.70	TEMP	4 6 9 11 8 6 7 6 9 1 8 9
TIME HOUR, MIN	1030. 1630. 1730. 1911.	TIME HOUR, MIN	1030. 1530. 1730.
DATE MO/DA/YR	10/19/82 10/20/82 10/21/82 10/21/82	DATE MG/DA/YR	10/19/82 10/20/82 10/21/82 10/21/82
STATION CODE	St S- 1 St S- 1 St S- 1 St S- 1	STATION	\$15-1 \$15-1 \$15-1

* BACKFLOW FROM INJECTION WELL

m
SAMPLES
۳,
=
ᄟ
5
>
7
ဟ
<u></u>
(J)
ES
TEST
•
>
Œ
E
=
_
O
Ō
RECO
=
ŭ.
ш

HARDNESS MG/LCACE	7.	93.	9.6	٠ دي دي	.05	000	4.5	. 40	63.	CB.	80	11.	10.	€ 9	60.	58.	5 c	40.	440.2	12.	17.	20.	22.	17.	40	15	17.	18.	74.	11.	20.	22+	15.	80	12.	10.	9	C G	11.	55.	
ALK PEG/L		4	•	4.	•	~	Ċ.	9	7	ر <u>ت</u>	٠	~	-		٦.	7	7	•	4.11	Ţ	ပ္	7	<u>ۍ</u>	ပ	ω	5	Š	,	Ċ	•	٠,	بد	3	¥,	•	æ	٥.	۲.	۲.	۲.	
SC4 MG/L	02.	02.	.00	C7.	90	01.	. o	96	06.	ن د ج	٦.	90	ct.	62.	0	0.4	12.	45	206.2	.20	C.1.	07.	C 2.	07.	56	63.	5	04.	95.	04.	69.	• • 0	550	0.4	98	04.	02.	55	63	04.	
7/9.4 CL	07.	17.	źt.	26.	30.	30.	30.	60.	61.	65.	7,5	01.	90.	24.	49.	55.	52.	ί.	364.3	81.	ე	86.	00	11.	5 E•	20.	1C.	21.	35.	37.	41.	4 8•	4	4.0	67.	73.	€8•	61.	54.	74.	
#6 #6/L	٠,	7.6	0.7	Ç.	Ç. 4	2.1	7.4	4.6	4.1	6.3	7.6	7.0	7.4	Ġ.	:	7.1	7.1	1.1	38.10	2.0	2 . 2	3 . 5	2.9	3.2	5.7	4.7	, ; ·	4.6	4.5	*	5.0	۵. د	4.0	4.6	J. C	5.0	ž•¢) • ¢	5.4	3.4	
CA #6/L	0.44	46.5	48.4	48.4	46.3	47.2	44.8	9.54	い・ちゃ	43.B	42.2	40.2	46.7	18.2	4.1.2	22.5	7.03	11.7	113.50	38.0	37.B	37.3	38.5	32.9	31.5	34.3	14.3	34.1	44.6	31.1	13.1	33.6	31.5	30.1	31.0	30.4	22.1	59.3	56.6	3.0	
**************************************	7	S.	`	٦.	٠,	7	'n.	0	Š	7.	٠.	7	٠,	0	7	7	4	4	5.14	,	7	1	5	4	41	۲.	,	۲.	۲.	0	7	•	*	٧.	*	4.	٧.	Ť	3	٥	
7 A A B C / L	27.0	•	38.0	36.6	39.0	40.6	φ. 0	55.C	4.5	4.5	60.0	77.0	72.4	70.0	7.40	0 y . C	be.	40.0	230.06	92.6	4.75		しゅたか	4.10	4.70	Ú. 90	16.0	4.47	10.4	21.0	7.17	4.61	4.97	35.C	26.0	26°C	79.8	34.6	20.07	15.1	
TIME HOUR, MIN	2 to 1	2	750	703	'n	.E ○ 4	Ę)	703	350	501	:0	'n,	0	5	5	735	105	405	1705.	7.7	40.0	2	733	Ċ	3	43	707	331	504	3	4)	2	400	2	300	20¢	S	3	0	3	
DATE ML/CA/YK	3/52/*	1/25/1	112918	1/36/6	1/36/0	1/3E/1	1/36/1	1/36/0	2/ 1/3	P/T /7	211 12	017 12	11 610	617 17	0/7 /7	0/5 /2	0/5 /7	6/ 3/3	79/5 /7	0/3 /7	2/ E/B	2/9 /7	611 113	110	0/1 /7	9/1 /7	2/ 6/0	2/3 /2	8/2 /2	0/0 /7	9/5 /7	2/4 /7	21 510	0/4 /7	2/17/12	8/01/7	4/13/0	2/57/7	135/0	2/22/3	
STALLAN	1. F-	4	(F- 5)	Lf- :	LF- 5	LF- 5	rr- 5	7 -17	1.f- 5	UF- U	1.F- 5	LF- 5	rt- 5	11-12	LF- 5	1-1-	7 -1-1		SEF- 50	7	F- 5	4 -41	-17	11- 5	U	the t	LF- 3	1.6- 2	LF- 5	LF- 5	LF- 5	1F- 5	LF- 5	LF- 5	£ - 51	6 -37	LF- 5	4 -71	LF- 5	Lf- 5	١

F. RECOVERY TEST SAMPLES (CONTINUED)

HARDNESS PG/LCACE		318	165	524	267	521	521.	513	524	523	522	530.	500	346	511.	523	520.	503	440	529	495	513.	545	541.	527.	526.	515.	546.	569.	565.	570.	567.	562.	553.	571.	567.	573	, 40 , 40 , 40	560.3	576	
ALK MEC/L	٠,		. '-	-	'-		'-	v	-			٠,	9	÷	_	4	ď	m	4)	87	_	ď	4	4	'n	Ň	w	ú	Ň	Ç	0	0	0	Ġ.	9	0	· [~	- E	2.70	• 30	9
\$C4 #67L	, 6	60	7	07.	ot.	E C	97	ġ	0	,	(n)	3	5.	9	5	3	. 40	4.	96		,	5.	Š	7	35	57.	÷	E	į		ŏ	ë	Š	8	N	7		9	164.4	0	ŗ
7/9# CL	6	7.7	67	83.	ÉВ	55	. 76	P.	0.5	7.4	3.5	2 C.	13,	30.	36.	. 40	20.	99	*	53.	00	76.	27.	41.	81.	9	50.	11.	22.	9	57.	77.	87.	94.	33.	5.	90	7	4006	6	
MG MG/L			رب د		7.6	.D			4	3.5	7	7	4	÷.	•	:	5.7	3.4	3.5	41	30	4.	4.4	÷	•	9.	5.	•	Đ.	9.		. 2	•	٠.	ع •	.0			68.40	7	•
CA #G/L	0.4	5.	22.1	3 4	-0	24.9	36.E	28.3	30.5	30.6	25.6	E • 6 7	27.4	80.8	28.0	3 C • E	31.4	30.1	0.60	28.7	17.9	7.7	20.7	16.4	12.0	12.5	12.0	17.7	9*67	14.9	16.2	16.5	12.3	10.6	17.1	15.3	16.7	14.7	111.00	15.0	
7 × 4 × 4 × 4 × 4 × 4 × 4 × 4 × 4 × 4 ×	-	1		~		7	~	7	"	٠,	٠,	*:			4	٠,	٠,	7	17	9	ď	<u></u>	41	~	5	ō	e.	6.7	7.0	0	1.2	1.3		1.3	1.8	φ. α	4.5	2 • 1	12.20	2 . 4	·
NA BOLL	44	41.	2407	4.74	34.6	47.6	23.6	50.0	30.00	7:5	20.4	75.0	9.79	D	73.6	200	72.5	5a.t	27.0	77.t	27.6	07.2	91.6	9.66	05.b	70°B	92.0	717	44.0	0.84	0.40	5.50	54.8	74.0	76.8	77.6	76.4	33.2	340.46	4.40	.3
Time Hours min	၁	7	ာ	3	~	<u>ာ</u>) TE	717	၁	J 0 0	S	302	7	3	ر ب	3	GT Q	<u>د</u>	3	ÜJC	3	2	2 00	0	700	6 J C	၁	700	30 2	3	ე <u>ი</u> ~	0	2	္က	007	2	200	2	700.	2	0.00
DATE ML/LA/TR	8/51/	_	8/57/	114/6	11510	11215	/15/B	A/CT/	/10/0	9/97/	110/8	115/0	11118	9/11/	9/17/	2117/	21171	12210	91771	16613	8/KT/	R/61/	12019	15018	121/8	/51/6	12410	124/8	125/8	6/57	126/8	197	127/8	127/8	1/8	1/6	218	218	2/ 3/83	3/8	773
STATION CLDE	<u>.</u>	1	F-	-1-1	1	<u>-</u>			· + 4.1	-4-	1	1	1	1	1	1		ļ	1	1 1 4	1	-f-	<u>.</u>	F 5	7	Ţ.		7.	T.		7	Ŧ.	1	7.	7	4	1	F .		<u>ل</u> ب	4

F. RECOVERY TEST SAMPLES (CONTINUED)

				^						
SIALLEN		TAME	, U	P CGN	AB CON	LAS	⋖	1	S	I S
CLUE	3.	HUUKARIN	CENI	ピン/ 50 カルコ	UNHOS/CR		F6/L	H6/L	1/94	1/91
LF- 5	0.41/69/0		ů	334	500	٠		•	99.	75.
11-12	0 11/25/10	2	ò	404	52	7	0	•	72	2
4 -11	2/57/11 7	Š	Š	422	563	٠	۲,	٠.	75	. p
LF- 5	9/15/TT 0	0	<u>, </u>	4.02	Š	٠	+	٥.	7.5	ۍ د
Lf- 5	8/18/11 0	. 47	ċ	423	540	٠	B	ં.	85	٠ ال
ا س	0 11/30/6	9	ò	7	35	•	- 3	£0.	36	60
11-	0 11/3c/b	5	ò	55.2	580	٠		•	46	000
. i-i-1	8/36/TT 0	3	ċ	004	540		ಾ	٠,	• 6 t	ů.
11-17	57 17F	1	ò	40.5	640	7	*	٠,	5.00	585
1F- 5	D/T /2T O	•	ò	3	5	7		<u>ن</u>	4	C32.
1 - 1 1	8/T /7T 0	3	ċ	7	650	٠.	4	਼	15.	683.
F- 1	0 12/ 2/8	1	ċ	3	730	٠	<u>د</u> ک	٠,	. 7.	C46.
-H-	171 7 TE	O	٠	ŭ	700	٠,	.3	٠.	. 14	C55
-F-	0 121 6/9	402	ن	562	240	਼	7	•	7.	317.
.F- 5	8/7 /7T C	2	ò	5	750	٠,	٠		77.	282.
LF- 5	u 121 310	3	ن	665	75C	7	J	•	77	160.
LF- 5	0 16/ 3/8	0.1	٥	909	766	•	į	•	.72	219.
c -47	U 121 3/8	10.7	•	613	800	٠	~	ာ.	'n.	23 B.
7-11	0/5 /77 7	20	Ĵ	40	40		J	•	5)	22b.
LF- 5	12/ C/8	001	9	010	502	ζ.	7	ಾ	-73	151.
L.F	0 101 010	ز	ن	ę	450	4	•	4	2	168.
0 -11	C 12/ 6/8	200		S	460	ı,	4.	¿)•	, 1 ¢	244.
LF- 5	C 121 113	325	•	t b	460	ب	43	O	.18	213.
LF- 5	U 12/ 7/8	2	٥	676	442	'n	41		£ 0.3	197.
LF- 5	C 121 7/8	ာ	3	~	468	-		ು	• 1 e	135.
LF- 5	J 12/ 7/0	2	٠	30.5	550	,_	÷	70* >	£ .	232.
LF- 5	2/9 /71 9	O	•	735	580	1~	_	30.	4.3	204.
LF- 5	6 121 E/B)		2	562	~	73	C	. E.E	257.
LF- 5	0/0 /77)	a	÷	712	079	T.	5	0	. 14	221.
LF- 5	P/9 /71 7	202	å	740	548	D	S	90.	25	255.
LF- 5	0 151 5/8	an.	• •	7	560	T)	~	S	.63	292.
11- 5	12/ 5/0	107	٥	157	909	ນ	Ð	O	.21	225.
. F-	0 151 5/0	4	٥	350	0.40	ນ	D	45.	٠ د د	250.
16. 5	016 177 0	3	٠ ن	~	710	E.	Ω	70.	115	235.
LF- 5	0 12/10/8	200		770	960	3	L)	0	84	245.
LFT 5	0 12/16/8	30	•	17	700	3	T)	40.	40	29B.
LF- 5	0 14/13/0	ပ္ပ		4	550	ې	CT1	-	9.5	245.
SLF- 24	161121 0	1360.	20.6	176ë.	1670.	7.17	.65	• 62	4.066	1275.0
1F- 5	J 12/13/0	S	9	5	040	. 4	•	.52	7.6	290.
LF- 5	1 16/13/3	S.	٥	ã	0.4°	ď	79.	* 0*	21	28 6 •

F. RECOVERY TEST DATA (CONTINUED)

1.015.SL	177.	0	7 2	1363.0	243	435.	314.	369.	460.	432.	435	456.	59B.	£0£.	c>e.	574.	552.	472.	500.	542.	546.	750.	82C.	618.	652	736.	620.	706.	70¢.	746.	728.	722.	774.	784.	594.	736.	000	200	. 4	 	1812.0	; ;
TC155 SR PG/L	. 17	,	ις: O	5.210	41	7.7	• 56	÷	.	• 0¢	CE	.58	: 52	.1¢	(T)	. 82		36	42	Ç	53	6.360		.42	.5.	.51	20.	H.	54.0	C.5C	C.86	C. 66	C.92	6.96	2.84	2.66) (C)	1 4	7 6	9 4	13.420	!
TDISS FE MG/L				.63		C				~											60.	20°	• 15	•10	£9.	• 03	•05		•03	*C4	£0.	03	ຫ ວ•	* C8	60.	40	9	9 0) (* } (*)) O	. 22	
TOTAL FE MG/L	CC.	2	~	90.	~	(7)	ൗ	o	2.3	• 2 4	CO.	\sim	į		O.	a	3	4	*	14.	ري	• 65	Ġ,	~	e)	ာ	S	• 29	£0.	• 24	> €0 •	_	7	• 25	_	.26	· N	22.	7.5		47.	
ГАВ РН			7	2	7	.11	4	*,	٥	٠	٠	۳,		٠,		Ü	9	÷	•	~		3	٠,	~	5	.72	۲.	7	Ö	7	ڊ	٠,	Ò	٠	~	30	Ö			• 0	7.16	
LAG COND	3	185	920	- 2⊃	3	3	S.	960	S	900	145	5	040	020	920	9	080	000	5	050	7	320	9	054	040	000	540	100	3	5	4 4	7	8	4	4	0	9	2		Ö	2440.	
SP CLNU UMHúS/CP	613.	277	27.7	r)	871	818	513	4 4	ŝ	822	627	643	845	740	623	715	715	74C	7 <u>i</u> 4	702	688	624	724	912	226	96	259	306	685	032	C 8 7	950	113	740	~	218	4	(C)	24.0		2207.	
TEMP	J		_	÷	ು	٠.	-1		. 1	-1				٠,	\sim	~	\sim	\sim	~	~	—	~	~	`-	•	~	-	~	•	•	\sim	\sim	~		~		~	.n	T)	-	28.0	
TIME HOURS BIN	,107	071	ر. دن	1700.	710.	Э.	1310.	-	760.	TCCC.	$\dot{\circ}$	1700.	*	1000	る	7	-	100.	•	S	J	~ C	700.	O	0	160C.	•	7.0	100.	0	0	•	0	1630.	1100.	1700.	100.	1766.	~	1700.	736.	
LATE MC/LA/YR	2	7	9/51/7	14117	1271	7	101	2115/13	110/0	2/15/0	116/3	-	0/11/7	11	7/7		P/17/7	17717	12218	16610	11918	9/57/	12018	12613	12119	21/8	12410	12418	D	8/57/	12010	726/8	15118	12112	9/1 /	9/1 /	1 218	8/7 /	3/8/	2/ 3/83	9/4/	
STATION Cour	1 1	<u>.</u>		1.1.1	LF-	-41	L F-	LF-	1	LF-	111		<u>+</u> -			<u>.</u> f	4	1	L F-	1	<u>-F</u>	- 4-1	1	LF-	7	LF-	- 4-	7	- - -		1	1	<u>.</u>	1	1	7	<u>+</u>		1	SLF- 50	-	

G. PACKER TEST VALUES

HARDNESS MG/LCACL	621.0 594.2 589.7	592.8 611.7 614.3			
ALK MEG/L	2.62 2.24 2.56	2.26 2.40 2.53	vs	PACKER PACKER PACKER	ELLHEAD LL LL
504 MG/L	176.2 161.2 177.4	184.2 199.2 193.0	COMPENTS	BELON ABCON BELON P	COMBO WELLHEAD OBS. WELL OBS. WELL
7/9# C1	1025.7 886.2 1015.3	954.9 1021.5 1077.8	TDISS SE MG/L	16.400 15.600 15.600	15.400
1/9# #6	07.77	78.80 80.00 81.60	TDISS FE MG/L	8 7 8 0 0 0	n n n
CA #6/L	114.20 112.70 108.90	107.50 113.10 111.50	T.DIS.SU MG/L	1474.0 2056.0 1488.0	2379.0 2143.0 2180.0
H6/L	14.80 16.00 13.60	14.80 15.60 16.00	LAD PH	7.63 7.53 7.55	7.45
NA MG/L	485.00 502.40 451.20	476.80 501.60	LAB CLNC UMHES/CM	3450. 3460. 3200.	3325. 3500.
TIRE HUUKARIN	1035. 1200. 1335.	1500. 1200. 1335.	TIME	1635. 1200. 1335.	1500.
	3/12/02 3/12/02 3/12/02	3/12/62 3/12/82 3/12/82	DATE MG/DA/YR	3/12/82 3/12/82 3/12/82	3/12/82 3/12/82 3/12/82
STATIEN Cúde		944		SLF- 50 SLF- 50 SLF- 50	0 T 4
STAT	SLF- SLF- SLF-	SLF- SLF- SLF-	STATIGN	SLF- SLF- SLF-	51.F- 51.F- 51.F-

Aquifer Test Data and Analyses

APPENDIX 4a. Aquifer Test No. 1

PUMPED WELL: SLF-50

DATE: 1/5/82 to 1/6/82

CASING DEPTH: 600 ft

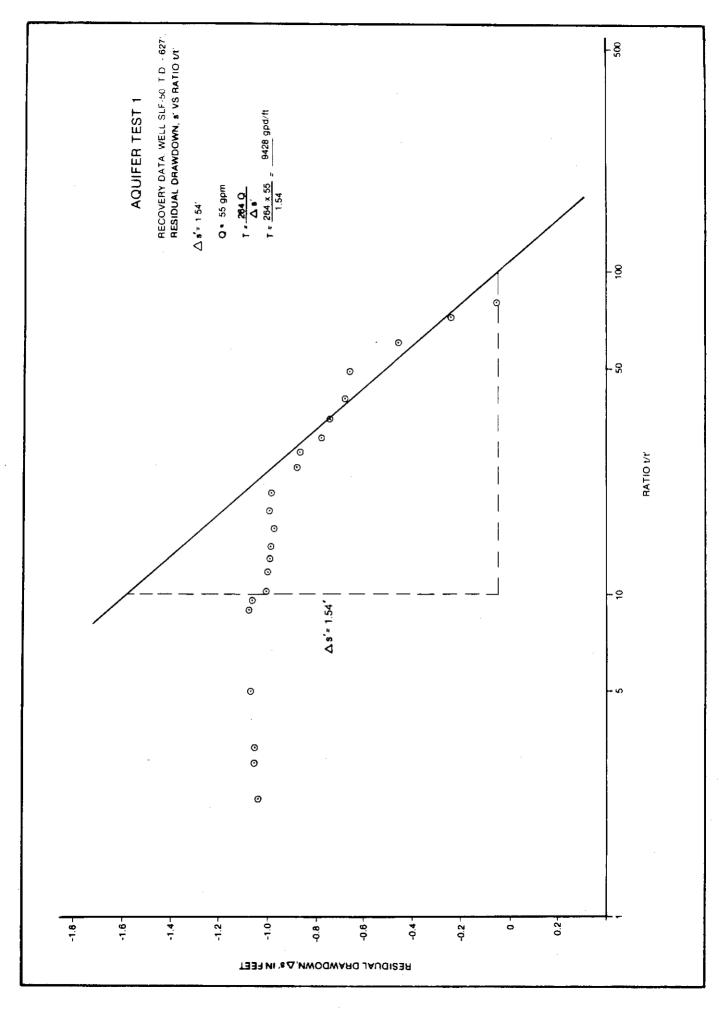
WELL DEPTH: 627 ft

WELL DISCHARGE: 55 gpm

RECOVERY DATA, WELL SLF-50

TIME SINCE PUMPING STARTED t (MINS)	TIME SINCE PUMPING STOPPED t'(MINS)	RATIO t/t'	DEPTH TO WATER (FT ABOVE MP)	RESIDUAL DRAWDOWN s'(FT)
121.5 122.0 122.5 123.0 123.5 124.0 124.5 125.0 126.0 127.0 128.0 129.0 130.0 131.0 133.0 134.0 135.0 150.0 180.0	1.5 2.0 2.5 3.0 3.5 4.0 4.5 5.0 6 7 8 9 10 11 13 14 15 30 45 60	31.0 61.0 49.0 41.0 35.3 31.0 27.7 25.0 21.0 18.1 16.0 14.3 13.0 12.0 10.2 9.6 2.0 5.0 3.7 3.0	8.28 8.69 8.88 8.90 8.97 9.00 9.08 9.10 9.20 9.21 9.22 9.22 9.22 9.22 9.22 9.22 9.23 9.29 9.30 9.29	-0.05* -0.46 -0.65 -0.67 -0.74 -0.77 -0.85 -0.87 -0.98 -0.98 -0.98 -0.99 -1.00 -1.06 -1.07 -1.06 -1.05 -1.05
210.0	90	2.3	9.27	-1.04

*Note: Minus sign indicates recovery above initial static level.



APPENDIX 4b. Aquifer Test No. 2

PUMPED WELL: SLF-50

DATE: 1/21/82

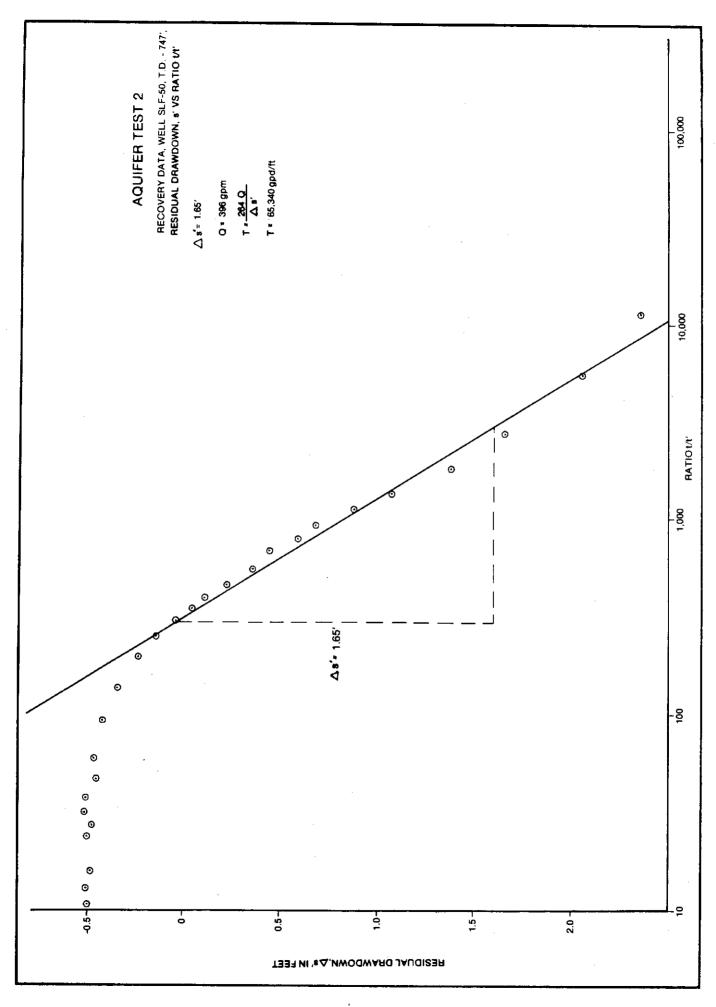
CASING DEPTH: 600 ft

WELL DEPTH: 747 ft

WELL DISCHARGE: 396 gpm

RECOVERY DATA, WELL NO. SLF-50

STARTED STOPPED RATIO DEPTH TO WATER DR t (MINS) t'(MINS) t/t' (FT ABOVE MP) s'	
2790.25 0.25 11161 6.39 2790.50 0.50 5581 6.69 2791 1.0 2791 7.09 2791.5 1.5 1861 7.39 2792 2.0 1396 7.69 2792.5 2.5 1117 7.89 2793 3.0 931 8.07 2793.5 3.5 798.1 8.17 2794 4.0 698.5 8.31 2795 5.0 559 8.40 2796 6.0 466 8.54 2797 7.0 399.6 8.63 2798 8.0 349.8 8.69 2799 9.0 311 8.80 2800 10.0 280 8.82 2801 11.0 254.6 8.90 2802 12.0 233.5 8.93 2803 13.0 215.6 8.95 2804 14.0 200.3 8.99 2805 15.0 187.0 9.01 2810 20.0	32.56 2.36 2.36 2.06 1.66 1.36 1.06 0.58 0.44 0.21 0.05 -0.07 -0.15 -0.24 -0.24 -0.24 -0.24 -0.33 -0.24 -0.33 -0.39 -0.45 -0.50 -0.50 -0.50 -0.50 -0.50



APPENDIX 4c. Aquifer Test No. 3

PUMPED WELL: SLF-50

DATE: 2/2/83

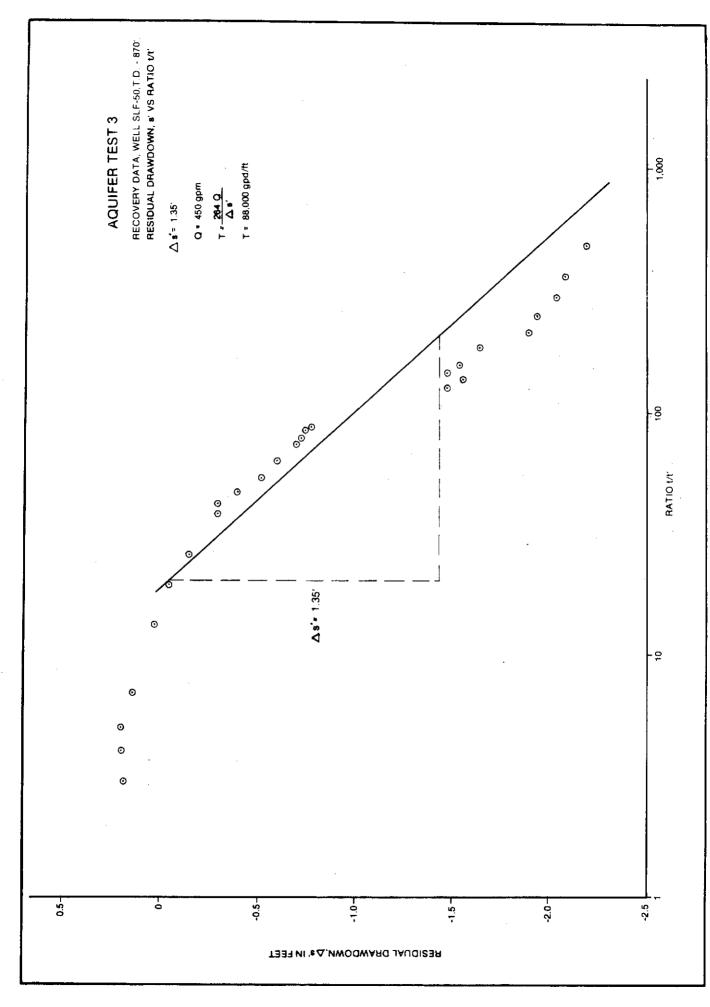
CASING DEPTH: 600 ft

WELL DEPTH: 870 ft

WELL DISCHARGE: 450 gpm

RECOVERY DATA, WELL NO. SLF-50

TIME SINCE PUMPING STARTED t (MINS)	TIME SINCE PUMPING STOPPED t'(MINS)	RATIO t/t'	RESIDUAL DRAWDOWN s'(FT)
364	0	. .	26.90
365.00	1.00	365.00	2.07
365.25	1.25	292.20	2.02
365.50	1.50	243.66	1.92
365.75	1.75	209.00	1.88
366.00	2.00	183.00	1.63
366.25	2.25	152.77	1.53
366.50	2.50	146.60	1.46
366.75	2.75	133.36	1.55
367.00	3.00	122.33	1.47
368.25	4.25	86.64	0.76
368.50	4.50	81.88	0.74
368.75	4.75	77.63	0.72
369	5.0	73.80	.69
370	6.0	61.66	.59
371	7.0	53.00	.51
372	8.0	46.5	.40
373	9.0	41.44	.30
374	10.0	37.40	.30
379	15.0	25.20	.15
384	20.0	19.20	.05
394	30.0	13.13	+.03
424	60.0	7.06	+.14
454	90.0	5.04	+.18
484	120.0	4.03	+.19
544	180.0	3.02	+.17



APPENDIX 4d. Aquifer Test No. 4

PUMPED WELL: SLF-50

DATE: 2/10/82 to 2/11/82

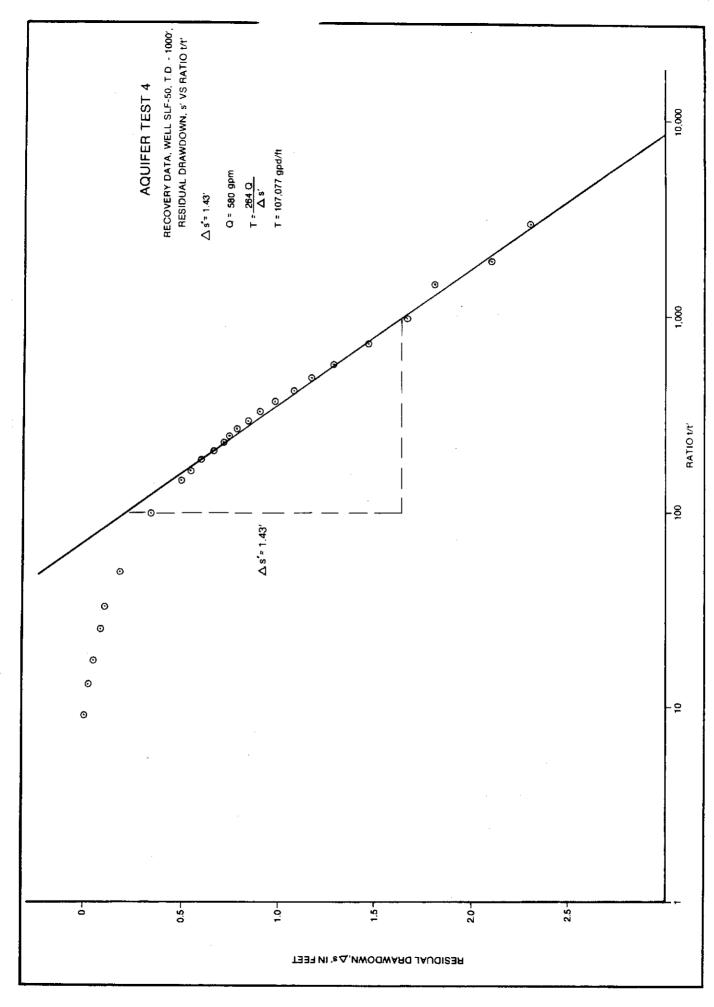
CASING DEPTH: 600 ft

WELL DEPTH: 1000 ft

WELL DISCHARGE: 580 gpm

RECOVERY DATA, WELL NO. SLF-50

TIME SINCE PUMPING STARTED t (MINS)	TIME SINCE PUMPING STOPPED t'(MINS)	RATIO t/t'	DEPTH TO WATER (FT ABOVE MP)	RESIDUAL DRAWDOWN s'(FT)
1457	0	-	-20.94	30.75
1457.25	0.25	5829	8.55	1.30
1457.5	0.50	2915	7.55	2.30
1457.75	0.75	1943.6	7.75	2.10
1458.0	1.00	1458	8.05	1.80
1458.5	1.50	972	8.19	1.66
1459.0	2.00	729.5	8.38	1.47
1459.5	2.50	583.8	8.56	1.29
1460	3 _	486.6	8.67	1.18
1460.5	3.5	417.3	8.76	1.09
1461	4	365.3	8.87	0.98
1461.5	4.5	324.8	8.94	0.91
1462	5 5.5	292.4	9.01	0.84
1462.5		265.9	9.06	0.79
1463	6	243.8	9.10	0.75
1464	/	109.1	9.17	0.68
1465	6 7 8 9	183.1	9.24	0.61
1466		162.9	9.30	0.55
1467	10	146.7	9.34	0.51
1472	15	98.1	9.50	0.35
1487	. 30	49.6 33.4	9.66 9.74	0.19 0.11
1501 1517	45 60	25.3	9.77	0.08
1517	90	17.2	9.8	0.05
1562	120	13.01	9.83	0.03
1622	180	9.01	9.85	0.00
1022	100	3.01	9.00	0.00



APPENDIX 4e. Aquifer Test No. 5 (8/24/82 to 8/26/82)

Table . Data on Wells Used in Pumping Test

WELL NUMBER	DEPTH (FT)	CASING DEPTH (FT)	CASING DIAMETER (INCHES)	DISTANCE FROM PUMPED WELL (FT)
SLF-49	893	56 0	6	428.34
SLF-50	775	600	6	147.90
SLF-51	775	600	6	-

PUMPING DATA, AQUIFER TEST NO. 5

ELAPSED TIME (MINS)	DRAWDOWN SLF-49 (FT)	DRAWDOWN SLF-50 (FT)	DRAWDOWN SLF-51 (FT)	PUMPING RATE (GAL/MIN)
0 0.25 .50	0 0 .01	0 0.19 0.56	0 41.11	0 393 455
.75	.01	0.84		413
1.00	.01	1.04	40.30	396
1.25 1.50	.01 .01	1.19 1.31		469 463
1.75	.01	1.42		450
2.00	.02	1.51	40.44	444
2.25	.02	1.59		437
2.50	.02	1.66	41.32	440
2.75	.02	1.73	10.00	432
3.00	.02	1.78	42.09	413
3.25 3.50	.02 .02	1.83	20.07	406 403
3.75	.02	1.86 1.90	39.97	403 388
4.00	.02	1.93	38.59	375
4.25	.02	1.96	30.33	369
4.50	.02	1.98		380
4.75	.03	1.99		375
5.00	.02	2.00	36.7	385
6.00	.025	2.03	34.84	377
7	.025	2.01	32.99	378
8	.03	2.01	31.33	369
9	.03	2.01	30.90	366
10	.03	2.01	30.79	355
11.00	.03	2.01	30.92	360
12	.03	2.02	30.94	358
13	.03	2.04	31.03	360
14	.03	2.05	31.08	360
15	.03	2.07	31.17	358 360
20 25	.03 .03	2.12 2.16	31.24 32.72	360 369
				302

PUMPING DATA, AQUIFER TEST NO. 5 (Continued)

ELAPSED	DRAWDOWN	DRAWDOWN	DRAWDOWN	PUMPING
TIME	SLF-49	SLF-50	SLF-51	RATE
(MINS)	(FT)	(FT)	(FT)	(GAL/MIN)
30 60 90 120 150 180 240 300 360 420 480 540 600 660 780 840 900 960 1020 1320 1380 1440 1500 1620 1680 1740 1860 1920 1980 2220 2280 2340 2460 2700 2760 2820 2880 3000 3060 3120 3180 3180 3120 3180 3180 3180 3180 3240 3300	.03 .035 .035 .035 .035 .02 .02 .01 .00 .01 .02 .04 .06 .07 .08 .08 .11 .11 .10 .09 .07 .04 .02 .00 .01 .02 .00 .01 .02 .04 .05 .07 .08 .08 .07 .08 .08 .07 .08 .09 .07 .09 .07 .08 .09 .07 .09 .07 .08 .07 .08 .07 .08 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .07 .09 .09 .09 .09 .09 .09 .09 .09 .09 .09	2.22 2.28 2.25 2.24 2.23 2.28 2.29 2.37 2.33 2.30 2.28 2.29 2.30 2.31 2.32 2.33 2.34 2.32 2.33 2.34 2.32 2.31 2.27 2.27 2.27 2.27 2.27 2.27 2.27 2.2	31.98 32.06 32.12 32.32 33.69 33.34 33.39 33.19 33.19 33.24 34.18 34.73 34.63 34.73 34.84 34.73 34.93 34.73 34.93 35.31 35.38 35.31 35.31 35.31 35.31 34.17 34.01 34.16 34.31	372 369 360 357 369 385 380 380 380 383 383 383 388 388 388 389 389 389 389

PUMPING DATA, AQUIFER TEST NO. 5 (Continued)

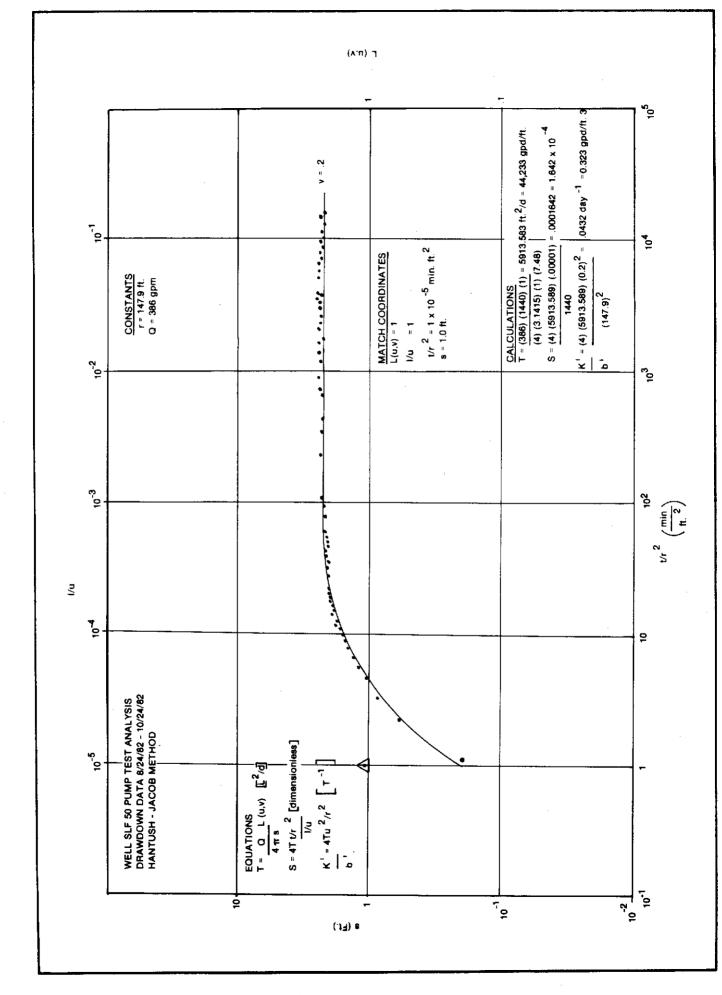
ELAPSED TIME (MINS)	DRAWDOWN SLF-49 (FT)	DRAWDOWN SLF-50 (FT)	DRAWDOWN SLF-51 (FT)	PUMPING RATE (GAL/MIN)
3360	.02	2.17	34.27	384
3420			34.21	
3480		•	34.34	
3540	.02	2.17	34.39	384
3600	.05	2.18	34.43	386
3660	.07	2,22	34.59	386
8720	.09	2.24		388
3780	.08	2.24		38 8
4020	.08	2.21	34.74	390
4080	.07	2.19	34.73	390
4140	.08	2.21	34.64	3 88
4200	.08	2.22	34.52	386
4260	.09	2.23	34.53	383
4320	.10	- ;	34.29	

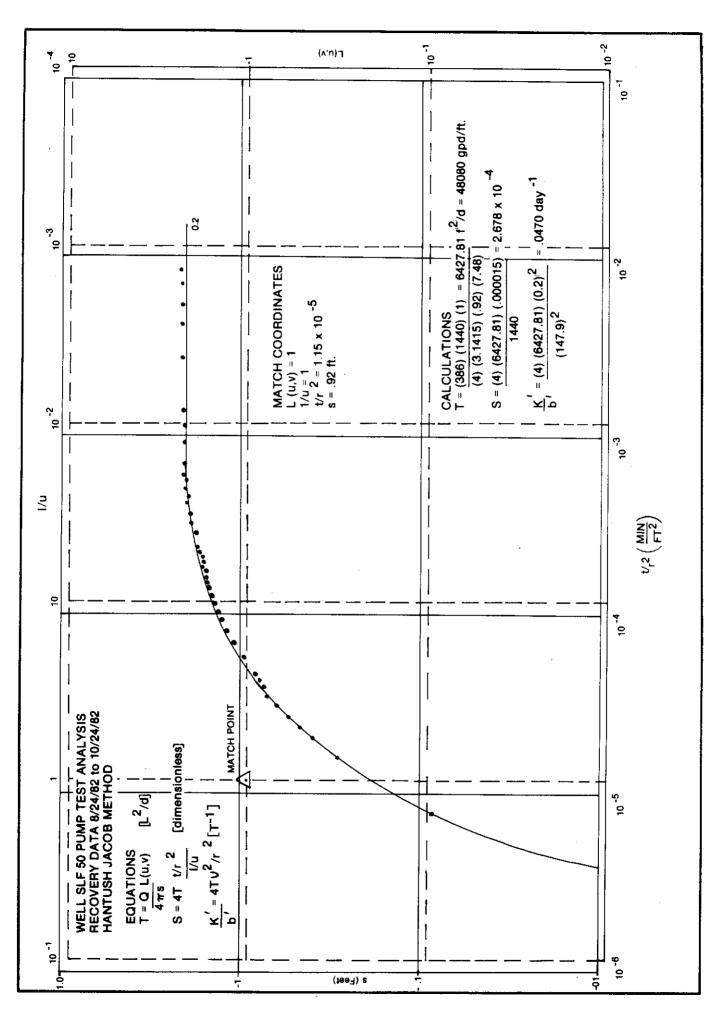
RECOVERY DATA, AQUIFER TEST NO. 5

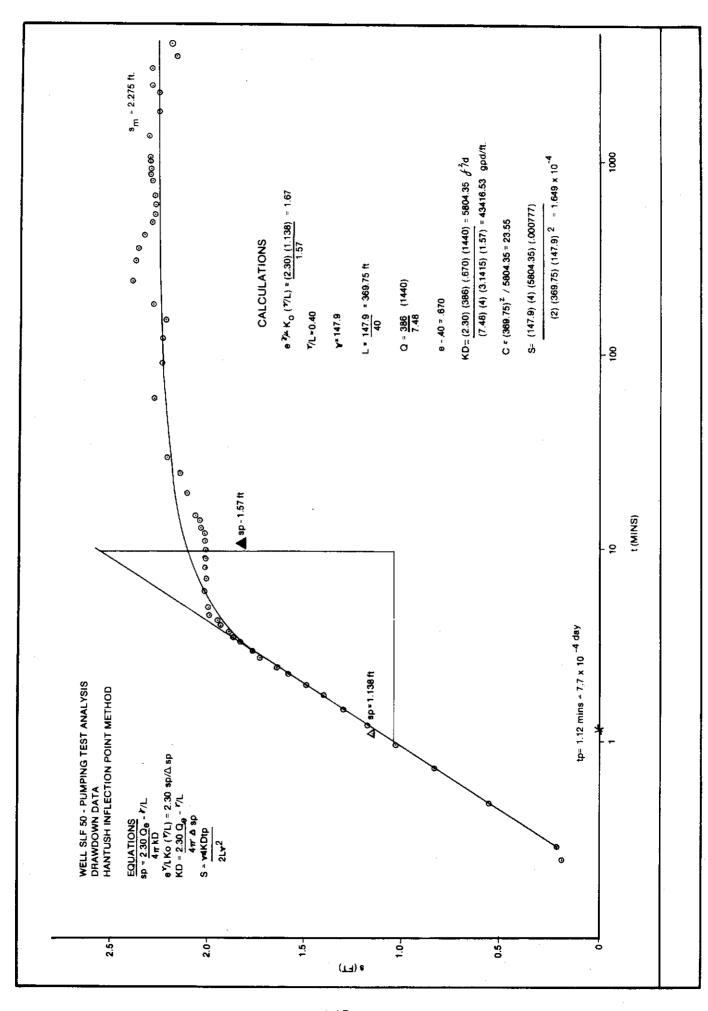
TIME SINCE PUMPING STARTED t (MINS)	TIME SINCE PUMPING STOPPED t'(MINS)	RATIO t/t'	RESIDUAL DRAWDOWN SLF-49 (FT)	RESIDUAL DRAWDOWN SLF-50 (FT)	RESIDUAL DRAWDOWN SLF-51 (FT)
4320.00 4320.08 4320.17	0 .08 .17	54001.0 25412.8	.10	2.22 2.14	0
4320.25 4320.33	.25 .33	17281.0 13091.9	.10	2.06 1.94	1.81
4320.42 4320.50 4320.58	.42 .50 .58	10286.7 8641.0 7449.3	.10	1.84 1.76 1.69	1.46
4320.67 4320.75 4320.83	.67 .75 .83	6448.8 5461.0 5205.8	.10	1.6 1.53 1.49	1.33
4320.92 4321.00 4321.25	.92 1.00 1.25	4696.7 4321.0 3457.0	.10	1.44 1.39 1.26	1.12 1.00
4321.50 4321.75 4322.00	1.50 1.75 2.00	2881.0 2469.6 2161.0	.09 .085 .085	1.14 1.05 .97	0.89 0.81 0.73
4322.25 4322.50 4322.75	2.25 2.50 2.75	1921.1 1729.0 1571.9	.085 .08 .08	.91 .85 .80	0.68 0.62 0.57
4323.00 4323.25 4323.50	3.00 3.25 3.50	1441.0 1330.2 1235.2	.08 .08 .08	.76 .72 .67	0.52 0.49 0.45
4323.75 4324.00	3.75 4.00	1153.0 1081.0	.08 .08	.65 .62	0.41 0.39

RECOVERY DATA, AQUIFER TEST NO. 5 (Continued)

TIME SINCE PUMPING . STARTED t (MINS)	TIME SINCE PUMPING STOPPED t'(MINS)	RATIO t/t'	RESIDUAL DRAWDOWN SLF-49 (FT)	RESIDUAL DRAWDOWN SLF-50 (FT)	RESIDUAL DRAWDOWN SLF-51 (FT)
4324.25 4324.75 4325.00 4326 4327 4328 4329 4330 4331 4332 4333 4334 4335 4340 4345 4350 4380 4410 4440 4470	4.25 4.50 4.75 5.00 6.00 7.00 8.00 9.00 10.00 11.00 12 13 14 15 20 25 30 60 90 120 150	1017.5 961.0 210.5 865.0 421.0 618.1 541.0 433.0 393.7 361.0 333.3 309.6 289.0 217 173.8 145.0 73.0 49.0 37.0 29.8	.08 .08 .075 .07 .07 .07 .065 .065 .06 .06 .06 .06 .055 .05	.59 .56 .54 .51 .44 .38 .33 .29 .26 .23 .22 .19 .18 .16 .10 .06 .03 04 06 08	0.36 0.33 0.31 0.28 0.20 0.14 0.09 .05 .01 02 05 07 09 11 16 20 22 28 30 30
4500	180	25.0	.02	10	30







APPENDIX 5

Summary of Packer Test, SLF-50

APPENDIX 5. Packer Test, Well SLF-50

Introduction

A packer test was performed at well SLF-50 on March 11 and 12, 1982 to provide additional information on the hydrologic properties of the Floridan Aquifer System. At the time of testing the well was 1000 feet deep, and was cased to a depth of 600 feet.

A packer is a mechanical device which is lowered into the wellbore and expanded at a given depth to provide a seal against the borehole wall. The purpose of the packer test was to:

- a) Isolate the potential injection horizon from the lower portion of the borehole.
- b) Determine the degree of interconnection between the potential injection horizon and the lower portion of the borehole.
- c) Determine differences in potentiometric head and water quality.
- d) Determine relative flow contributions from the two intervals under natural flowing conditions.

The setting depth of the packer was chosen based on hydrogeologic and borehole conditions. Geophysical, lithologic, and aquifer test data indicated that the interval between 600 feet and 775 feet was potentially the most productive zone in the sequence penetrated by the borehole. This interval consisted of medium-grained coquinoid limestones with well developed moldic and intergranular porosity. The Caliper Log of the well indicated considerable washout in this interval, probably due to poor cementing of the fossil fragments in the limestone matrix. The Flowmeter Log indicated three zones of substantial flow contribution in this interval. Pumping test data indicated significantly higher transmissivity values compared to the lower section of the strata penetrated.

Final choice of the setting depth was based on the borehole condition. Successful setting of the packer requires a relatively smooth section of borehole which has a diameter not more than 8 inches. The interval between 760 feet and 770 feet met these criteria. The 4 foot long expanding element was set between 766 feet and 770 feet depth below ground surface.

Equipment and Methods

The packer used was a 5 5/8 inch O.D. retrievable tool (TAM International Model Tamset) with a 4 foot long inflatable rubber element. The tool was lowered through the 6-inch PVC casing and into the open bore. Centralizers were placed on the drill stem close to the packer and also within the casing to hold the tool in the center of the hole and ensure smooth reentry into the casing on retrieval. At the setting depth, the tool was inflated and the circulating sleeve opened, using hydraulic pressure transmitted through the drill pipe. This effected a seal within the borehole and allowed water from below the packer to flow upward through the drill pipe. Groundwater from the strata isolated above the packer flowed through the annular space between the open bore or casing and the drill pipe. The tool is equipped with a deflating mechanism for retrieval.

Shut in of the drill stem and annular space were effected using a solid plug and a sandwich seal respectively. Shut-in heads were measured by means of manometer tubes tapped into the plug and seal. Manometer tubes were also installed on the monitoring well SLF-51 which was 775 feet deep and located 147.9 feet northeast of the exploratory well. Water samples and water quality data were taken after discharging for 2 hours and then allowing levels to stabilize. The two zones were allowed to discharge alternately and the effects on the other zone and on the observation wells were recorded. Table 5-1 summarzies data from these tests.

Static potentiometric level in the zone below the packer was 0.24 feet above the level in the upper zone, indicating some degree of isolation between these zones. This was further confirmed by data from the discharge tests. As shown on Table 5-1, discharging water from the lower zone caused no drawdown in the upper zone for a drawdown in the lower zone of 15 feet. The negligible effect observed at the monitor well could be due to measurement errors or natural fluctuations in water level, since this well was open to the upper zone only. Discharging water from the upper zone had a considerable effect on levels in the monitoring well (0.94 feet drawdown), but little effect on the water level in the lower zone (0.13 feet drawdown).

Water quality data were ambiguous but indicate that differences in quality between the zones probably exist. No consistent pattern in specific conductance, chloride, or total dissolved solids content was found which could be explained in terms of the separate zones isolated by the packer. Samples from well SLF-51 and the lower zone in well SLF-50 (depth 766 feet) showed the highest chloride concentrations. The annulus sample from the upper zone above 775 feet showed the lowest chloride concentration of 888 mg/l. The wellhead sample at SLF-50, after removal of the packer, had a chloride concentration intermediate between these. This tended to indicate that the upper zone was less mineralized than the lower zone. However, specific conductance and total dissolved solids concentrations were inconsistent with the chloride values. No measurable differences in temperature were observed. This may be due to temperature equalization during upward flow of water from the lower zone through the drill stem. However, temperature logs run prior to the packer test indicated an increase in temperature with depth from a value of 83.50F at 600 feet to approximately 85.80F at 1000 feet.

TABLE 5-1

WATER LEVEL AND WATER QUALITY DATA - PACKER TEST, WELL SLE-50

Wellhead - SLF-50 (after removal of packer)

Wellhead - Obs. Well SLF-51

82⁰

82⁰

WATER LEVEL AND WATER QUALITY DATA - PACKER TEST, WELL SLF-50													
ACTIVITY	WATER LEVEL OR DRAWDOWN IN UPPER ZONE (FT NGVD)	WATER LEVEL OR DRAWDOWN IN LOWER ZONE (FT NGVD)	WATER LEVEL OR DRAWDOWN SLF-51 (FT NGVD)										
All wells shut in (static condition).	42.86	43.10	42.34										
Lower zone pumped at 30 gallons per minute for 1 hour.	(0.00)*	(15.0)*	(0.06)*										
Upper zone discharged by free-flow at 200 gpm for 1 hour.	(10.60)*	(0.13)*	(0.94)*										
*NOTE: Numbers in bra	ckets indicate drawdow	n.											
WATER QUALITY DATA, PA	CKER TEST												
SAMPLING POINT	<u>т^оғ рн</u>	SPEC. COND. umhos/cm	C1 TDS mg/1 mg/1										
Drill stem - SLF-50 (lower zone)	82 ⁰ 7.63 7.55	3450 3200	1025 1474 1015 1888										
Annulus - SLF-50 (upper zone)	82 ⁰ 7.53	3400	888 2058										

7.45

6.17

3325

3500

3500

954.9

1077.8 1021.5 2379

2180

2143

In summary, the packer test confirmed that the interval above 766 feet produced the major flow to the borehole and that there was good lateral hydraulic connection within this interval between the exploratory well and the monitor well. This zone is separated from the lower strata by a low permeability confining layer. There is a small difference in heads between these two intervals. It would be expected that vertical leakage between the upper zone and lower zone would be small, compared to horizontal flow. Water quality differences between the two zones were relatively small and would therefore not be expected to have any significant effect on the injection/recovery process.

APPENDIX 6

Data from Injection/Recovery Tests

APPENDIX 6a. DATA FROM INJECTION PHASE OF INJECTION/RECOVERY TEST, WELL SLF-50, ST. LUCIE

REMARKS		Pump adjusted to 500 gpm at 1240.												.374 total gallons X 1000.							
OBS.WELL PRESSURE SLF-51 (ft TOC)	19.22	19.20	19.46	19,45	19.44	19.37	19,36	19.32	19.27	19.27	19.25	19.20	19.17	19.13	19.12	19.12	18.98	18.94	18.95	19.00	19.10
SPECIFIC CONDUCTANCE (umhos/cm)	1,220	1,220	1,230	1,221	1,210	1,230	1,240	1,272	1,260	1,260	1,250	1,290	1,330	1,320	1,320	1,320	1,210	1,320	1,310	1,300	1,310
INJECTION PRESSURE SLF-50 (psi)	28	29	31	32	33	33	33	33	34	34	35	34	34	34	34	34	34	35	35	36	35
Qin (gpm)	475	475	200	200	200	200	200	200	200	200	200	200	200	200	200	200	475	475	475	475	475
ELAPSED TIME (min)	09	120	180	240	300	360	420	480	540	540	009	099	720	780	840	006	1260	1320	1380	1440	1500
TIME	1114	1214	1314	1414	1514	1614	1714	1814	1914	1914	2014	2114	2214	2314	2414	0114	0714	0814	0914	1014	1114
DATE	10/19/82	10/19/82	10/19/83	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/19/82	10/20/82	10/20/82	10/20/82	10/20/82	10/20/82	10/20/82

APPENDIX 6a (Continued)

REMARKS						Pump adjusted to reduce injection pressure.			Pump adjusted to reduce injection pressure.						Pump adjusted.						
OBS.WELL PRESSURE SLF-51 (ft TOC)	19.16	19.12	19.12	19.01	18.96	18.68	18.53	18.48	17.79	•	18.00	17.98	18.00	17.99	17.80	18.01	18.00	17.98	17.98	17.95	18.01
SPECIFIC CONDUCTANCE (umhos/cm)	1,310	1,310	1,310	1,285	1,310	1,310	1,320	1,330	1,300	1,300	1,300	1,350	1,350	1,350	1,330	1,330	1,330	1,320	1,320	1,300	1,310
INJECTION PRESSURE SLF-50 (psi)	37	39	38	38	40	32	37	36	28	34	34	34	34	34	30	38	38	39	40	40	39
Qin (gpm)	475	475	475	450	450	400	375	350	280	325	325	325	325	325	300	325	300	300	300	300	300
ELAPSED TIME (min)	1560	1620	1680	1800	1860	1920	1980	2040	2160	2180	2220	2280	2340	2400	2700	2760	2820	2880	2940	3000	3060
TIME	1214	1314	1414	1614	1714	1814	1914	2014	2214	2234	2314	2414	0114	0214	0714	0814	0914	1014	1114	1214	1314
DATE	10/20/82	10/20/82	10/20/82	10/20/82	10/20/82	10/20/82	10/20/81	10/20/82	10/20/82	10/20/82	10/20/82	10/20/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82

APPENDIX 6a (Continued)

REMARKS			Water sample taken.		Total gallons injected 1,279,300.	Recovery halted to clean well by backflow. Backflow started at 1851.	Water murky after 10 min. T=760F.	Backflow completed.	Injection restarted.											Injection halted to redevelop well. Total gallons injected 1,476,200.	Backflow started.
OBS.WELL PRESSURE SLF-51 (ft TOC)	17.97	17.94	17.92	17.88	I					17.87	17.99	17.96	17.90	17.88	17.89	17.58	17.59	17.72	17.73		
SPECIFIC CONDUCTANCE (umhos/cm)	1,310	1,270	1,240	1,270			1,330			1,270	1,320	1,310	1,330	1,320	1,340	1,330	1,330	1,320	1,310		
INJECTION PRESSURE SLF-50 (psi)	40	40	41	40	41					38	38	36	37	38	34	34	38	39	39		
Qin (gpm)	300	300	275	280	275					325	280	300	300	280	250	250	250	250	250		
ELAPSED TIME (min)	3120	3240	3300	3360	3390	3392	3397	3426	3432	3436	3600	3660	3720	3780	3840	4140	4200	4260	4320	4324	4326
TIME	1414	1614	1714	1814	1844	1846	1851	1920	1926	1920	2214	2314	2414	0114	0214	0714	0814	0914	1014	1018	1020
DATE	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/21/82	10/22/82	10/22/82	10/22/82	10/22/82	10/22/82	10/22/82	10/22/82	10/22/83	10/22/83

APPENDIX 6a (Continued)

REMARKS	Backflow stopped.	Surging with pump started. Water dark brown at start, clear after 15 min.	End of surging.	Pumping resumed.		Injection terminated. Total gallons injected 1,488,100.
OBS.WELL PRESSURE SLF-51 (ft TOC)				•	17.45	18.06
SPECIFIC CONDUCTANCE (umhos/cm)			1,290		1,210	1,250
INJECTION PRESSURE SLF-50 (psi)					35	34
Qin (gpm)					300	300
ELAPSED TIME (min)	4335	4345	4386	4397	4398	4500
TIME	1029	1039	1120	1131	1132	0114
DATE	10/22/83	10/22/83	10/22/83	10/22/83	10/22/83	10/22/83

APPENDIX 6b. Data from Recovery Test, Well SLF-50, St. Lucie County

DATE	TIME	VOLUME DISCH. GALS.	SPECIFIC CONDUCT. umhos/cm	CHLORIDE CONCEN. mg/l	TEMP.	REMARKS
11/29/82 11/29/82 11/29/82 11/29/82 11/29/82 11/29/82 11/29/82	1150 1153 1220 1250 1350 1450 1550	0 6100 10000 21200 31400 40800	1332.0 1372.0 1372.0 1372.0 1372.0 1386.0	- 220 230 - - - 250	78.8 78.8 78.8 78.8 78.8 78.8	Water Sample #1.
11/29/82 11/29/82 11/29/82 11/29/82	1620 1650 1720 1750	45700 51100 56200 60600	1402.0 1414.0 1414.0 1420.0	262 - - 282	78.8 78.8 78.8 78.8	Water Sample #2. Water Sample #3;
11/30/82 11/30/82 11/30/82 11/30/82 11/30/82	0703 0733 0803 0903 1003	62000 16100 72200 82500 92900	1430.0 1421.0 1421.0 1420.0 1416.0	277 271 - -	80.6 78.8 78.8 78.8 79.0	recovery halted. Recovery resumed. Water Sample #4. Water Sample #5.
11/30/82 11/30/82 11/30/82 11/30/82	1103 1203 1303 1403	103200 113600 123900 135300	1424.0 1434.0 1438.0 1450.0	278 - - 288	79.0 79.0 79.0 79.0	Water Sample #6. Water Sample #7;
11/30/82 11/30/82	1503 1603	144800 155100	1449.0 1461.0	-	79.1 79.1	water clear. Water Sample #8; recovery halted.
12/01/82 12/01/82 12/01/82	0705 0735 0805	167400 171700 177000	1434.0 1490.0 1497.0	285 298 -	79.1 79.1 79.1	Recovery resumed. Water Sample #9; water clear.
12/01/82 12/01/82 12/01/82	0905 1005 1105	187700 198300 209000	1497.0 1497.0 1503.0	308	79.2 79.2 79.2	Water Sample #10; water clear.
12/01/82 12/01/82 12/01/82 12/02/82	1505 1605 1645 0700	251600 262200 270100 270200	1522.0 1529.0 1543.0	- 322 -	79.4 79.4 79.4 79.4	Recovery halted. Recovery resumed.
12/02/82 12/02/82 12/02/82 12/02/82 12/02/82	0705 0735 0805 0905 1005	271100 277500 281800 292700 303600	1553.0 1553.0 1552.0 1547.0 1548.0	328 328 - - -	79.2 79.2 79.2 79.2	Water Sample #12.
12/02/82 12/02/82 12/02/82 12/02/82	1105 1205 1305 1405	315500 325800 336600 348500	1554.0 1556.0 1556.0 1565.0	- - - -	79.2 79.4 79.4 79.4	Water Sample #13. Water Sample #14.

APPENDIX 6b. Data from Recovery Test, Well SLF-50, St. Lucie County (Continued)

DATE	TIME	VOLUME DISCH. GALS.	SPECIFIC CONDUCT. umhos/cm	CHLORIDE CONCEN. mg/l	TEMP.	REMARKS
12/02/82 12/02/82 12/02/82	1505 1605 1705	358700 370400 381200	1565.0 1574.0 1595.0	- - -	79.4 79.4 79.4	Water Sample #15; recovery halted.
12/03/82	0700	381300	1605.0		- 77 0	Recovery resumed.
12/03/82 12/03/82 12/03/82 12/03/82	0705 0735 0805 0905	382200 387500 392000 402900	1605.0 1605.0 1605.0 1606.0	- - -	77.2 79.4 79.4 79.4	Water Sample #16.
12/03/82 12/02/82 12/03/82	1005 1105 1205	413700 425500 436400	1606.0 1606.0 1612.0	- - -	79.6 79.6 79.6	Water Sample #17.
12/03/82 12/03/82 12/03/82 12/03/82	1305 1405 1505 1605	447300 458100 468000 478900	1612.0 1613.0 1625.0 1634.0	- - -	79.6 79.6 79.6 79.6	Water Sample #18.
12/03/82	1705	489800	1644.0	-	79.6	Water Sample #19; recovery halted.
12/06/82 12/06/82	1015 1020	489900 490800	1624.0	- 375	- 79.0	Recovery resumed.
12/06/82 12/06/82 12/06/82	1030 1100 1200	492900 498800 510000	1556.0 1618.0 1615.0	351 382 -	79.4 79.5 79.6 79.6	Water Sample #20.
12/06/82 12/06/82 12/06/82 12/06/82	1300 1400 1500 1600	521800 533100 544900 557400	1620.0 1638.0 1638.0 1648.0	389	79.6 79.6 79.6	Water Sample #21.
12/06/82	1700 0700	569000 569200	1650.0	391 -	79.6	Water Sample #22; recovery halted. Recovery resumed.
12/07/82 12/07/82 12/07/82 12/07/82	0705 0735 0805 0905 1005	570100 575900 581600 593000 605500	1667.0 1668.0 1668.0 1665.0 1672.0	396 398 - -	79.3 79.8 79.8 79.9 80.0	Water Sample #23.
12/07/82 12/07/82 12/07/82 12/07/82	1105 1205 1305	616000 627400 638900	1676.0 1676.0 1679.0	407 -	80.0 80.0 80.0	Water Sample #24.
12/07/82 12/07/82 12/07/82 12/07/82	1405 1505 1605	651400 661900 674400	1679.0 1678.0 1688.0	415 - -	80.0 80.0 80.0	Water Sample #25.
12/07/82	1705	684900	1702.0	420	80.0	Water Sample #26; recovery halted.
12/08/82 12/08/82	0700 0705	685000 686200	- 1714.0	-	80.0	Recovery resumed.
12/08/82	0735	691500	1715.0	426	80.0	Water Sample #27.

APPENDIX 6b. Data from Recovery Test, Well SLF-50, St. Lucie County (Continued)

DATE	TIME	VOLUME DISCH. GALS.	SPECIFIC CONDUCT. umhos/cm	CHLORIDE CONCEN. mg/l	TEMP.	REMARKS
12/08/82 12/08/82	0805 0 9 05	694700 703000	1714.0 1708.0	-	80.0 80.0	
12/08/82 12/08/82	1005 1105	711200 719700	1696.0 1707.0	- 422	80.0 80.1	Water Sample #28.
12/08/82 12/08/82	1205 1305	728100 736600	1704.0 1704.0	- -	80.1 80.1	
12/08/82	1405	745100	1712.0	424	80.1	Water Sample #29.
12/08/82 12/08/82	1505 1605	753600 762100	1713.0 1731.0	- -	80.1 80.1	
12/08/82	1705	770700	1740.0	436	80.1	Water Sample #30; recovery halted.
12/09/82	0700	770800	1740.0	-	-	Recovery resumed.
12/09/82	0705 0735	771800 777400	1743.0	- 127	80.4	Water Sample #31.
12/09/82 12/09/82	0805	777400 782900	1746.0 1749.0	437 -	80.1 80.1	•
12/09/82	0905	793900	1750.0	- .	80.1	
12/09/82	1005	804700	1751.0	-	80.2	
12/09/82	1105	815200	1757.0	440	80.2	Water Sample #32.
12/09/82	1205	825300	1758.0	-	80.2	
12/09/82	1305	835900	1760.0	- 420	80.2	Hatau Camala #22
12/09/82 12/09/82	1405 1505	846100 856100	1750.0 1749.0	438	80.4 80.2	Water Sample #33.
12/09/82	1605	867400	1727.0	-	80.4	
12/09/82	1705	875800	1768.0	_	80.1	
12/09/82	1805	886400	1777.0	435	80.1	Water Sample #34;
12/10/02	0707	000600	1705 0		70.7	recovery halted.
12/10/82 12/10/82	0707 0808	888600 896400	1725.0 1765.0	- -	79 .7 78 . 8	Recovery resumed.
12/10/82	0908	906800	1770.0	-	80.6	
12/10/82	1008	915900	1770.0	-	80.6	Water Sample #35.
12/10/82	1108	926300	1763.0	-	80.6	
12/10/82	1206	935700	1766.0	-	80.6	
12/10/82	1306	945500	1768.0	-	80.6	
12/10/82	1408	956000	1786.0	-	80.6	Unhar Carry la #26
12/10/82	1506	965900	1770.0	•	80.6	Water Sample #36;
12/13/82	1020	965900	1740.0	448	79.8	recovery halted. Water Sample #37;
12, 13, 02	1020	303300	1740.0	440	75.0	recovery resumed.
12/13/82	1100	973500	1763.0	_	80.4	, 400, 0.9
12/13/82	1200	984400	1776.0	-	80.4	
12/13/82	1300	996100	1768.0	469	80.4	Water Sample #38.
12/13/82	1400	1005500	1848.0	-	80.4	
12/13/82	1500	1014300	1846.0	467	80.4	Water Cample #20
12/13/82 12/13/82	1600 1700	1023100 1032100	1798.0 1854.0	467 -	80.3 80.3	Water Sample #39.
12/13/82	1800	1040300	1822.0	<u>-</u>	80.4	Water Sample #40;
,v, v=		20.000				recovery halted.

APPENDIX 6b. Data from Recovery Test, Well SLF-50, St. Lucie County (Continued)

DATE	TIME	VOLUME DISCH. GALS.	SPECIFIC CONDUCT. umhos/cm	CHLORIDE CONCEN. mg/l	TEMP.	REMARKS
12/15/82	0710	1149800	1871.0	498	80.4	Water Sample #45;
12/15/82 12/15/82	0810 0910	1157400 1166300	1842.0 1828.0	-	80.6 80.6	recovery resumed.
12/15/82 12/15/82	1010 1110	1175900 1185000	1818.0 1824.0	493 -	81.0 81.0	Water Sample #46.
12/15/82 12/15/82 12/15/82	1210 1310 1410	1195400 1205100 1212900	1817.0 1817.0 1811.0	494 -	81.0 81.0 81.0	Water Sample #47.
12/15/82 12/15/82	1510 1610	1223300 1231600	1804.0 1812.0	· <u>-</u>	81.0 81.0	
12/15/82	1710	1240900	1815.0	500	81.0	Water Sample #48; recovery halted.
12/16/82	0700	1242100	1837.0	518	81.0	Water Sample #49; recovery resumed.
12/16/82 12/16/82	0800 0900	1254500 1266500	1839.0 1827.0	-	81.0 81.0	•
12/16/82 12/16/82 12/16/82	1000 1100 1200	1278300 1289200 1300900	1825.0 1825.0 1825.0	514 - -	81.0 81.0 81.0	Water Sample #50.
12/16/82 12/16/82 12/16/82	1300 1400 1500	1312700 1325500 1337300	1827.0 1825.0 1825.0	507 - -	81.0 81.0 81.0	Water Sample #51.
12/16/82 12/16/82	1600 1700	1348100 1359900	1838.0 1843.0	- 521	81.0 81.0	Water Sample #52;
12/17/82	0700	1363400	1849.0	520	81.0	recovery halted. Water Sample #53; recovery resumed.
12/17/82 12/17/82	0800 0900	1371600 1381500	1849.0 1847.0	<u>-</u> -	81.0 81.0	•
12/17/82	1000	1390500	1842.0	532	81.0	Water Sample #54; recovery halted.
12/21/82 12/21/82 12/21/82 12/21/82	1000 1015 1115 1215	1390500 1393000 1405000 1415600	1623.0 1723.0 1720.0	435 -	81.6 81.2 81.2	Recovery resumed. Water Sample #55.
12/21/82 12/21/82 12/21/82	1315 1415 1515	1427200 1438700 1450300	1715.0 1715.0 1715.0 1700.0	433 - -	81.4 81.4 81.4	Water Sample #56.
12/21/82 12/21/82 12/22/82	1615 1715 0630	1461900 1474500 1474500	1713.0 1719.0	527 - -	81.2 81.2	Water Sample #57. Recovery halted. Recovery resumed.
12/22/82 12/22/82 12/22/82	0700 0800 0900	1478800 1490300 1499900	1740.0 1720.0 1714.0	546 - 546	81.0 81.0 81.2	Water Sample #58. Water Sample #59.
12/22/82	1000	1511400	1715.0	-	81.4	

APPENDIX 6b. Data from Recovery Test, Well SLF-50, St. Lucie County (Continued)

DATE	TIME	VOLUME DISCH. GALS.	SPECIFIC CONDUCT. umhos/cm	CHLORIDE CONCEN. mg/l	TEMP.	REMARKS
12/22/82	1030	1515700	1702.0	555	81.4	Water Sample #60; recovery halted.
1/19/83 1/19/83 1/19/83	1000 1100 1400	1515700 1528900 1562600	- 1688.0 1817.0	- 472 549	- 81.4 81.4	Recovery resumed. Water Sample #61.
1/19/83 1/20/83 1/20/83	1700 0700 1000	1596800 1757600 1792200	1824.0 1924.0 1904.0	551 602 593	81.4 81.5 81.5	Water Sample #62. Water Sample #63.
1/20/83 1/20/83 1/20/83 1/21/83	1300 1700 0700	1826700 1872800	1888.0 1912.0 1922.0	615 645 669	81.5 81.5 81.8	Water Sample #64. Water Sample #65.
1/21/83 1/21/83	1000 1300	2034800 2069600 2105400	1932.0 1943.0	670 686	81.8 82.0	·
1/21/83 1/24/83	1600 1030	2139300 2139300	1960.0 -	700 -	82.0 -	Water Sample #66 recovery halted. Recovery resumed.
1/24/83 1/24/83 1/24/83	1100 1400 1700	2142800 2169200 2196800	1892.0 1940.0 1962.0	668 684 701	81.7 82.0 82.0	Water Sample #67. Water Sample #68.
1/25/83 1/25/83 1/25/83	0700 1000 1300	2318400 2343400 2368400	1995.0 2028.0 2032.0	717 726 732	82.0 81.6 81.7	Water Sample #69.
1/25/83 1/26/83 1/26/83	1700 0700 1000	2402000 2519000 2544200	2032.0 2087.0 2067.0	732 748 745	81.7 81.6 81.7	Water Sample #70. Water Sample #71.
1/26/83 1/26/83 1/27/83	1300 1700 0700	2569400 2604500 2720900	2076.0 1096.0 2113.0	756 759 778	81.9 81.9 81.8	Water Sample #72. Water Sample #73.
1/27/83 1/27/83	1000 1300	2746100 2771800	2103.0 2122.0	763 782	82.0 82.0	
1/27/83 2/01/83	1630 1100	2802500 2802500	2140.0 2210.0	791 706	82.0 81.9	Water Sample #74; recovery halted. Water Sample #75;
2/01/83 2/01/83	1400 1700	2827100 2852500	2187.0 2218.0	797 789	81.7 81.1	recovery resumed. Water Sample #76.
2/02/83 2/02/83 2/02/83	0700 1100 1400	2971500 3006800 3033600	2198.0 2220.0 2210.0	805 797 808	82.0 82.2 81.9	Water Sample #77.
2/02/83 2/03/83 2/03/83	1700 0700 1100	3060700 3191200 3227800	2203.0 2248.0 2218.0	815 818 823	82.4 82.4 82.4	Water Sample #78. Water Sample #79.
2/03/83 2/03/83 2/04/83	1400 1700 0730	3253900 3280000 3411500	2202.0 2231.0 2207.0	832 840 848	82.0 82.0 82.4	Water Sample #80. Water Sample #81.